

# MICROWAVE RADAR METEOROLOGY Foundations and Applications

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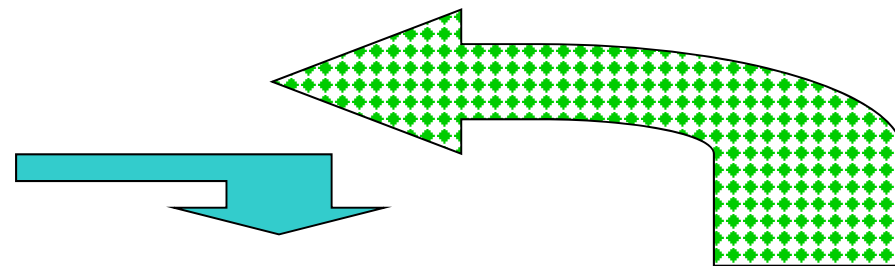
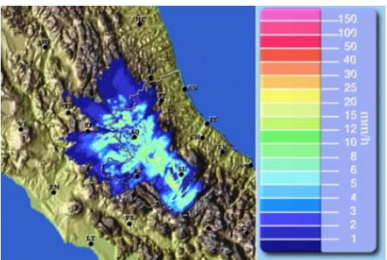
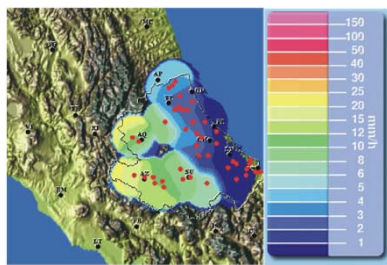
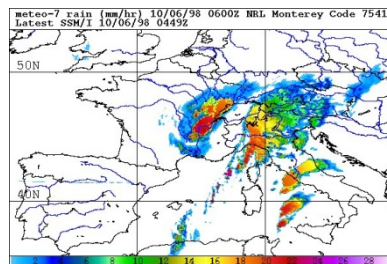
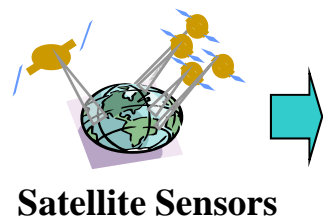
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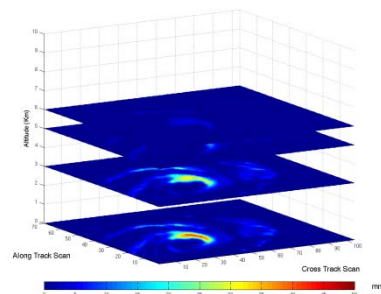


# RAINFALL FORECAST

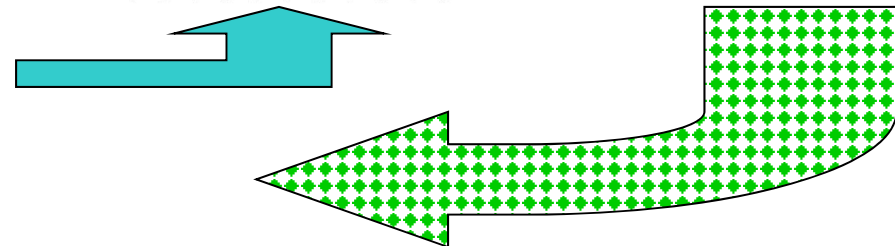
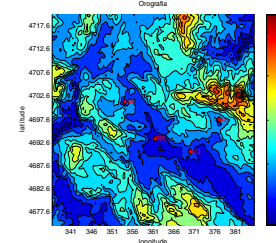
## A conceptual diagram



Atmospheric Forecast

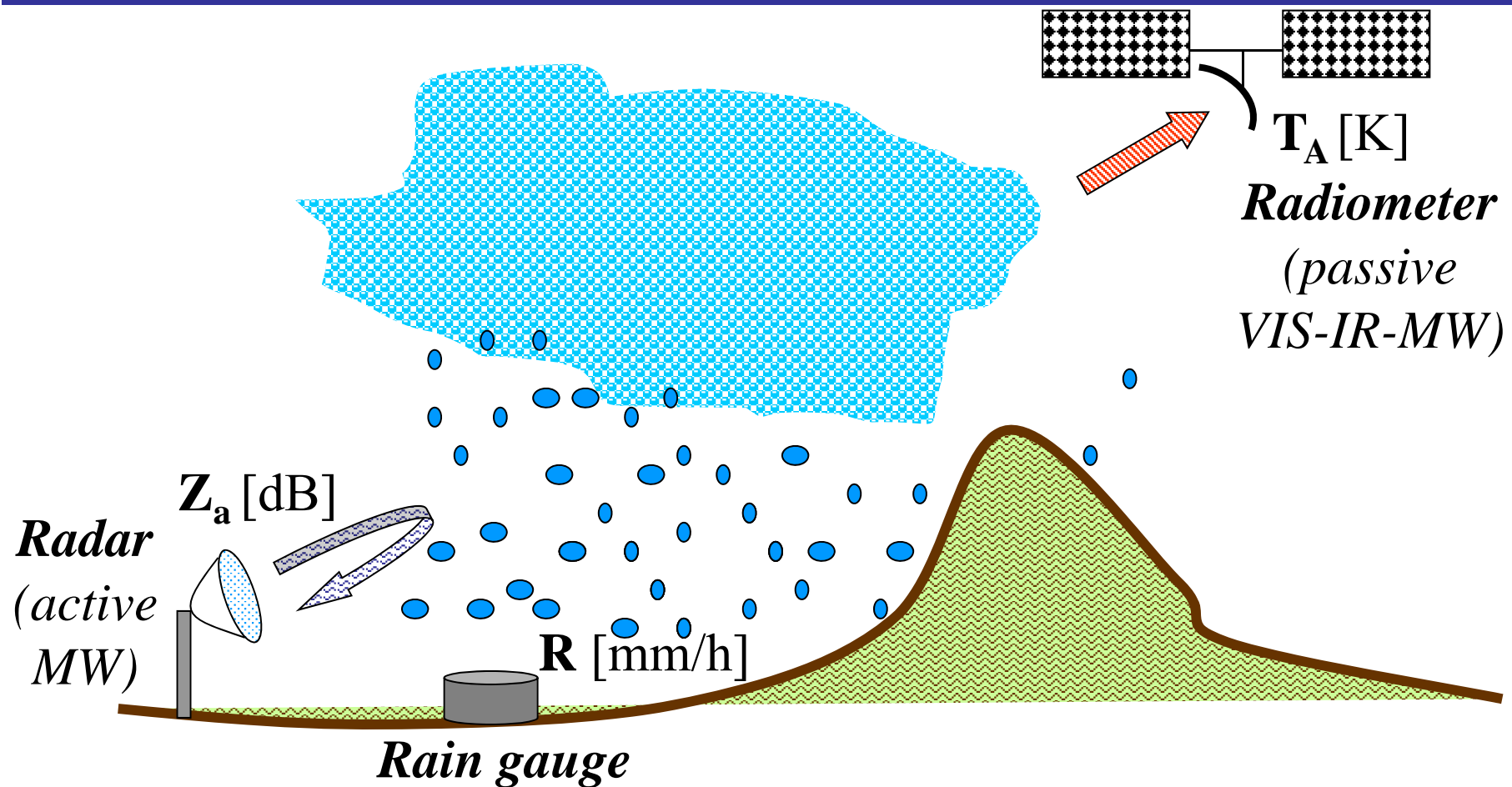


Hydrogeological Forecast



# RAINFALL SENSING

## Points of view



# RADAR METEOROLOGY

## Lecture's contents

- **Radar sensor**
  - Pulsed, Doppler, and polarimetric systems
  - Receiver sensitivity, antenna specifications and radar volume resolution
- **Radar equation**
  - Atmospheric refraction and attenuation
  - Radar equation for single and distributed scatterers
- **Radar signal**
  - Signal statistics and decorrelation
  - Noise reduction techniques
- **Radar applications**
  - Clouds and precipitation
  - Rainfall backscattering and polarimetric measurables
- **Radar products**
  - Radar measurements and error budget
  - Examples of radar measurements and estimates

# RADAR METEOROLOGY

## Historical perspective

### ➤ Radio era

- 1886: H. Hertz experiments e.m. wave reflection
- 1902: G. Marconi carries out radio-propagation experiments
- 1922: object detection with continuous wave instruments
- 1932: object detection with pulsed wave instruments

### ➤ Radar era

- 1938: construction of radars for aircraft detection and ranging
- 1941: first use of radars for cloud observations
- 1960s: use of radars for quantitative rainfall estimation
- 1970s: use of Doppler radars at VHF-UHF for turbulence and wind

### ➤ Computer era

- 1980s: impact of computers on data acquisition and processing
- 1990s: use of digital receivers and pulse forming
- 1990s: use of radars for weather nowcasting
- 2000s: spaceborne radars and sensors' synergy

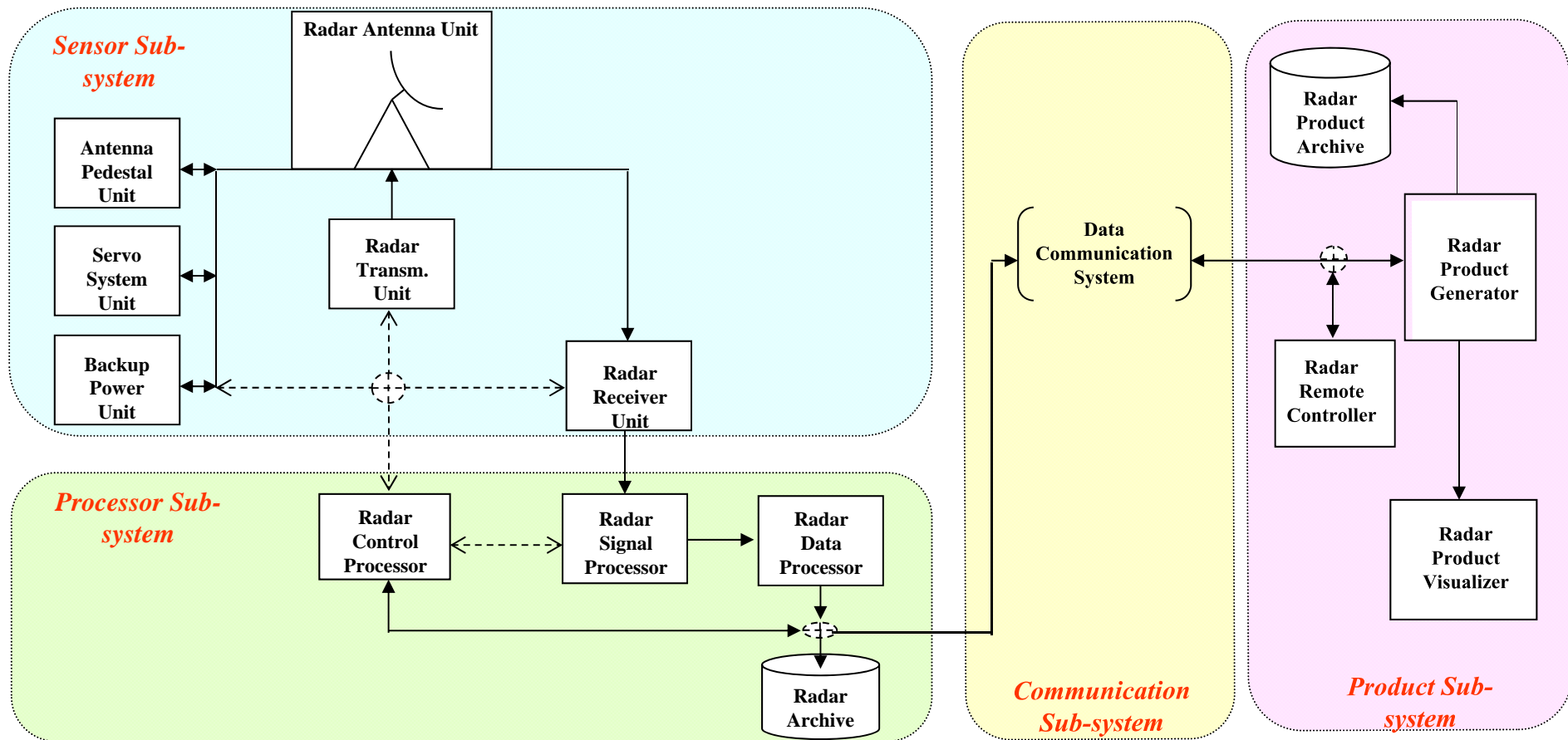
# RADAR METEOROLOGY

## Context and objectives

- **Radar (RAdio Detection and Ranging)**
  - Pulsed incoherent and coherent radars (Doppler or MTI)
  - Continuous-wave and frequency modulation (FM-CW)
- **Meteorological radars**
  - Rain (weather) radars: rainfall monitoring from ground and space
  - Cloud radars: cloud monitoring from ground and space
  - Stratosphere-Troposphere (ST) radars: wind profiling from ground
- **Weather (rain) radars**
  - Monitoring of three-dimensional (3-D) structure of rainfall and winds
    - Covering large areas (100-400 km) around ground site
    - Measuring of e.m. backscattering due to cloud hydrometeor volumes
  - Advantages with respect to:
    - optical sensors (e.g, higher penetration, any meteo condition)
    - radiometers (e.g., ranging capability, )
    - rain-gauges (e.g., larger effective coverage, rain structure)
  - Disadvantages: high power, complexity, maintenance

# RADAR METEOROLOGY

## A system approach



# RADAR METEOROLOGY

## Lecture's contents

### ➤ Radar sensor

- Pulsed, Doppler, and polarimetric systems
- Receiver sensitivity
- Antenna specifications
- Radar volume resolution

### ➤ Radar equation

- Atmospheric refraction and attenuation
- Radar equation for single and distributed scatterers

### ➤ Radar signal

- Signal statistics and decorrelation
- Noise reduction techniques

### ➤ Radar applications

- Operational problem overview
- Clouds and precipitation
- Rainfall backscattering and polarimetric measurables
- Example of radar measurements and estimates



# RADAR SENSOR

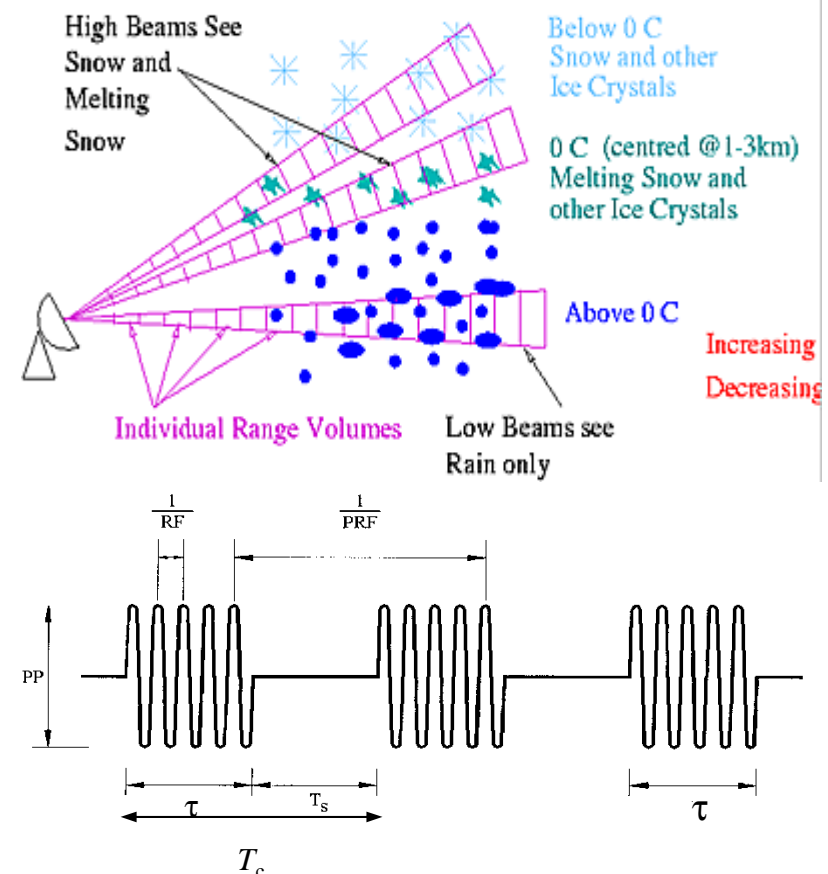
## Pulsed microwave systems

### ➤ Principle

- Send a train of short, **high-power pulse** of e.m. energy at high frequency (GHz)
- E.m. energy, captured by atmospheric objects, is **partially absorbed and re-irradiated** (scattered) into many directions among which that of radar (backscattered: radar echo)
- **Angular resolution**: derived from the antenna pointing direction and limited by its beam width
- **Range resolution**: derived from the two-way time employed by radar pulse echo (antenna-object-antenna) and depending by medium light velocity

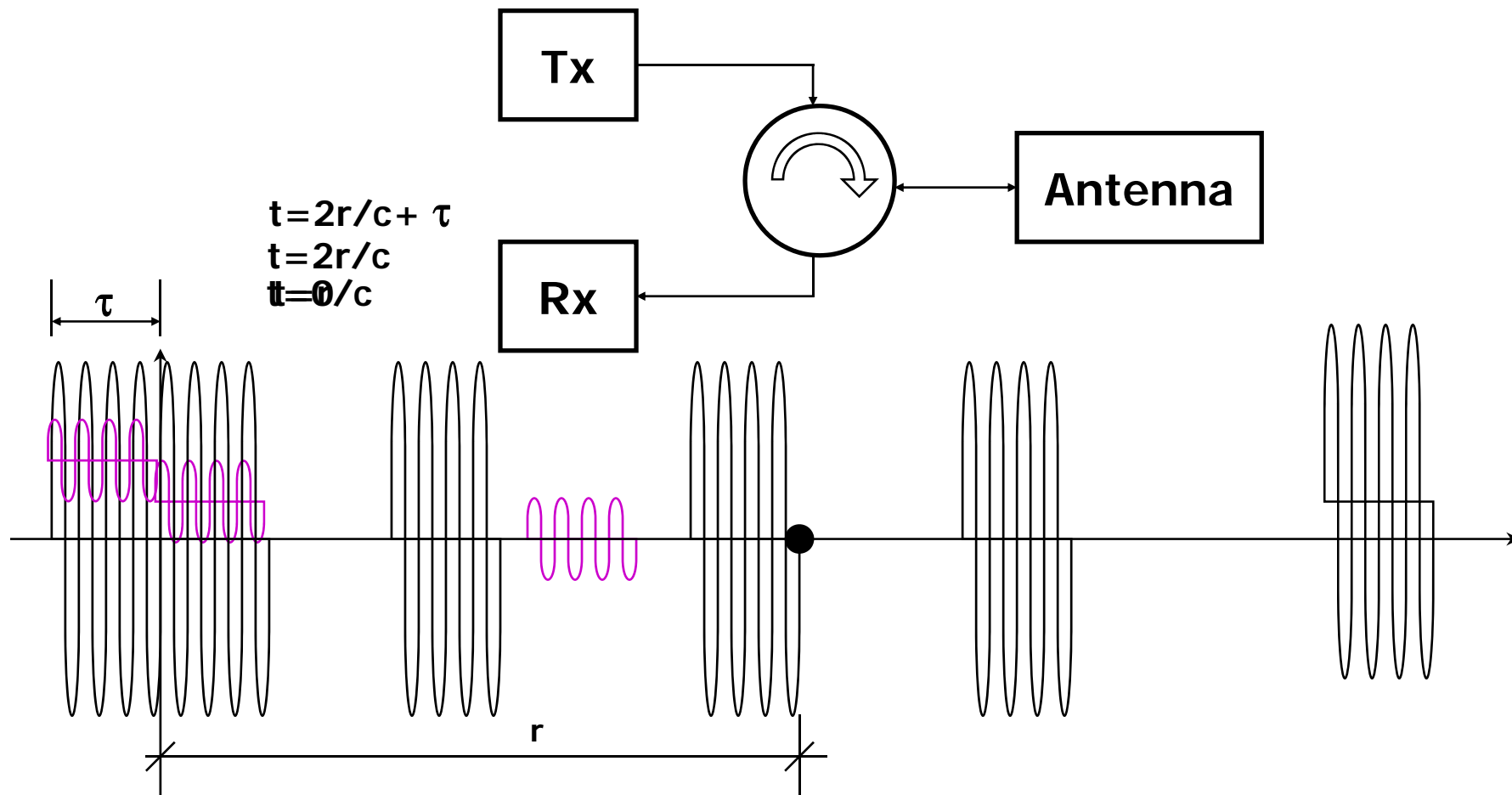
### ➤ Legend

- PRF: Pulse Repetition Frequency
- RF: Radio Frequency ( $f_0$ )
- PP: Peak-to-peak power
- $P_t$ : Peak power of the transmitter



# RADAR SENSOR

## Principle of pulsed systems



# RADAR SENSOR

## Principle of operation

### Range resolution

$$\Delta r = \frac{c\tau}{2}$$

### Maximum unamb. range

$$r_{\max} = \frac{cT_c}{2}$$

### Repetition frequency

$$f_r = \frac{1}{T_c}$$

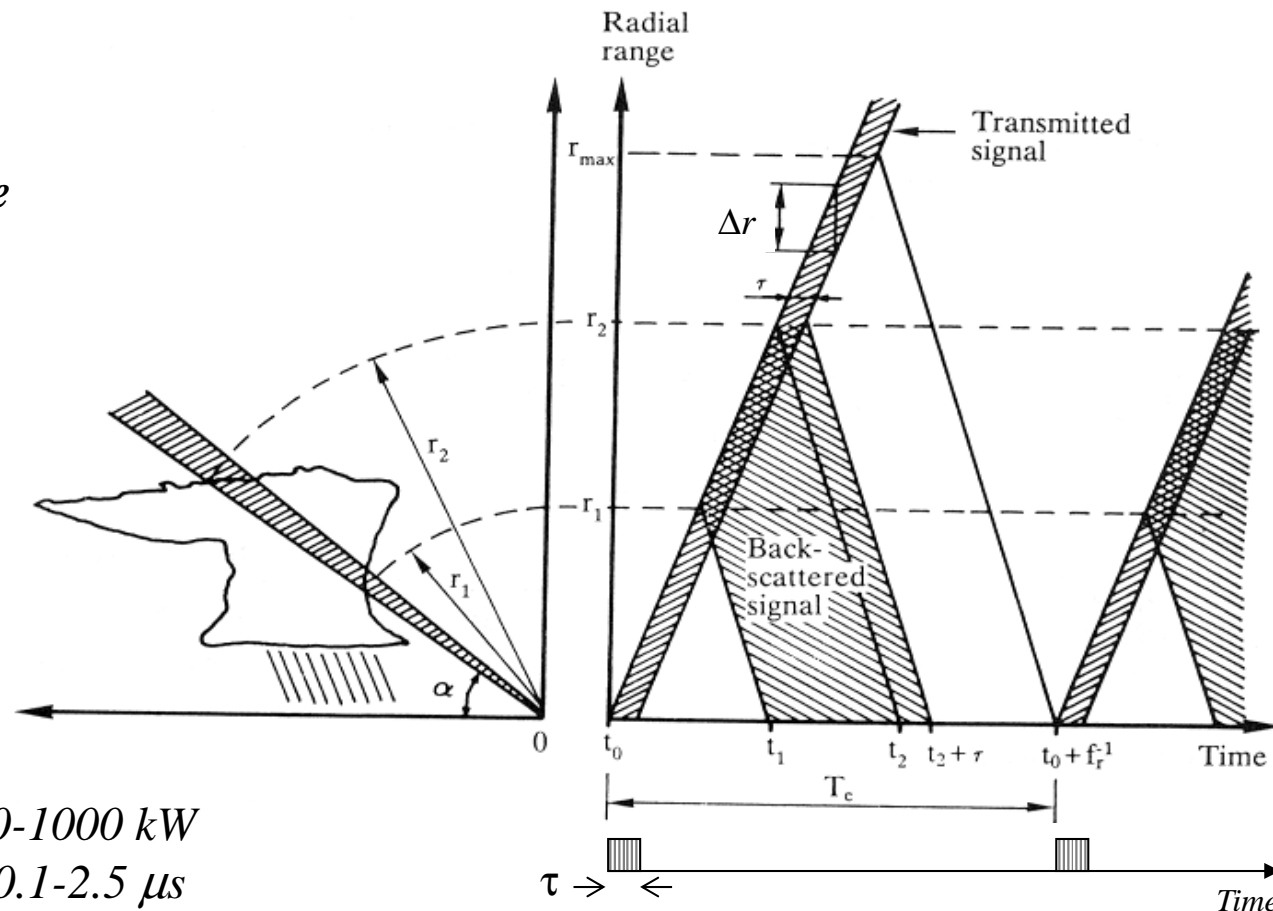
### Average TX Power

$$P_{\text{avg}} = P_t \left( \frac{\tau}{T_c} \right) = P_t d_c$$

### Typical values:

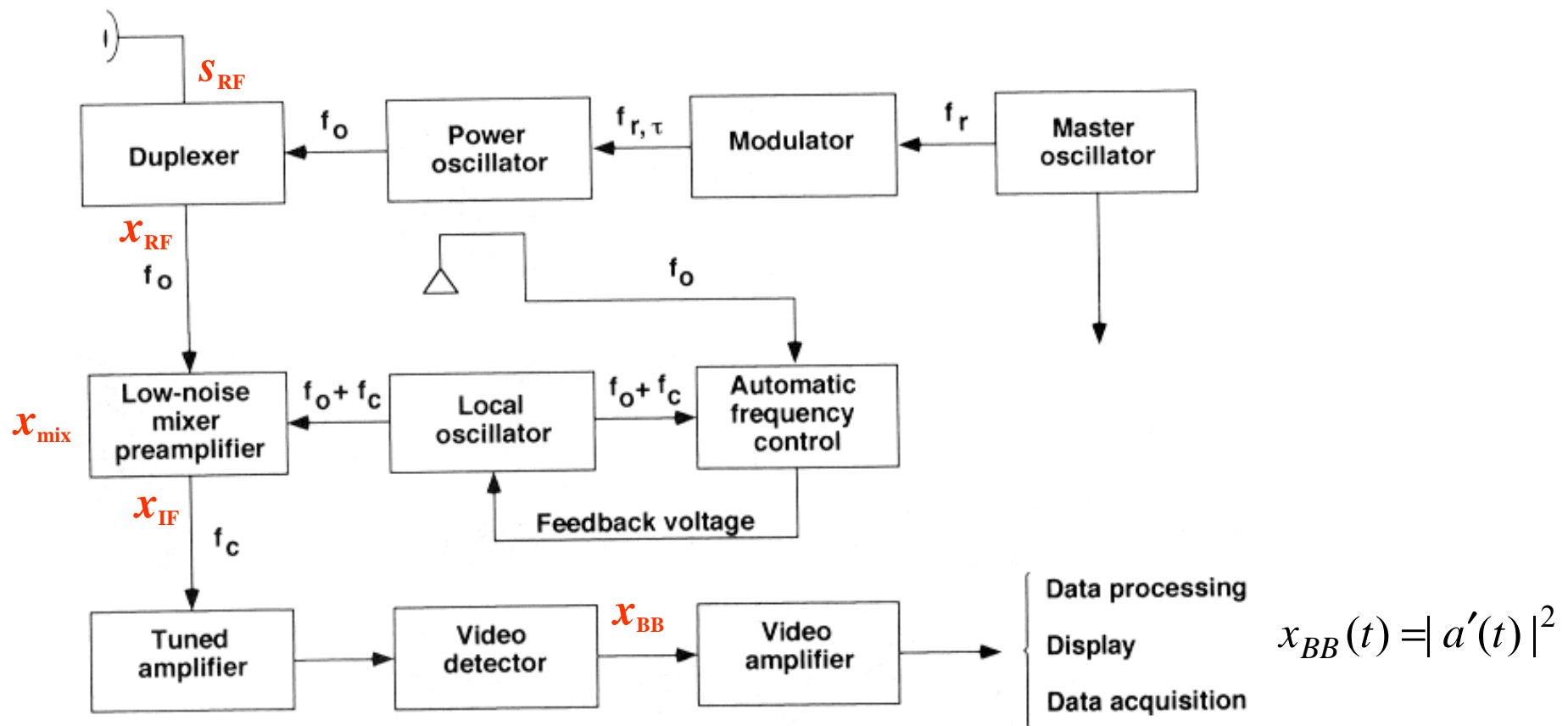
$f_0 \rightarrow 3-90 \text{ GHz}$ ;  $P_t \rightarrow 100-1000 \text{ kW}$

$f_r \rightarrow 250-2500 \text{ Hz}$ ;  $\tau \rightarrow 0.1-2.5 \mu\text{s}$



# RADAR SENSOR

## Pulsed incoherent scheme



# RADAR SENSOR

## Pulsed incoherent system

### ➤ Radar components

- Master oscillator (MO): system reference oscillator
- Modulator: device which determines the form e repetition of RF pulse
- Power oscillator: microwave tube (magnetron or klystron amplifier)
- Duplexer: 3-port circulator which separates TX from RX signal
- Local oscillator (LO): device which controls mixer frequency
- Mixer: device which mixes RF signal with LO doing amplification and filtering
- Amplifier: band-pass amplifier and filter for noise reduction of IF signal
- Video detector: detector of IF envelope power at baseband

### ➤ Signal processing

$s_{RF}(t) = [A \cos(2\pi f_0 t)] \text{rect}(f_r t)$	<i>Transmitted Radio Frequency (i.e., GHz)</i>
$x_{RF}(t) = a(t) \cos[2\pi f_0 t + \varphi(t)]$	<i>Received Radio Frequency</i>
$x_{mix}(t) = a(t) \cos[2\pi f_0 t + \varphi(t)] \{b \cos[2\pi(f_0 + f_c)t]\}$	<i>Mixer signals</i>
$x_{IF}(t) = \frac{a(t)b}{2} \cos[2\pi f_c t + \varphi(t)] = a'(t) \cos[2\pi f_c t + \varphi(t)]$	<i>Intermediate Frequency (i.e., MHz)</i>
$x_{BB}(t) =  x_{IF}(t) ^2 =  a'(t) ^2$	<i>Base-band signal (i.e., kHz)</i>

# RADAR SENSOR

## Doppler frequency shift

### Doppler frequency shift

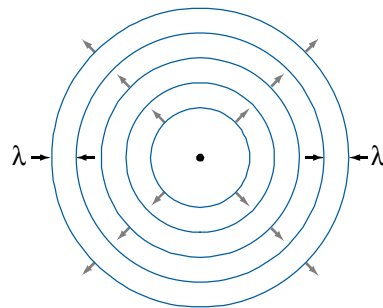
$$f_d = -\frac{u_r}{\lambda_0} = -\frac{\mathbf{u} \cdot \hat{\mathbf{r}}}{\lambda_0} = -\frac{u \cos \theta}{\lambda_0}$$

**If leaving target:**

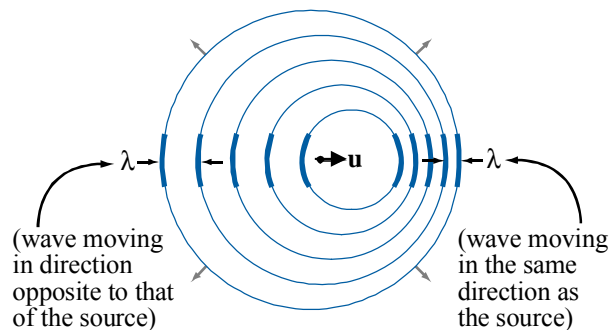
$$\theta < \pi/2 \Rightarrow f_d < 0$$

**If approaching target:**

$$\theta > \pi/2 \Rightarrow f_d > 0$$



(a) Stationary source



(b) Moving source

### Monostatic Radar

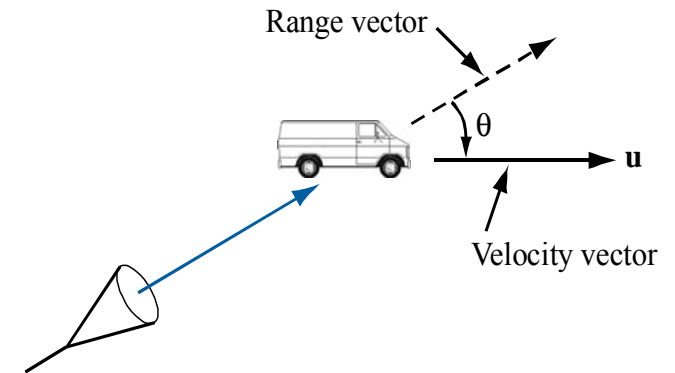
**Twice a Doppler frequency shift**

- from radar to target

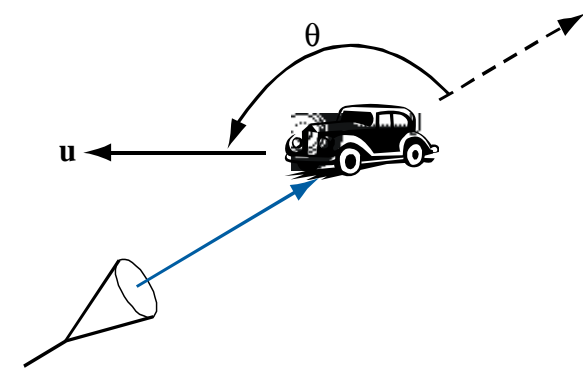
- from target to radar

$$f_{RF} = f_0 + f_d$$

$$f_d = -2 \frac{u_r}{\lambda_0} = -2 \frac{\mathbf{u} \cdot \hat{\mathbf{r}}}{\lambda_0}$$



(a)



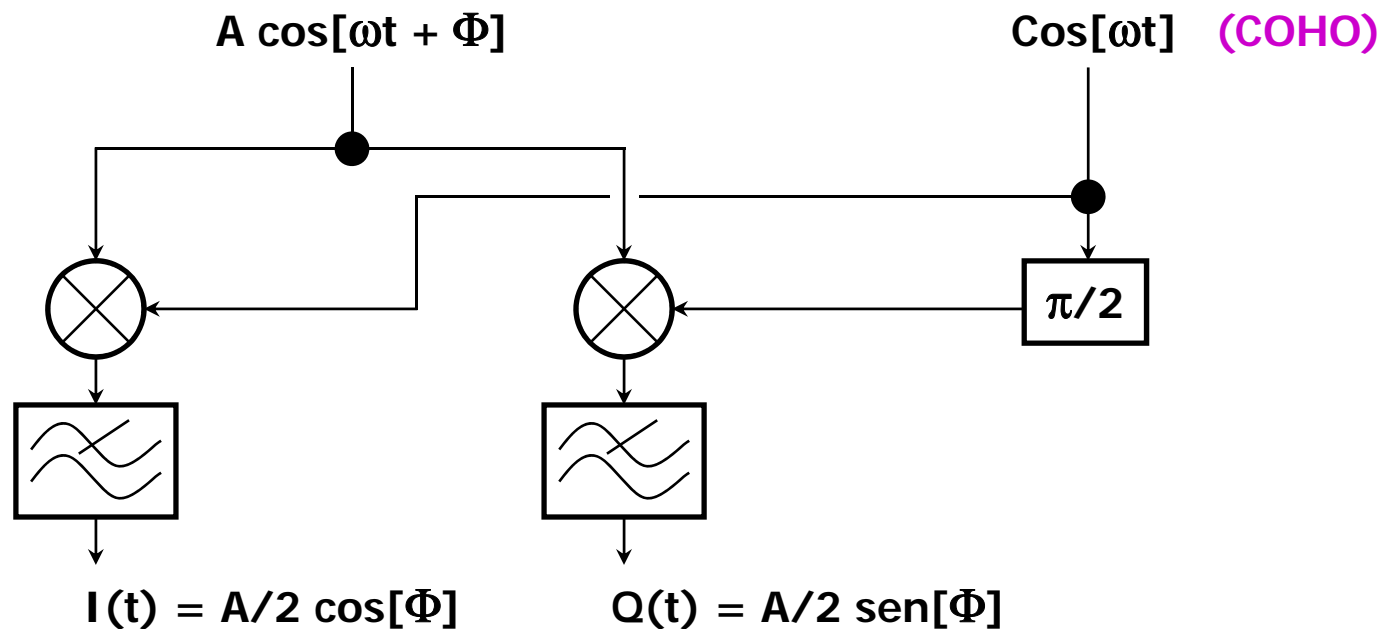
(b)

# RADAR SENSOR

## Pulsed coherent scheme

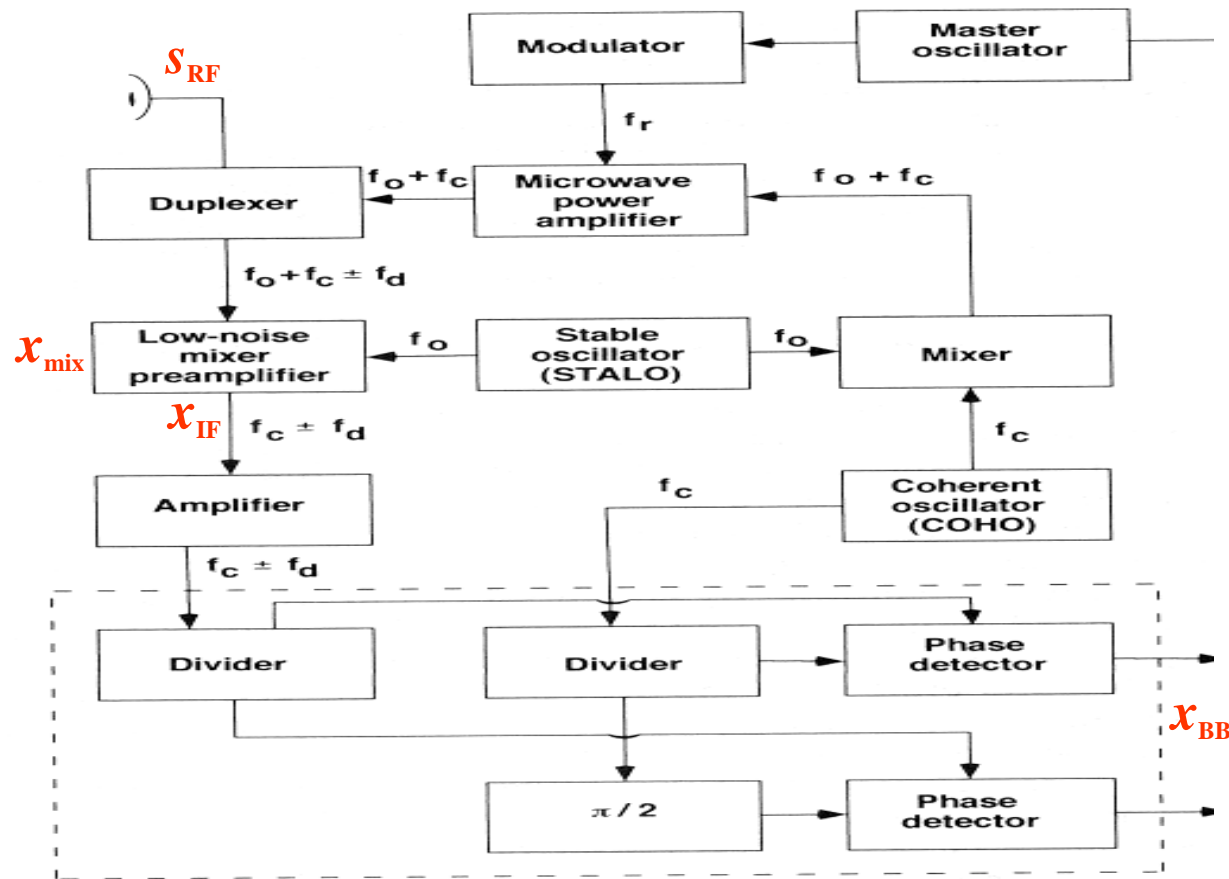
Per poter determinare la velocità Doppler del bersaglio puntiforme  
devo usare un

### RICEVITORE COERENTE



# RADAR SENSOR

## Pulsed coherent scheme



### Doppler frequency

$$\varphi(t) = -2 \left[ \frac{2\pi}{\lambda} r(t) \right] \Rightarrow$$

$$\frac{d\varphi(t)}{dt} \equiv \frac{4\pi}{\lambda} u_r(t) \equiv 2\pi f_d(t)$$

$$\Rightarrow f_d(t) = -\frac{2}{\lambda} u_r(t)$$

$$x_{BB}(t) = \begin{cases} I(t) = a'(t) \cos[\varphi(t)] \\ Q(t) = a'(t) \sin[\varphi(t)] \end{cases}$$

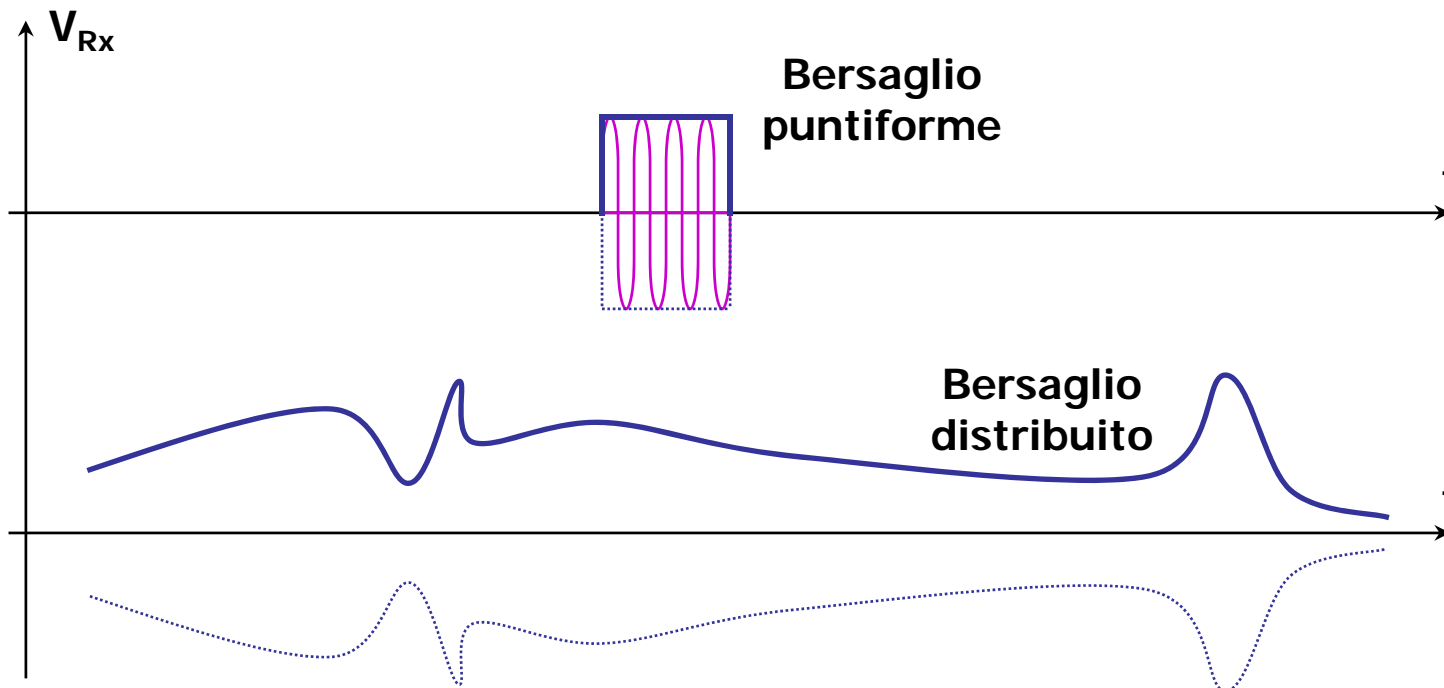
with  $V(t) = I(t) + jQ(t)$



# RADAR SENSOR

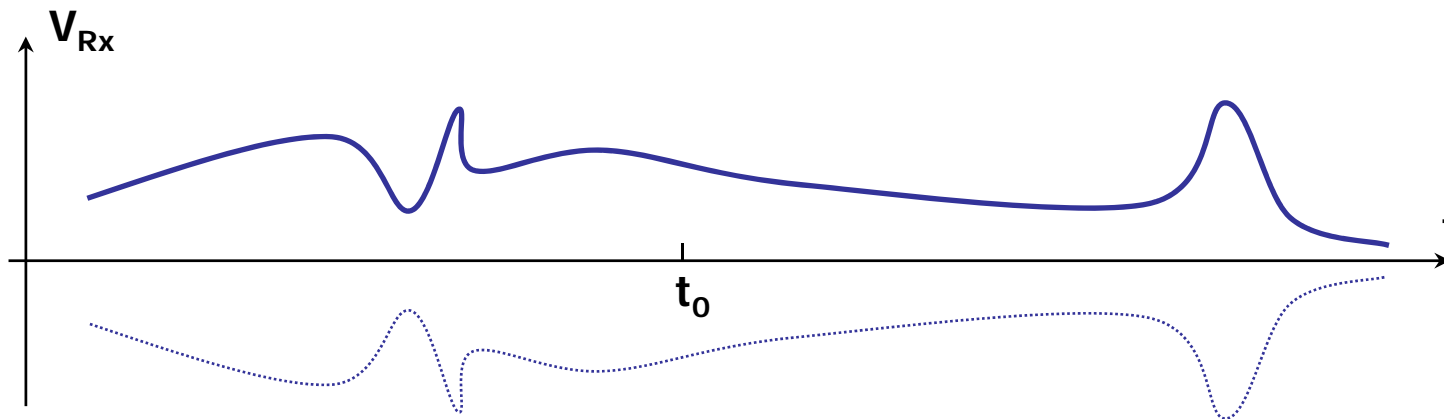
## Pulsed coherent scheme

- L'atmosfera è un bersaglio distribuito, ossia è costituita da un elevato numero di bersagli elementari (idrometeore)
- L'eco radar è la somma di tanti contributi elementari del tipo visto in precedenza (impulso rettangolare a RF)



# RADAR SENSOR

## Pulsed coherent system



L'eco ricevuto all'istante  $t_0$  è la somma dei contributi di tutti i bersagli elementari che si trovano ad una distanza dal radar compresa tra  $c \cdot (t_0 - \tau)/2$  e  $c \cdot t_0/2$

Il segnale ricevuto viene prima **campionato** (nel GPM-500C con una frequenza di 2.4 MHz, ossia ogni 417 ns) e quindi i campioni vengono processati.

Ogni campione è rappresentativo di un volume elementare (tronco di cono) con altezza pari a  $(c \cdot \tau)/2$

# RADAR SENSOR

## Pulsed coherent system

- Detection of received signal phase

$$\varphi(t) = -2 \left[ \frac{2\pi}{\lambda} r(t) \right] \Rightarrow \frac{d\varphi(t)}{dt} \equiv -\frac{4\pi}{\lambda} \frac{dr(t)}{dt} = -\frac{4\pi}{\lambda} u_r(t) \equiv 2\pi f_d(t) \Rightarrow f_d(t) = -\frac{2}{\lambda} u_r(t)$$

- Signal processing

$$s_{RF}(t) = [A \cos(2\pi(f_0 + f_c)t)] \text{rect}(f_r t) = [A \cos(2\pi f'_0 t)] \text{rect}(f_r t)$$

$$x_{RF}(t) = a(t) \cos[2\pi f_0 t + \varphi(t)] = a(t) \cos \left[ 2\pi f_0 t - \frac{2\pi}{\lambda} 2r(t) \right]$$

$$x_{mix}(t) = a(t) \cos[2\pi(f_0 + f_c)t + \varphi(t)] \{b \cos[2\pi f_0 t]\}$$

$$x_{IF}(t) = a'(t) \cos[2\pi f_c t + \varphi(t)] = a'(t) \cos[\varphi(t)] \cos(2\pi f_c t) - a'(t) \sin[\varphi(t)] \sin(2\pi f_c t)$$

$$x_{IF}(t) \equiv I(t) \cos(2\pi f_c t) - Q(t) \sin(2\pi f_c t) = \text{Re} \left\{ [I(t) + jQ(t)] e^{j2\pi f_c t} \right\} = \text{Re} \left\{ a'(t) e^{j\varphi(t)} e^{j2\pi f_c t} \right\}$$

$$x_{BB}(t) = \begin{cases} I(t) = a'(t) \cos[\varphi(t)] \\ Q(t) = a'(t) \sin[\varphi(t)] \end{cases} \quad \text{with } V(t) = I(t) + jQ(t)$$

# RADAR SENSOR

## Doppler velocity ambiguity

### ➤ Signal sampling (Shannon) theorem

- A signal with a finite energy (a maximum spectral frequency  $f_M$ ) is reconstructed from its temporal samples if the sampling frequency  $f_s$ :

$$f_M \geq \frac{f_s}{2} \Rightarrow T_M \leq \frac{T_s}{2}$$

### ➤ Unambiguous Doppler frequency

- If maximum frequency is  $f_r$ :

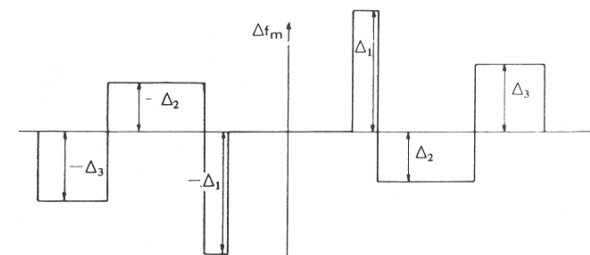
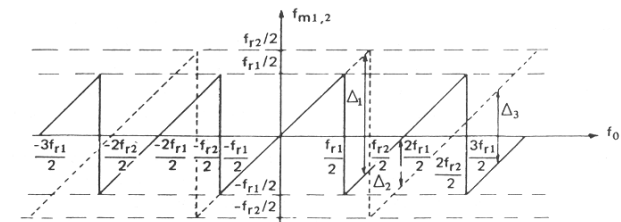
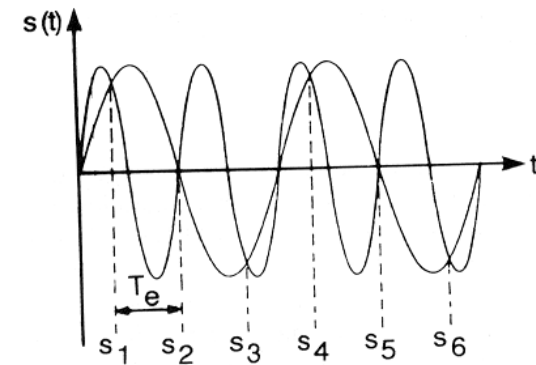
$$f_d \geq f_{d \max} = \frac{f_r}{2} \Rightarrow u_{r \max} = \pm f_r \frac{\lambda}{4}$$

- Velocity-range limit

$$u_{r \max} r_{\max} = [(f_r \lambda) / 4][c T_c / 2] = (c \lambda) / 8$$

### ➤ Ambiguity reduction techniques

- Doppler phase measured between 2 pulses
- Dual-PRF: double of unambiguous  $f_d$
- Phase and polarization pulse coding



# RADAR SENSOR

## Issues on coherent systems

### ➤ Microwave transmitter

- MAGNETRON: self-oscillating microwave tube with cylindrical structure (cross-field device), able to handle high power (up to 2000 kW peak). A high (up to 50 kV) DC voltage is applied to coaxial cathode (+) and anode (-) together with a static magnetic high field (up to 1 A/m). The emitted electron beam is spatially modulated (bunches) and output is coupled into one of ring resonant cavities. The latter determines the radio-frequency (RF) signal. Oscillation condition is governed by Hartree's curve. Overall efficiency is about 50%.
  - When input voltage is pulsed (as in radars), no inter-pulse coherence is guaranteed (even though phase information may be retained by injecting a CW signal).
- KLYSTRON: microwave-tube amplifier with a linear (linear-field) structure, able to handle medium power (up to 1000 kW). An electron beam is formed by an electron gun (thermoionic emission by cathode-anode at > 1000 K), passes through 2 or more resonant cavities in succession and is collected on a collector. Electric field of first cavity gaps induce electron bunches whose are "focused" on (load) resonant cavities which produce output RF signal.
  - Within pulsed radars, klystron is used as an amplifier excited by a local microwave oscillator. These scheme guarantees a high phase purity (low phase noise) and an inter-pulse coherence.

### ➤ Base-band detection

- Amplitude: logarithmic detector in order to deal with large amplitude (i.e.,  $a'(t)$ ) dynamics.
- Phase: linear detector to extract in-phase  $I(t)$  and quadrature  $Q(t)$  components

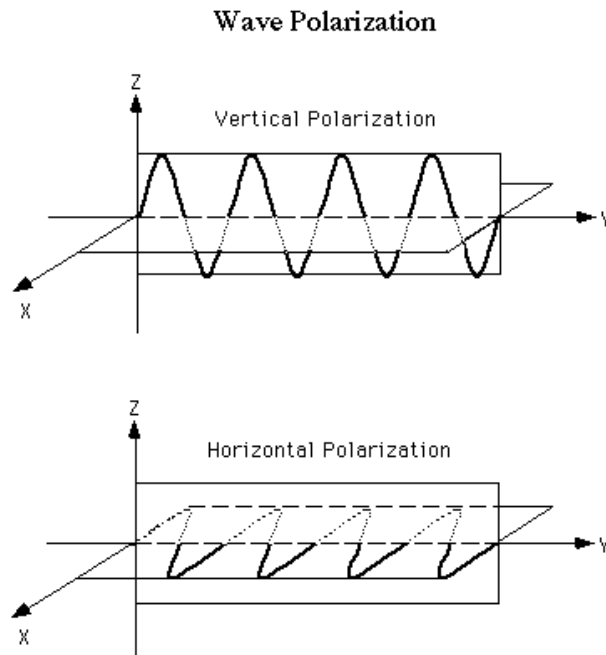
# RADAR SENSOR

## E.m. wave polarization

### ➤ Polarization states

Curve described in 3-D space by the free end of the monochromatic wave vector

- Horizontal H (w.r.t. ground)
- Vertical V
- Circular (LHC, RHC)

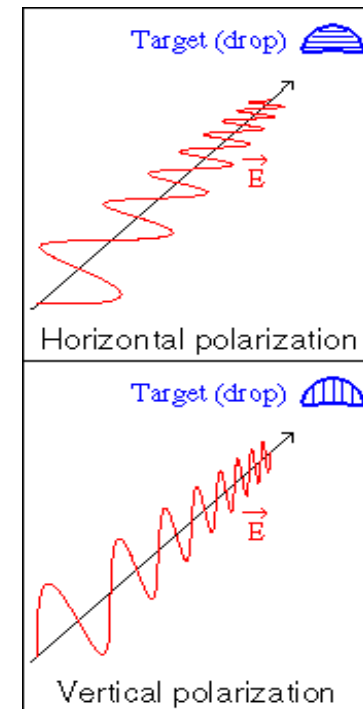


### ➤ Linear polarizations

$$\underline{E}(r, t) = \underline{E}_{0p}(t) \cos[2\pi f_0 t + \varphi(t)]$$

$$\underline{E}_{0v}(t) = E_{0v}(t) \hat{z}$$

$$\underline{E}_{0h}(t) = E_{0h}(t) \hat{x}$$

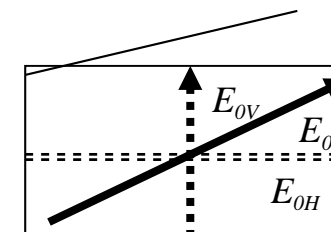
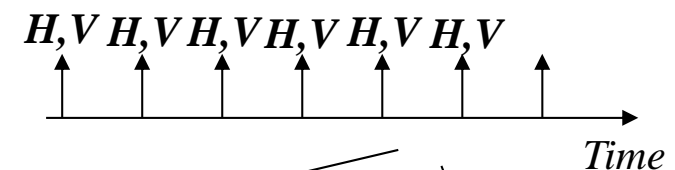
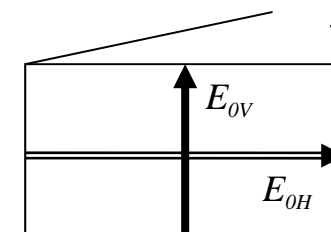
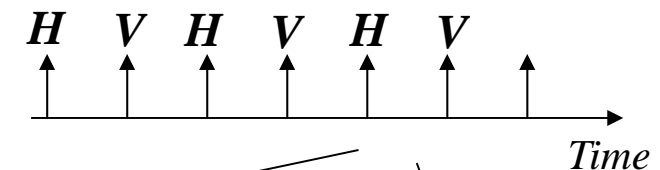


# RADAR SENSOR

## Issues on polarimetric radars

### ➤ Configuration

- Alternate transmission
  - Single-transmitter and single-receiver
  - Alternate transmission between H and V fields
  - ↑ Single receiver
  - ↑ Anomalous propagation removal
  - ↓ Polarization switching
  - ↓ H and V non-contemporary data
- Simultaneous transmission
  - Single transmitter and dual receiver
  - Simultaneous transmission of H and V fields
  - ↑ No polarization switching (costly and loss)
  - ↑ No delay in H and V acquisition data
  - ↑ Fast scanning (half time w.r.t. alternate trans.)
  - ↓ Cross-polarization effects in received data
  - ↓ Dual receiver and limitation in polarimetric features

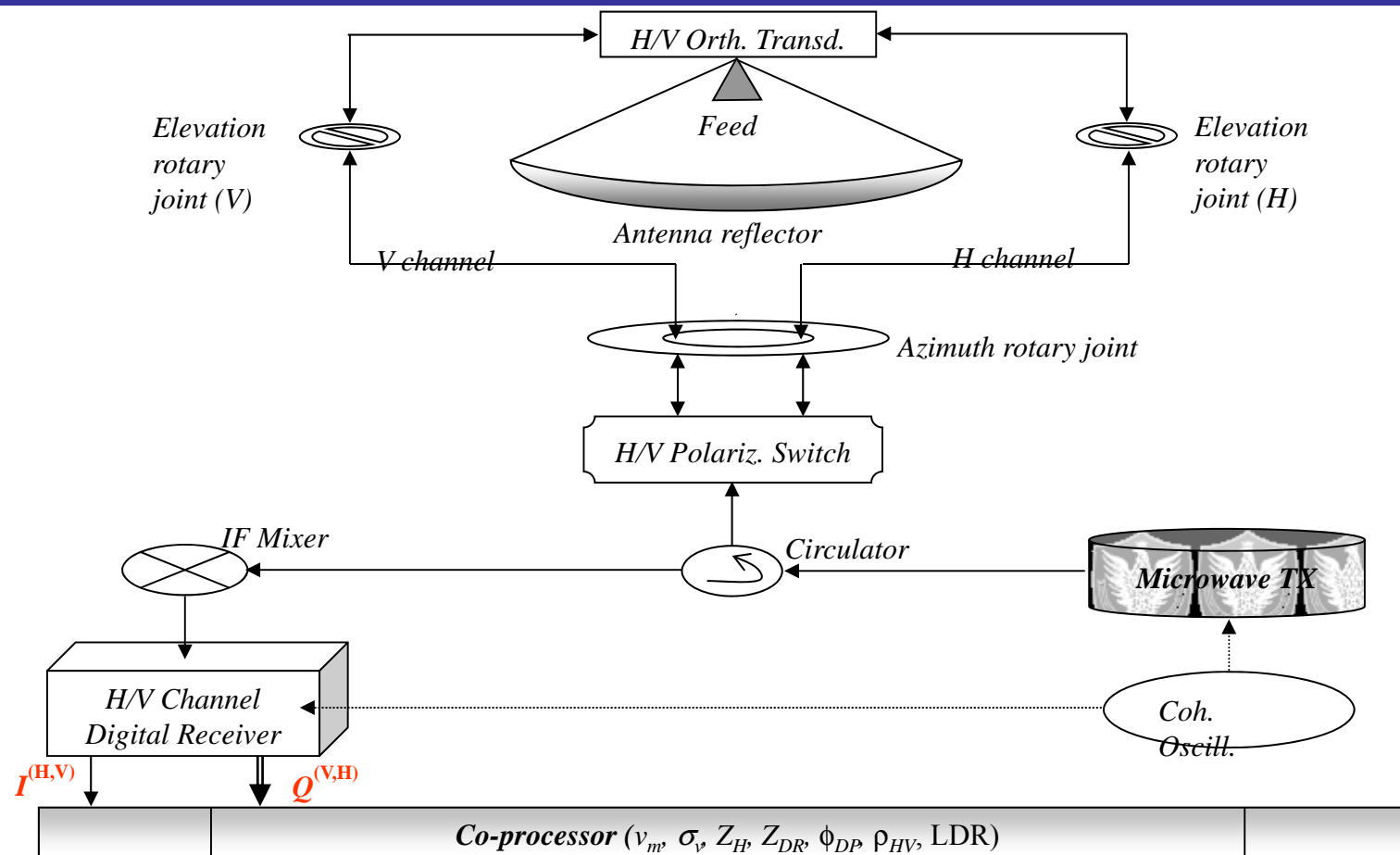


### ➤ Receiver

- Digital receiver at IF stage
  - Sampling of IF signal (A/D digitizers)
  - Digital processing of binary I/Q sequences

# RADAR SENSOR

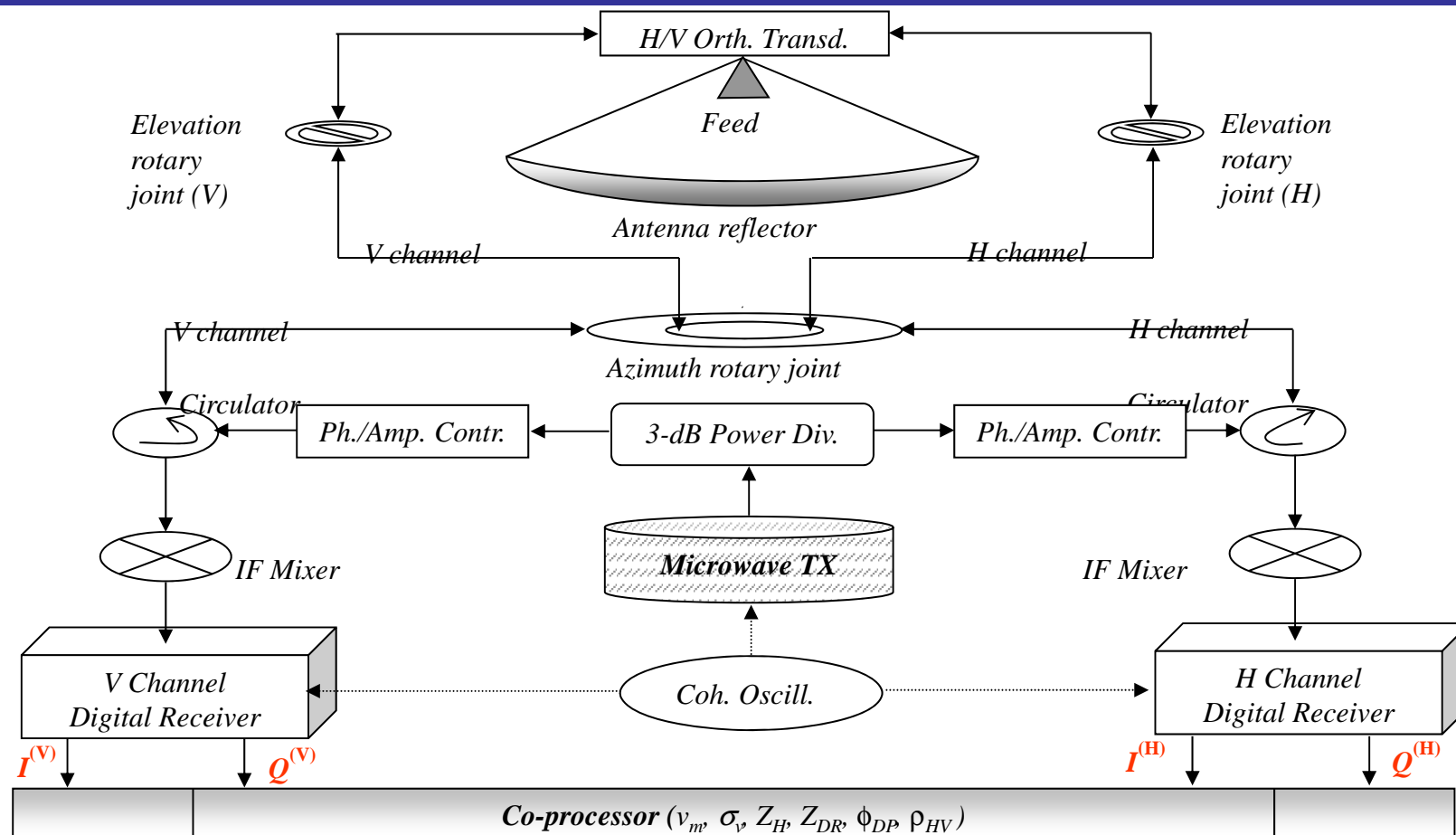
## Alternate polarim. system





# RADAR SENSOR

## Simultaneous polarim. system



# RADAR SENSOR

## Receiver sensitivity

### ➤ Noise sources

- cosmic, artificial (human) and atmospheric emission (H<sub>2</sub>O, O<sub>2</sub> at MW)
- internal, due to electron thermal mobility (white Gaussian noise at MW).

### ➤ Internal noise of receiver

- Receiver input power with a resistance load at T<sub>0</sub> and bandwidth B<sub>n</sub>:  $N_i = kT_0 B_n$
- Noise figure F<sub>n</sub> or equivalent noise temperature T<sub>e</sub>:  $T_e = T_0(F_n - 1)$



### ➤ Minimum detectable signal

- Minimum signal detectable over the noise receiver (e.g., -110 dBm = 10<sup>-11</sup> mW)

$$F_n = \frac{S_i / N_i}{S_0 / N_0} \Rightarrow S_{iMin} = F_n N_i \frac{S_0}{N_0} \Big|_{Min} = F_n (kT_0 B_n) \frac{S_0}{N_0} \Big|_{Min}$$

- Matched filters to increase S/N: for square pulses.  $B_n \approx 1/\tau$  and triangular shape

# RADAR SENSOR

## Aperture antenna basics

### ➤ Radar antennas

- Parabolic reflector antenna
- Prime-focus feeder (horn)
- Optional radome protection
- Optional polarization capability

### ➤ Antenna radiation

- Radiation pattern function

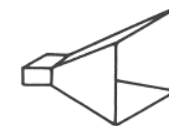
$$f(\theta, \varphi) = FFT[E_a(x, y)]$$

- Irradiated electric far-field (for  $r \geq L^2/\lambda$ )

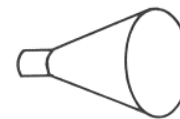
$$E(r, \theta, \varphi) = E_0 f(\theta, \varphi) \frac{e^{-jkr}}{r}$$

- Power flux density [ $W/m^2$ ]

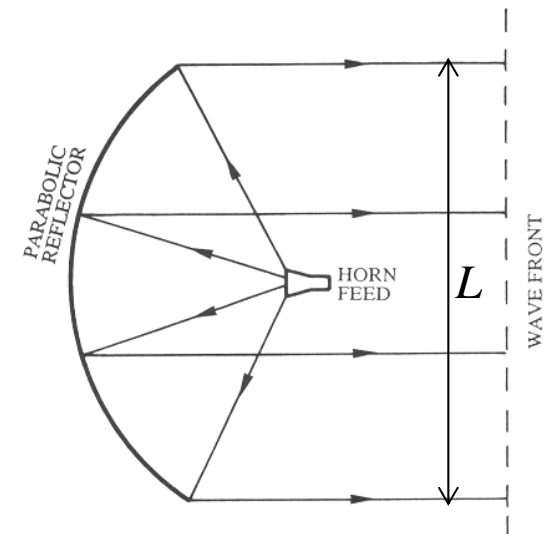
$$P(r, \theta, \varphi) = \frac{1}{2\eta_0} |E(r, \theta, \varphi)|^2 = \frac{E_0}{2\eta_0 r^2} |f(\theta, \varphi)|^2$$



(a) Pyramidal horn



(b) Conical horn



(c) Classical parabolic antenna

# RADAR SENSOR

## Antenna parameters

### ➤ Directivity and Gain

$$D(\theta, \varphi) = \frac{P(r, \theta, \varphi)}{W_T / 4\pi r^2} = \frac{|f(\theta, \varphi)|^2}{W_T / 4\pi}, \quad G(\theta, \varphi) = \eta_r D(\theta, \varphi)$$

### ➤ Effective area

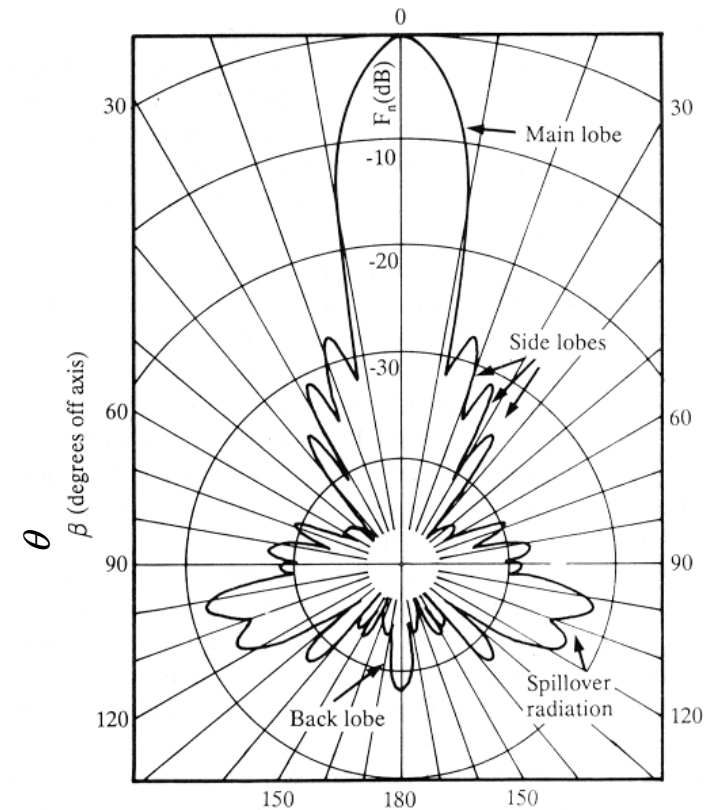
$$A_e(\theta, \varphi) = \frac{W_R}{P_i(r, \theta, \varphi)} = \frac{\lambda^2}{4\pi} D(\theta, \varphi) = D_M |f_n(\theta, \varphi)|^2$$

### ➤ Beamwidth (reflector antennas)

$$|f_n(\Theta_{3dB} / 2, \varphi)|^2 = 0.5 \Rightarrow \Theta_{3dB} \cong 70 \frac{\lambda}{L}, \quad D_M \cong \frac{4\pi}{(\Theta_{3dB})^2}$$

### ➤ Polarization capability

- Ortho-Mode Transducer (H/V) at horn feeder
- Low sidelobes (<30 dB) similar for H/V patterns
- Isolation between H/V patterns (no cross-pol.)
  - Azimuthally symmetric radiation pattern at H/V
- Effects of supporting struts and radome paneling



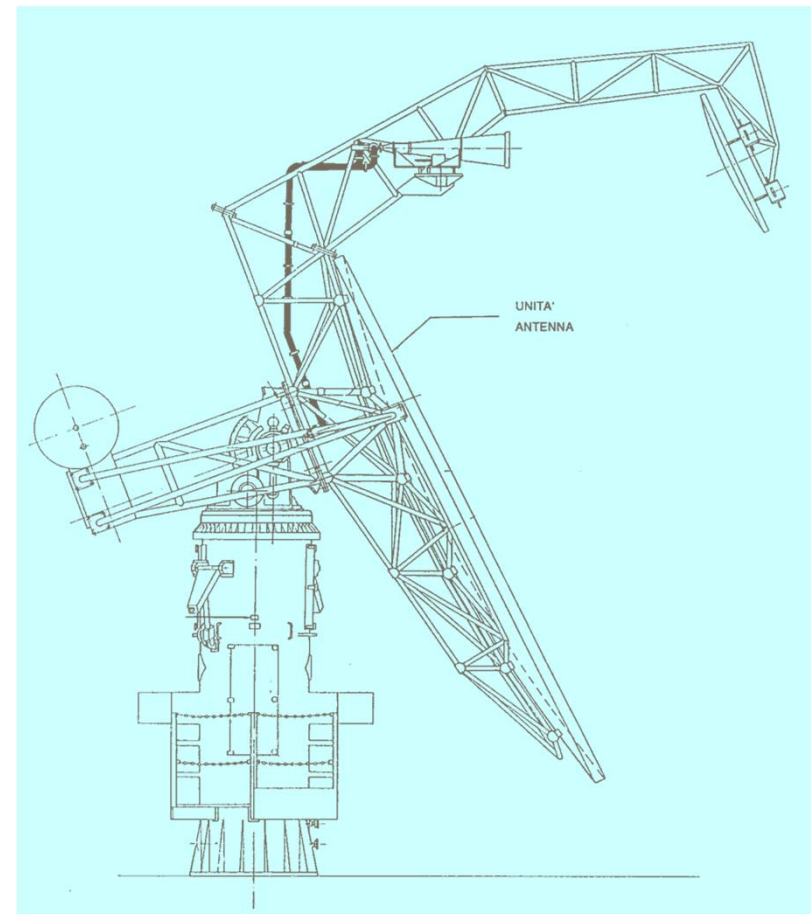
# RADAR SENSOR

## Antenna parameters

**Example:**  
**Radar GPM-500C**

**Cassegrain dual-offset**

- fascio di  $0.9^\circ$  (a -3 dB)
- velocità max 30 deg/s
- lobi secondari molto bassi
- ottima simmetria tra H e V



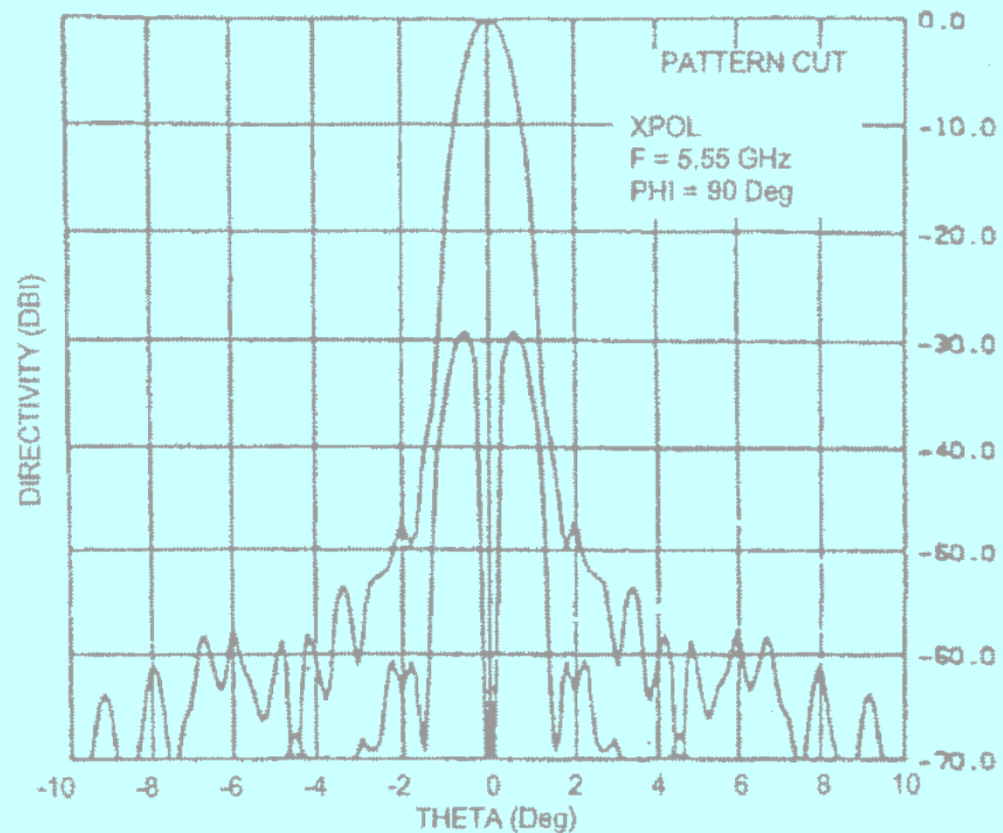
# RADAR SENSOR

## Antenna diagram

Example:  
Radar GPM-500C

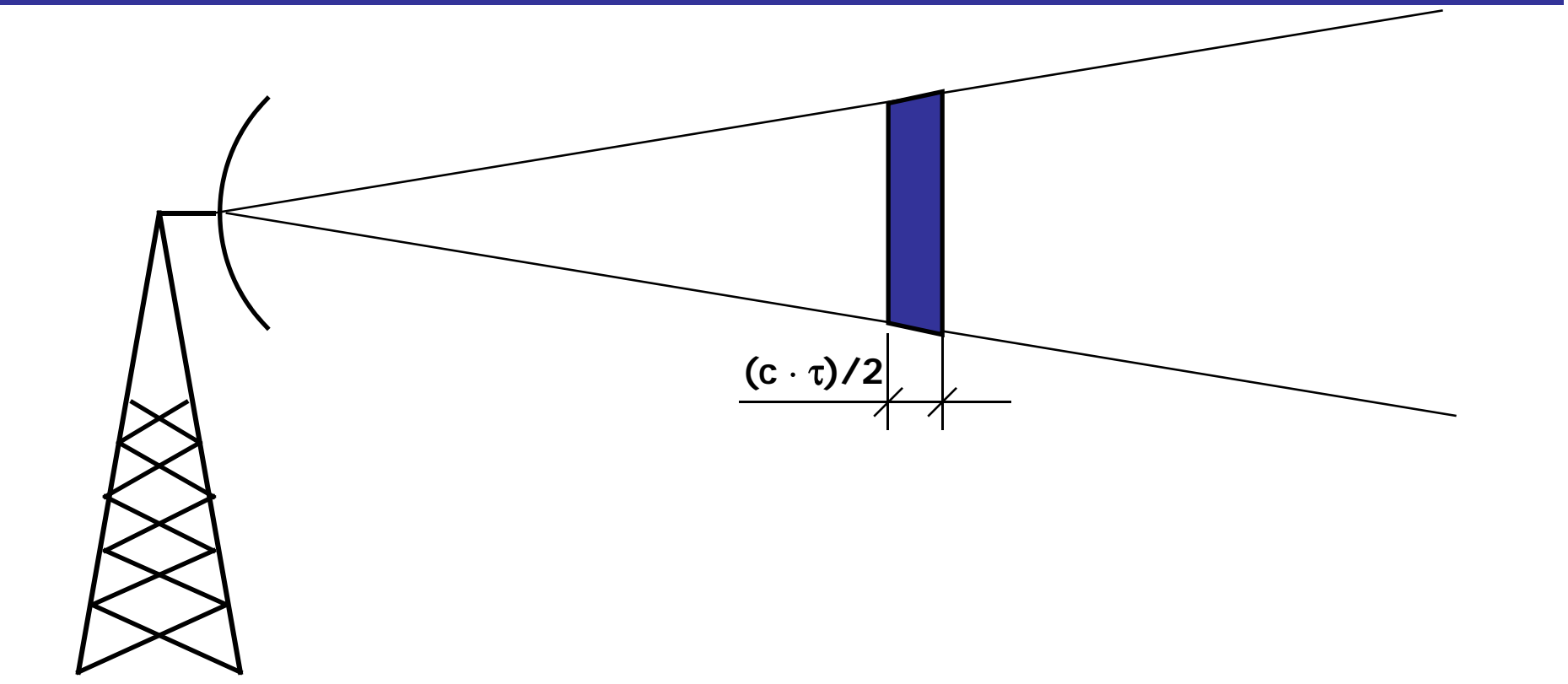
Cassegrain dual-offset

- fascio di  $0.9^\circ$  (a -3 dB)
- velocità max 30 deg/s
- lobi secondari molto bassi
- ottima simmetria tra H e V



# RADAR SENSOR

## Volume resolution



# RADAR SENSOR

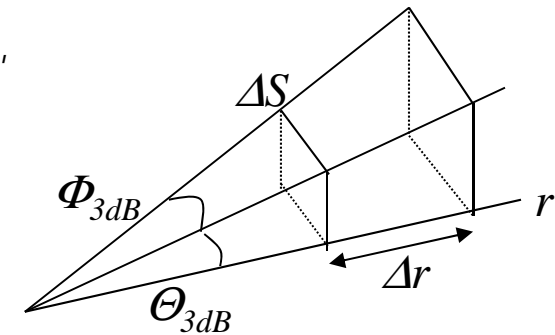
## Volume resolution

### ➤ Spatial resolution

- Capability of discriminating 2 objects in space (range, zenith and azimuth in spherical coordinates)

### ➤ Radial resolution

- Resolution in range:  $\Delta r = \frac{c\tau}{2}$ 
  - Matched receivers have a lower radial resolution
  - Use of pulse compression to increase spatial res.



### ➤ Transverse resolution

- Resolution in zenith-azimuth:  $\Delta\Omega = \frac{\Delta S}{r^2} \Rightarrow \Delta S = r^2 \Delta\Omega \cong r^2 \Omega_{3dB}$ 
  - Affected by side lobes contribution.

### ➤ Pulse volume resolution

$$V_{bin} = \Delta r \Delta S = (c\tau/2)r^2 \Omega_{3dB} \cong \frac{\pi}{4}(c\tau/2)r^2 \Theta_{3dB} \Phi_{3dB}$$

- Volume varies with range (for 1° and 1 μs, at 100 km Δr=150 m and ΔS=1745<sup>2</sup> m<sup>2</sup>)
  - Energy volume distribution can be not uniform in range (e.g., matched filters) and in transverse direction (e.g., directivity angular variation)



# RADAR METEOROLOGY

## Lecture's contents

### ➤ Radar sensor

- Pulsed, Doppler, and polarimetric systems
- Receiver sensitivity
- Antenna specifications
- Radar volume resolution

### ➤ Radar equation

- Atmospheric refraction and attenuation
- Radar equation for single and distributed scatterers

### ➤ Radar signal

- Signal statistics and decorrelation
- Noise reduction techniques

### ➤ Radar applications

- Operational problem overview
- Clouds and precipitation
- Rainfall backscattering and polarimetric measurables
- Example of radar measurements and estimates

# RADAR EQUATION

## Atmospheric refraction

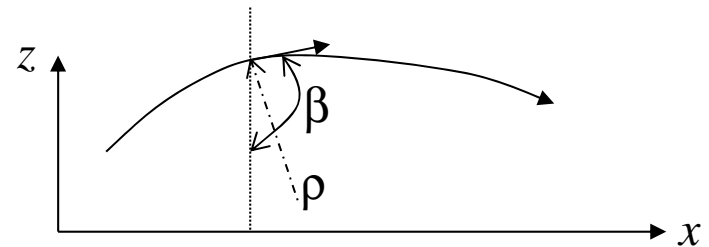
### ➤ Atmospheric refractivity

- Refractive index:

$$n = \frac{c}{v} = \sqrt{\epsilon_r} = n' - jn''$$

- Refractivity:

$$N = (n' - 1)10^6 = f(p/T, e/T^2)$$



### ➤ Optical rays

- Geometrical optics equation:

$$\frac{1}{\rho} = \frac{10^6}{n} \frac{dN}{dz} \cos \beta, \quad N(z) = N_0 e^{-z/H}$$

with  $\rho$  ray curvature.

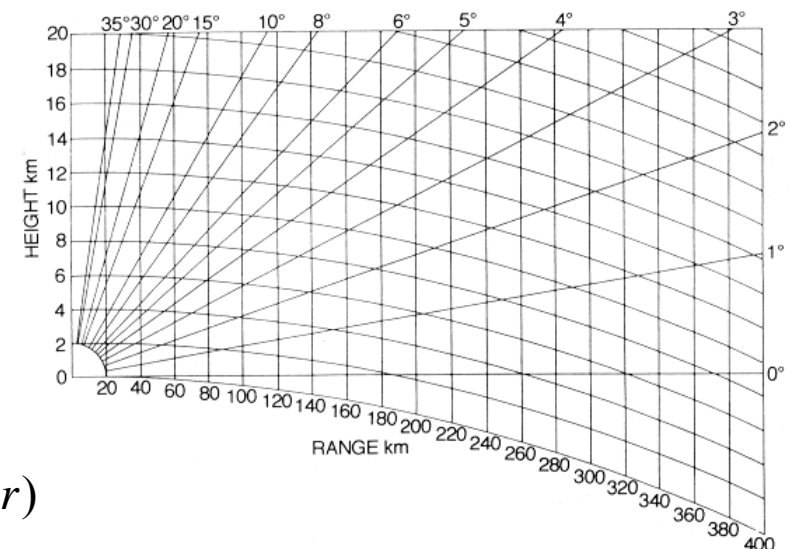
- If standard atmosphere,  $dN/dz = -40$  [N/km]
  - Earth effect. radius  $R_e = 4/3 R_T = 4/3 (6370) = 8500$
  - Rays are rectilinear and Earth flatter

### ➤ Specific attenuation

- For two-way path, the attenuation factor L:

$$dW = -2\alpha W dr \Rightarrow W(r) = W_0 e^{-2 \int_0^r \alpha dr} = W_0 L^2(r)$$

*Atmosphere with  $N_0=313$  and  $H=7$  km*



# RADAR EQUATION

## Single scatterer form

- Power flux density upon scatterer

$$P_i(r, \theta, \varphi) = \frac{W_T}{4\pi r^2} G_M |f_n(\theta, \varphi)|^2 L(r)$$

- Backscattering radar cross section

$$\sigma_b(-\theta, -\varphi, \theta, \varphi) = 4\pi r^2 \frac{P_r(r, \theta, \varphi)}{P_i(r, \theta, \varphi)}$$

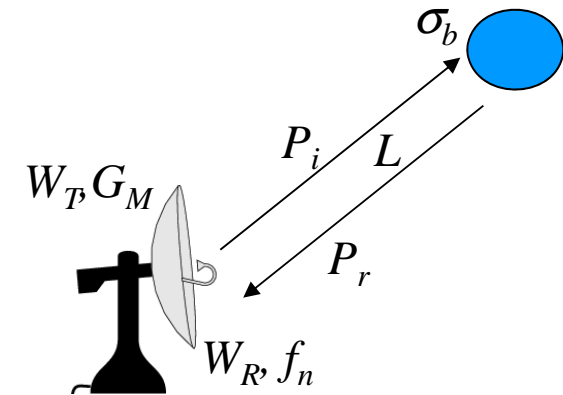
- Received power

$$P_r(r, \theta, \varphi) = \frac{\sigma_b}{4\pi r^2} P_i = \frac{\sigma_b}{4\pi r^2} L \frac{W_T}{4\pi r^2} G_M |f_n(\theta, \varphi)|^2 L \Rightarrow W_R = A_e P_r = \frac{\lambda^2}{4\pi} G P_r$$

- Radar equation for a single scatterer

$$W_R = \left( \frac{W_T G_M^2 |f_n(\theta, \varphi)|^4 \lambda^2}{(4\pi)^3} \right) \sigma_b \frac{L^2}{r^4} = \left( \frac{W_T \eta_r^2 A_e^2}{4\pi r \lambda^2} \right) \sigma_b \frac{L^2}{r^4} \Rightarrow \boxed{W_R = C_1 L^2 \frac{\sigma_b}{r^4}}$$

- For a point target,  $f_n=1$  and  $W_R \propto 1/r^4$ . Constant  $C_1$  depends on radar specs.



# RADAR EQUATION

## Effect of distributed scatterer

### ➤ Distributed scatterers

- Set of a large number  $N_{part}$  of equal-size scatterers, simultaneously present in the same resolution volume  $V_{bin}$  with randomly distributed phase and totally filling the volume ( e.g. raindrops).
- Volumetric reflectivity [ $m^{-1}$ ]

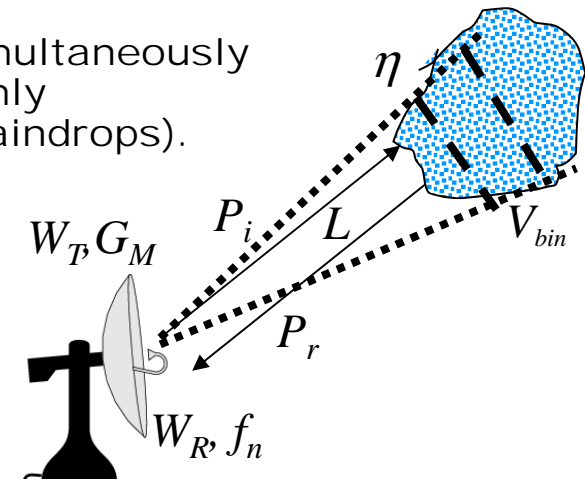
$$\eta = \left( \frac{\sum_{i=1}^{N_{part}} \sigma_{bi}}{dV} \right) \Rightarrow \sum_{i=1}^{N_{part}} \sigma_{bi} = dV \left( \frac{\sum_{i=1}^{N_{part}} \sigma_{bi}}{dV} \right) = dV \eta$$

### ➤ Particle and total received power

$$dW_{Ri} = \left( \frac{W_T G_M^2 \lambda^2}{(4\pi)^3} \right) \frac{|f_n(\theta, \varphi)|^4}{r_i^4} \sigma_{bi} = C \frac{|f_n(\theta, \varphi)|^4}{r_i^4} L^2 \sigma_{bi}$$

$$dW_R = \sum_{i=1}^{N_{part}} dW_{Ri} = \sum_{i=1}^{N_{part}} C \frac{|f_n(\theta, \varphi)|^4}{r_i^4} L^2 \sigma_{bi} \cong C \frac{|f_n(\theta, \varphi)|^4}{r^4} L^2 \sum_{i=1}^{N_{part}} \sigma_{bi} \frac{dV}{dV} \cong C \frac{|f_n(\theta, \varphi)|^4}{r^4} L^2 \eta dV$$

$$W_{Rtot} = \int_{V_{bin}} dW_R = \int_V C \frac{|f_n(\theta, \varphi)|^4}{r^4} L^2 \eta dV \cong \langle W_R \rangle$$



# RADAR EQUATION

## Distributed scatterer form

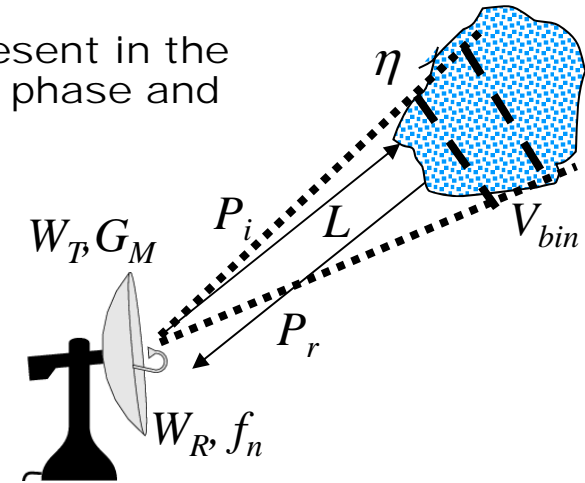
### ➤ Distributed scatterers

- Set of a large number of scatterers, simultaneously present in the same resolution volume  $V_{bin}$  with randomly distributed phase and totally filling the volume ( e.g. raindrops).

### ➤ Total (average) received power

$$\langle W_R \rangle = \left( \frac{W_T G_M^2 \lambda^2}{(4\pi)^3} \right) L^2 \int_{V_{bin}} \frac{|f_n(\theta, \varphi)|^4}{r^4} \eta dV = \left( \frac{W_T G_M^2 \lambda^2}{(4\pi)^3} \right) L^2 \eta \Delta r \int_{V_{bin}} \frac{|f_n(\theta, \varphi)|^4}{r^2} d\Omega$$

where  $dV = dr(r^2 d\Omega)$ ,  $V_{bin}$  : radar resolution volume



### ➤ Radar equation for volume scattering

$$\int_V \frac{|f_n(\theta, \varphi)|^4}{r^2} d\Omega \cong \frac{1}{r^2} \frac{\pi \Theta_{3dB} \Phi_{3dB}}{8 \ln 2} \quad (\text{Probert - Jones correction})$$

$$\langle W_R \rangle = \left( \frac{W_T G_M^2 \lambda^2}{(4\pi)^3} \right) L^2 \eta \Delta r \frac{1}{r^2} \frac{\pi \Theta_{3dB} \Phi_{3dB}}{8 \ln 2} \Rightarrow \boxed{\langle W_R \rangle = C_2 L^2 \frac{\eta}{r^2}}$$

# RADAR SIGNAL

## Signal fluctuations

### ➤ Power fluctuations of radar echo

- Complex envelope due to i-th scatterer:

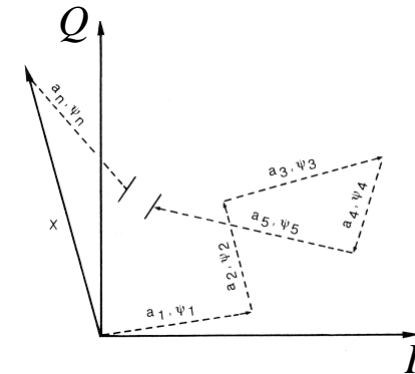
$$V_i(t) = A_i e^{j\varphi_i(t)}, \quad \varphi_i(t) = -4\pi r_i(t) / \lambda + \psi_i$$

- Total complex envelope due to distributed scatterers:

$$V(t) = I(t) + jQ(t) = \sum_i V_i(t) = \sum_i A_i e^{j\varphi_i(t)} = A(t) e^{j\varphi(t)}$$

- In-phase and quadrature components

$$\begin{cases} I(t) = \text{Re}[V(t)] = A(t) \cos \varphi(t) \\ Q(t) = \text{Im}[V(t)] = A(t) \sin \varphi(t) \end{cases} \quad \text{with} \quad A = \sqrt{I^2 + Q^2}, \quad \varphi = \text{artg} \frac{Q}{I}$$



### ➤ Radar echo power

$$W_R(t) = kV(t)V(t)^* = A^2(t) = k \sum_i A_i^2 + k \sum_{i \neq j} A_i A_j e^{j(-4\pi/\lambda)(r_i - r_j)} = \begin{cases} W_{Rdc} + W_{Rac} \\ \langle W_R \rangle + \tilde{W}_R \end{cases}$$

- From one pulse to next, power fluctuation  $W_{ac}$  are related to scatterer random displacement (velocity or Doppler frequency).
- It results  $\langle W_{ac} \rangle = 0$  and  $W_{ac}$  decreases doing averages on independent samples (such that displacement  $> \lambda$ ). For estimating  $\Delta\Phi$ , samples must be correlated to avoid ambiguity.

# RADAR SIGNAL

## Statistics of power fluctuations

### ➤ Distribution of $I, Q$

- For the central limit,  $V(t)$  is Gaussian so that  $I$  and  $Q$
- $A$  and  $\Phi$  are independent random variables with  $\Phi$  uniformly distributed in  $0-2\pi$ .

$$p(I, Q) = \frac{1}{\sqrt{2\pi\sigma^2}} e^{-(I^2+Q^2)/2\sigma^2}, \quad p(I, Q)dIdQ = p(A, \varphi)dAd\varphi, \quad dIdQ = AdAd\varphi$$

### ➤ Marginal distribution of $A, \varphi$

$$p(A, \varphi) = \frac{A}{\sqrt{2\pi\sigma^2}} e^{-A/2\sigma^2} \Rightarrow \begin{cases} p(A) = \frac{A}{\sqrt{2\pi\sigma^2}} e^{-A/2\sigma^2} & \text{Rayleigh pdf} \\ p(\varphi) = 1/2\pi & \text{Uniform pdf} \end{cases}$$

### ➤ Square-law detector statistics of received power

$$W_R = A^2, \quad dW_R = 2AdA, \quad p(W_R)dW_R = p(A)dA \Rightarrow p(W_R) = \frac{1}{2\sigma^2} e^{-W_R/2\sigma^2} \quad \text{Exponential pdf}$$

# RADAR SIGNAL Decorrelation time

## ➤ Doppler spectrum

- Duration  $t_d$  needed to ensure sample independence depends on volume and wavelength
- Doppler spectrum is of Gaussian type for uniform regions:

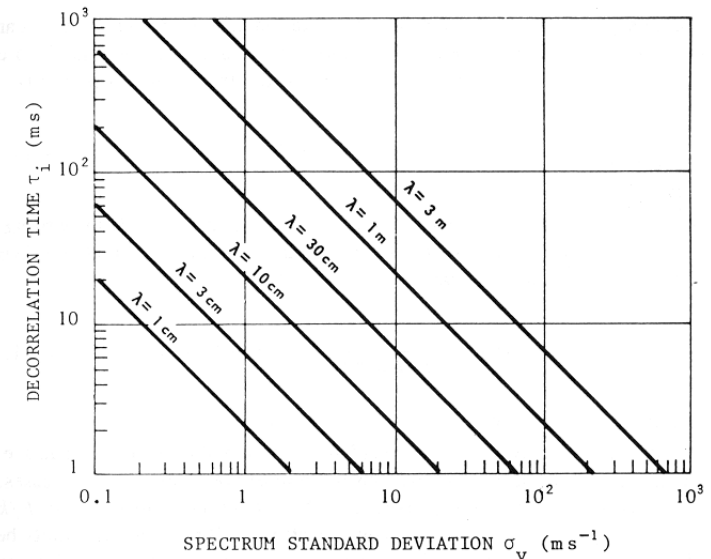
$$p(f_d) = p_0 e^{-f_d^2 / 2\sigma_f^2}; \quad f_d = (2/\lambda)u_r$$

- Time autocorrelation function:

$$R(t_d) = R_0 e^{-t_d^2 / 2\sigma_t^2}, \quad \sigma_t = \frac{1}{2\pi\sigma_f} = \frac{1}{2\pi(2/\lambda)\sigma_u}$$

## ➤ Decorrelation time $t_{inc}$

$$R(t_{inc}) / R_0 = 0.02 \Rightarrow \boxed{t_{inc} = 2\lambda / \sigma_v}$$



*Snow:*  $\sigma_u = 0.5 \text{ m/s}$

*Rain:*  $\sigma_u = 1 \text{ m/s}$



# RADAR SIGNAL

## Time-space data integration

### ➤ Signal variance reduction

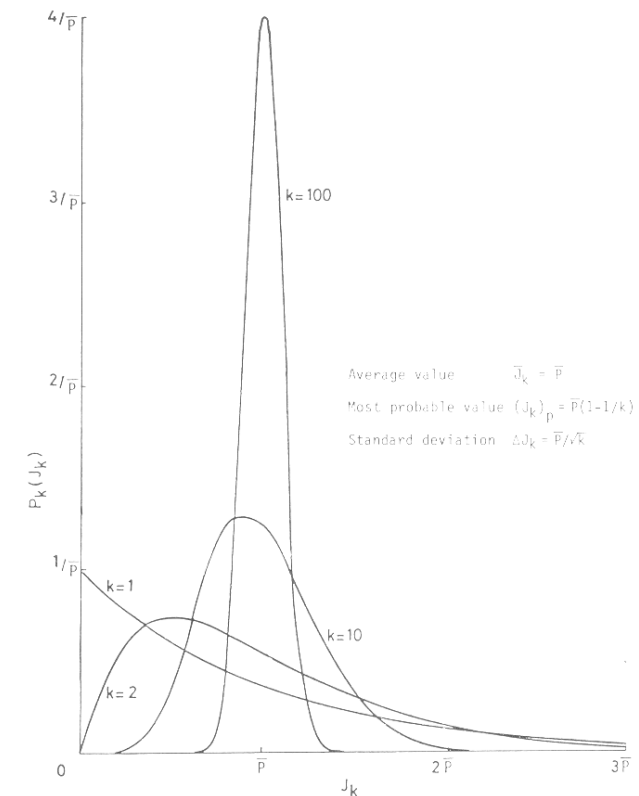
- For meteorological appliatio,  $W_{dc}$  is of interest so that  $W_{ac}$  must be reduced
- How many independent samples of received power must be averaged to accurately estimate  $W_{dc} = \langle W \rangle$  using a square-law detector?

$$J_k = \frac{1}{k} \sum_{n=1}^k W_{Rn} \quad \text{with} \quad p(J_k) dJ_k = p(A_n^2) d(A_n^2)$$

- The final pdf is almost Gaussian for  $k > 10$
- For  $k=30$ ,  $\langle W \rangle$  error estimate is about 1 dB

### ➤ Spatial integration

- For fast scannig, spatial averaging is performed together with time integration
- Reduction of measurement spatial resolution (e.g., range similar to transverse)



# RADAR METEOROLOGY

## Lecture's contents

### ➤ Radar sensor

- Pulsed, Doppler, and polarimetric systems
- Receiver sensitivity
- Antenna specifications
- Radar volume resolution

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- Atmospheric refraction and attenuation
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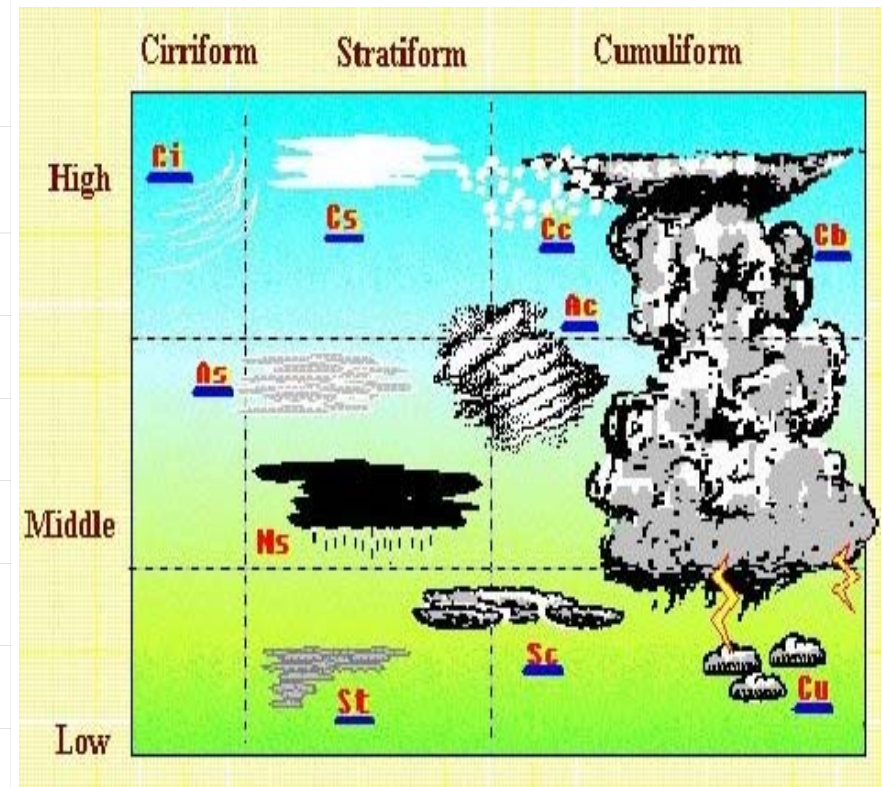
### ➤ Radar applications

- Operational problem overview
- Clouds and precipitation
- Rainfall backscattering and polarimetric measurables
- Example of radar measurements and estimates

# RADAR APPLICATIONS

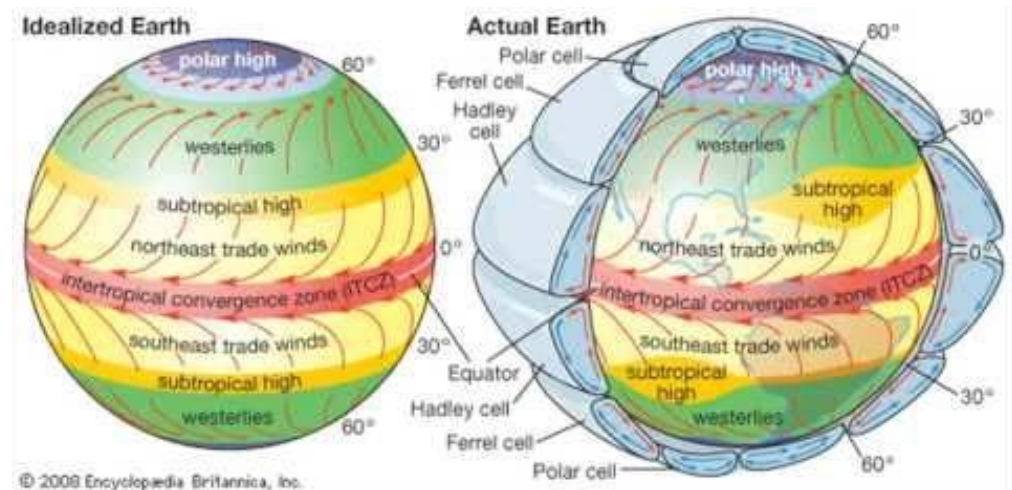
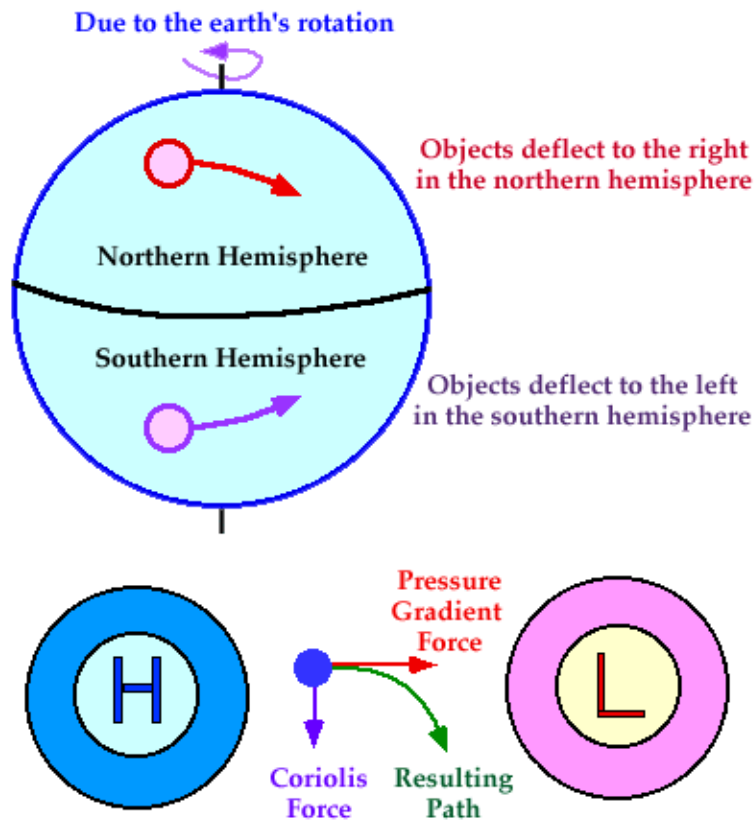
## Clouds and precipitation

<i>Genus</i>	<i>Sym.</i>	<i>Region</i>	<i>Species</i>	<i>RR (mm/h)</i>
<i>Cumulus</i>	Cu	Vert. develop. (0-6 Km)	<i>Mediocris</i> <i>Congestus</i> <i>Incus</i>	0 < 30 < 60
<i>Cumulonimbus</i>	Cb	Vert. develop. (0-12 km)		10÷100
<i>Stratus</i>	St	low (0-2 km)		< 2
<i>Stratocumulus</i>	Sc	low (0-2 km)		< 5
<i>Nimbostratus</i>	Ns	middle (2-6 km)		< 15
<i>Altostratus</i>	As	middle (2-6 Km)		< 2
<i>Alto cumulus</i>	Ac	middle (2-6 km)		0
<i>Cirrostratus</i>	Cs	high (6-12 km)		0
<i>Cirrocumulus</i>	Cc	high (6-12 km)		0



# DYNAMIC METEOROLOGY

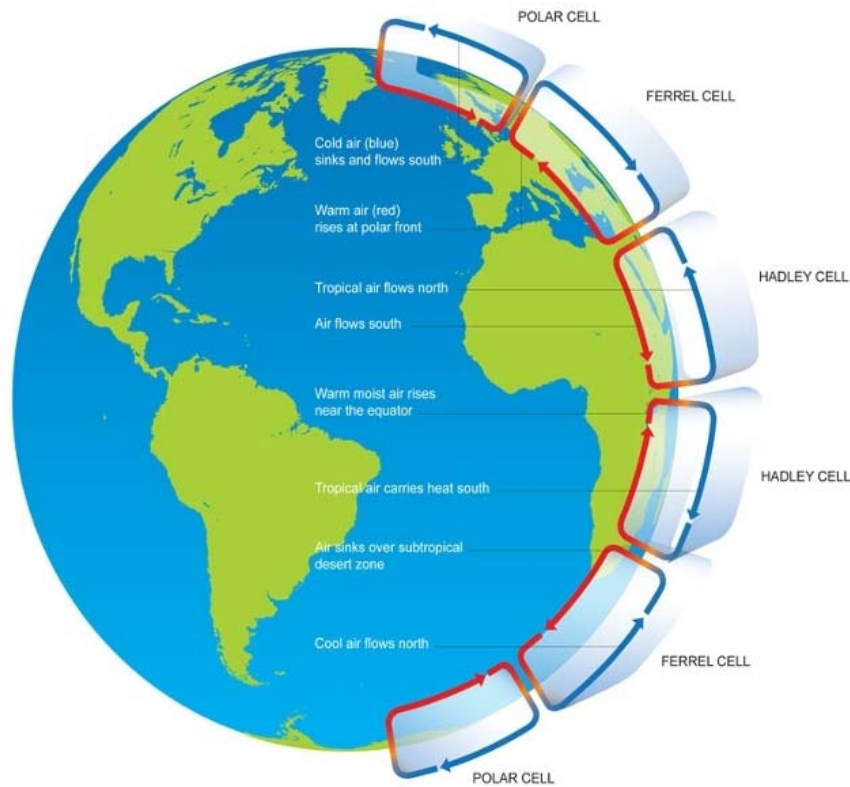
## Air mass flows and winds



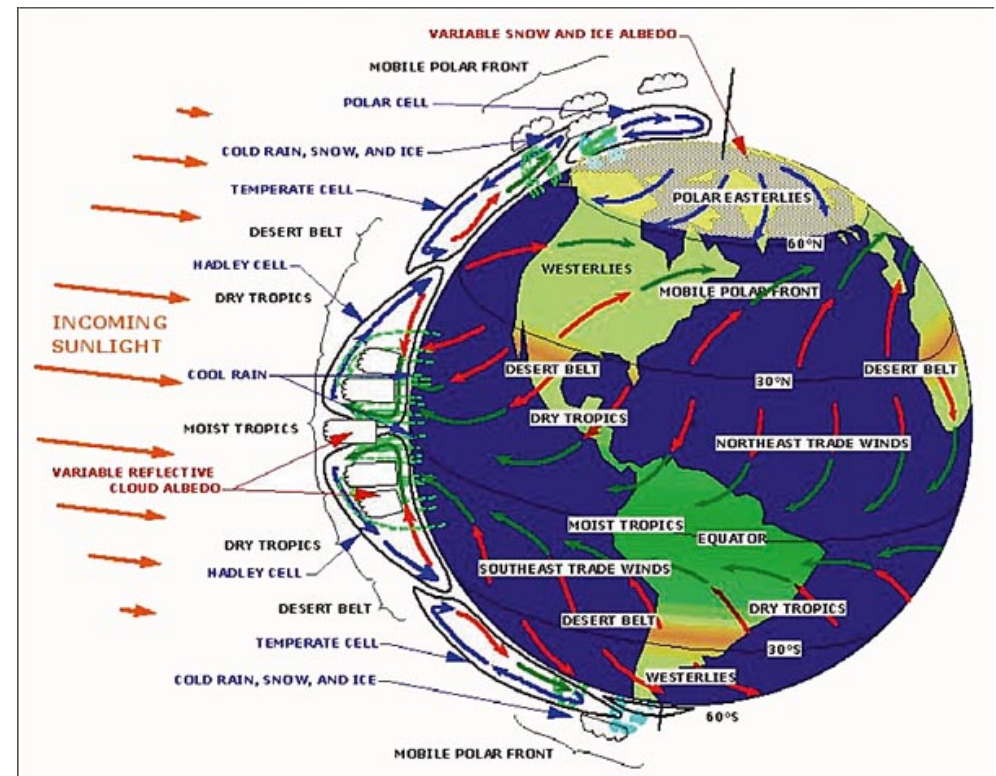


# DYNAMIC METEOROLOGY

## Circulating cells

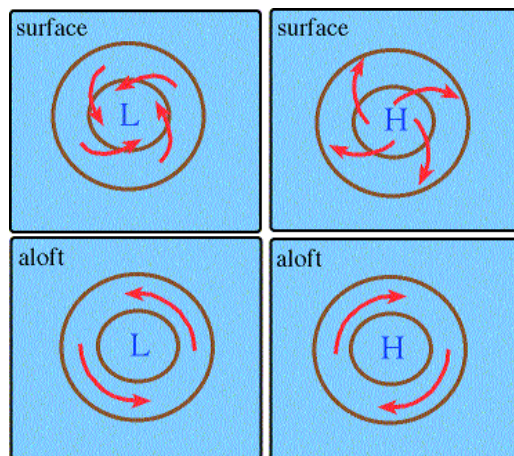
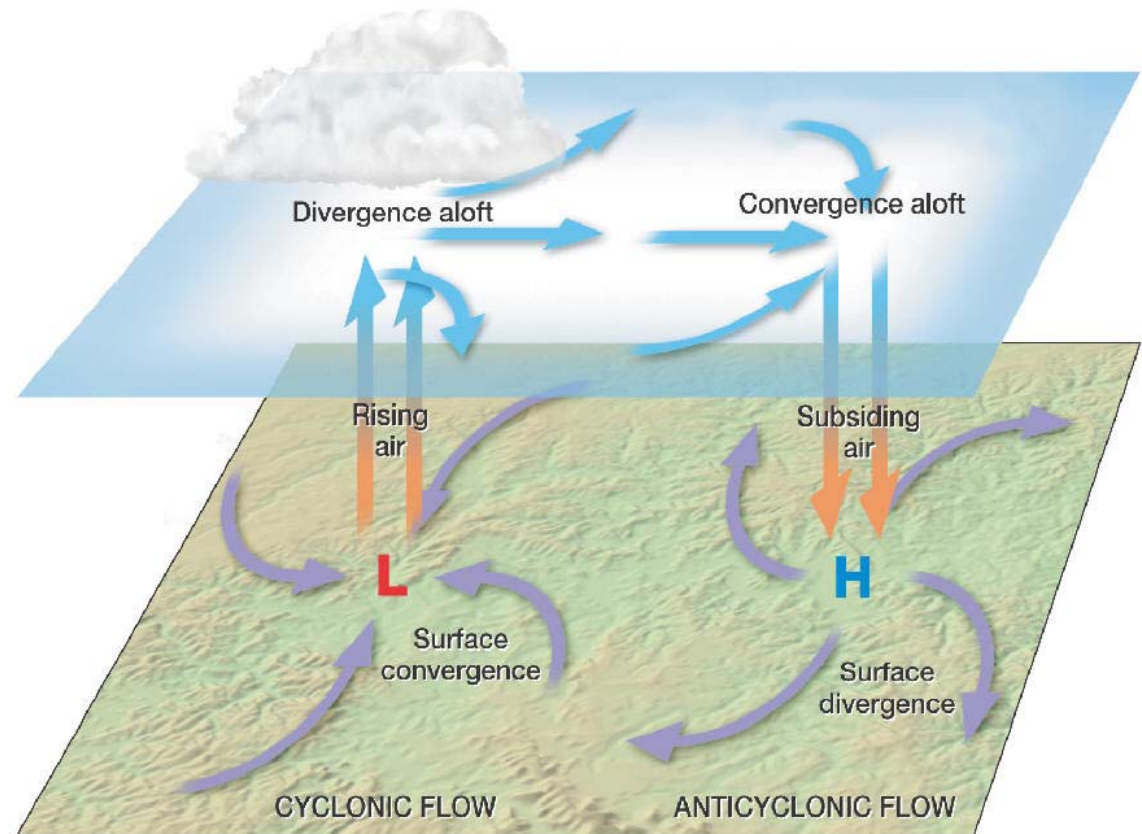
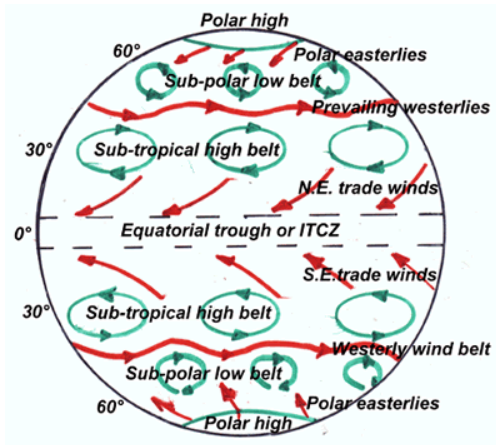


**CIRCULATING CELLS**  
 The Hadley cells have the most regular pattern of air movement, and produce extreme wet weather at the equator and extreme aridity on the deserts. The polar cells are the least well-defined.



# DYNAMIC METEOROLOGY

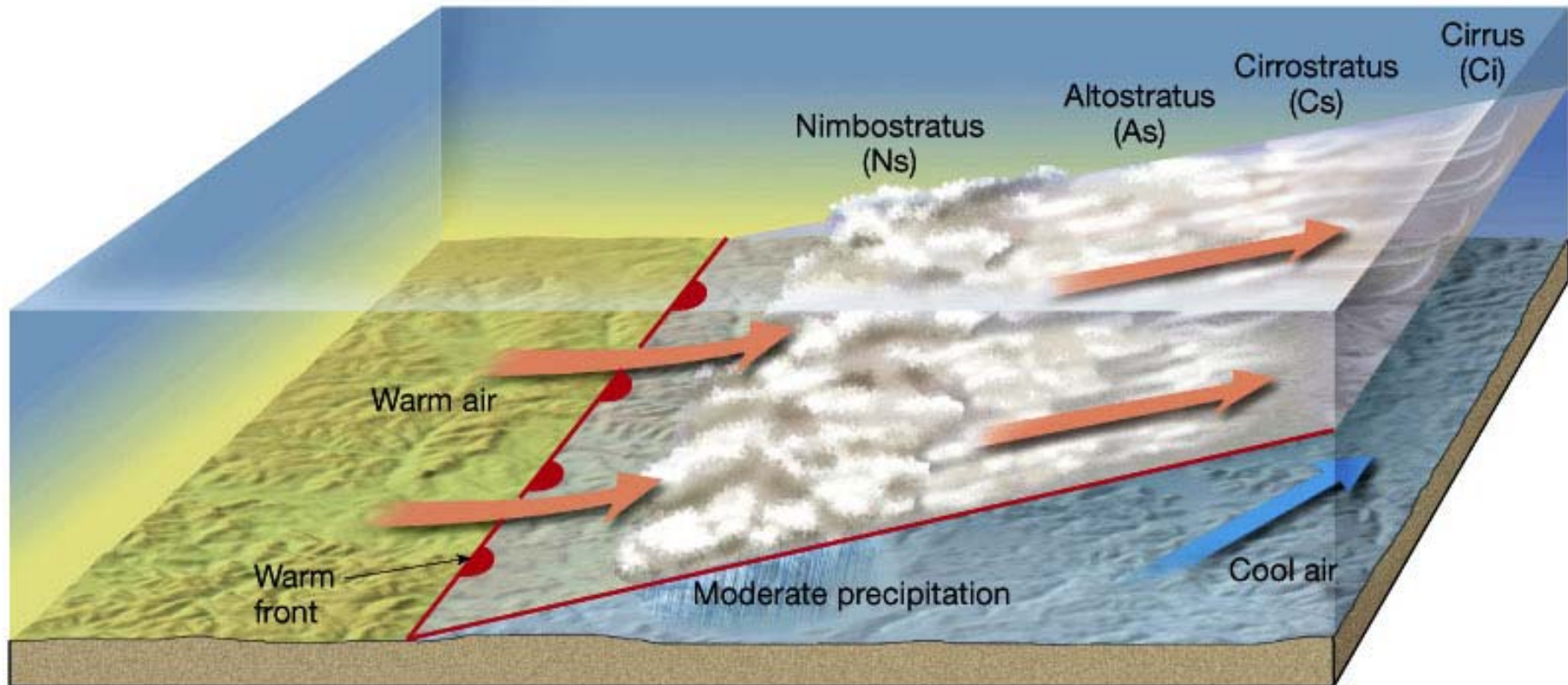
## Cyclones and anticyclones





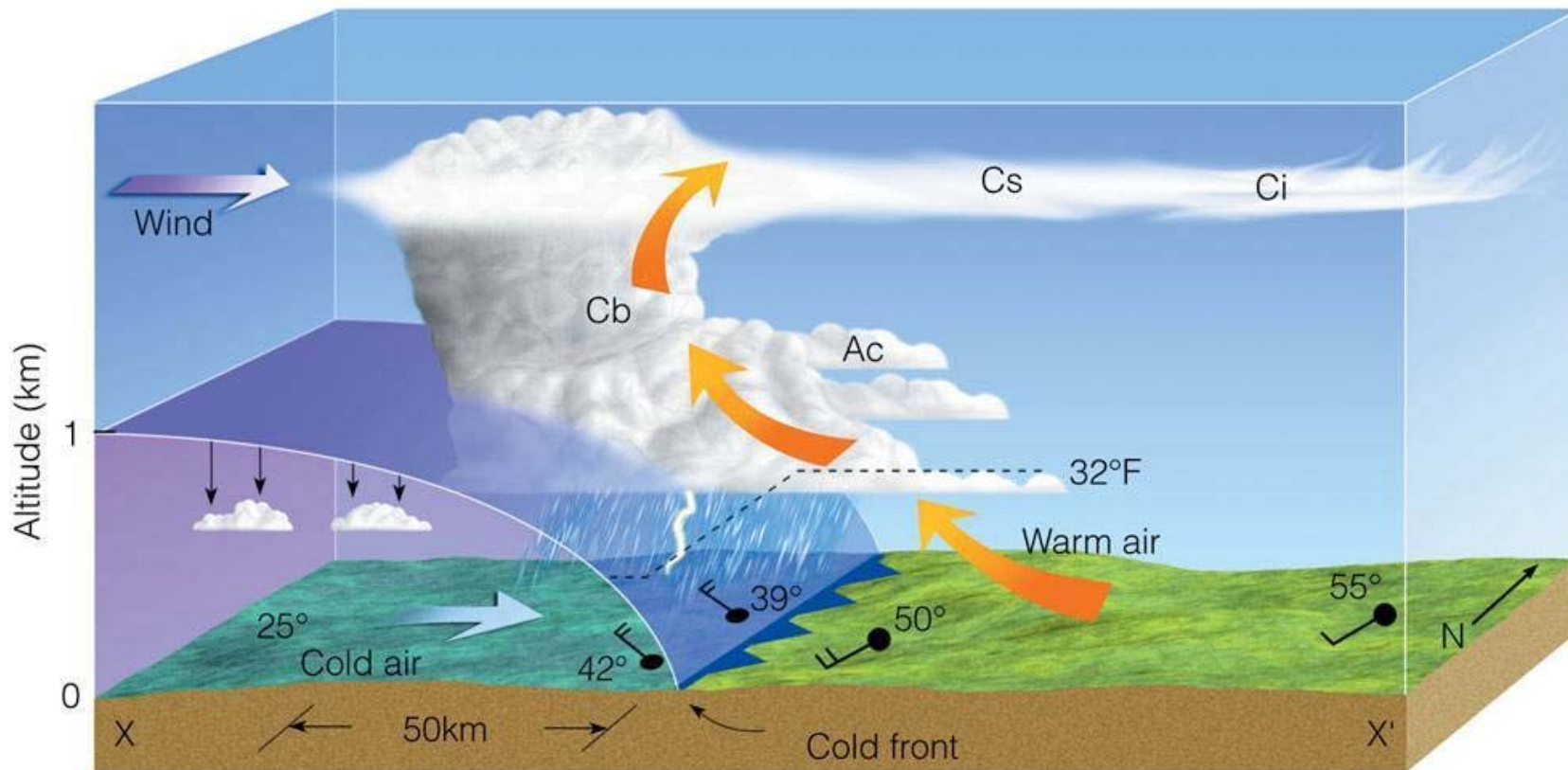
# CLOUD GENERA

## Warm fronts



# CLOUD GENERA

## Cold fronts

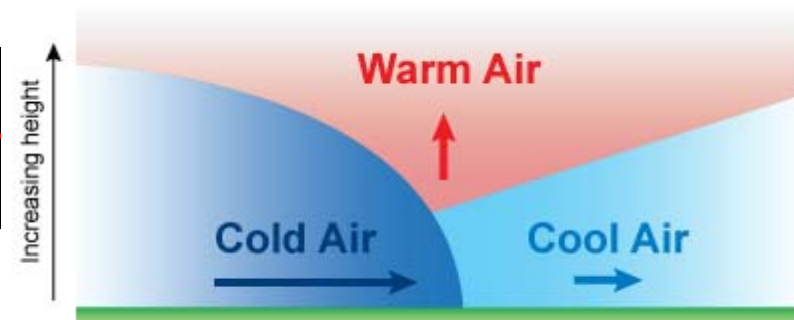
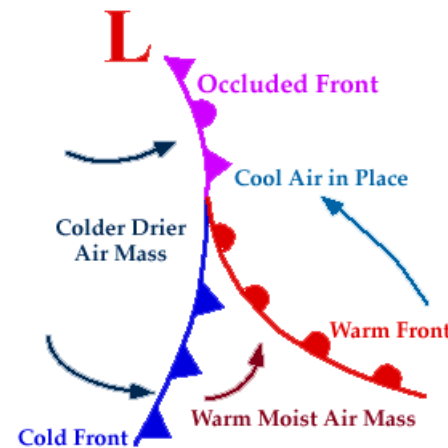
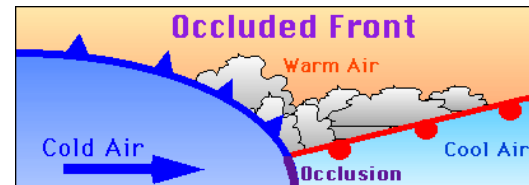
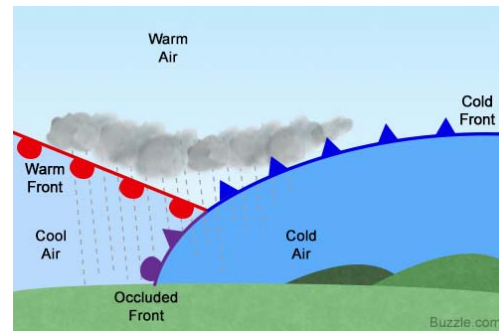
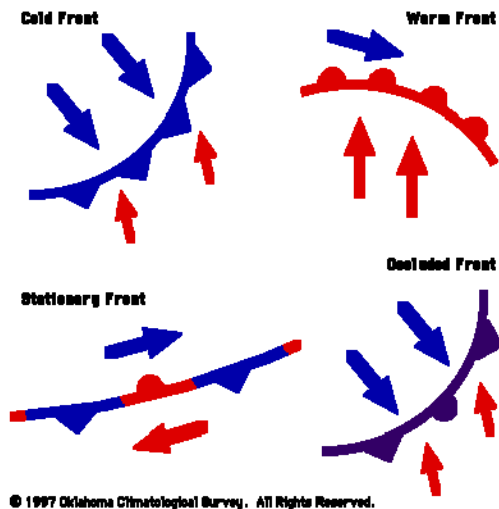


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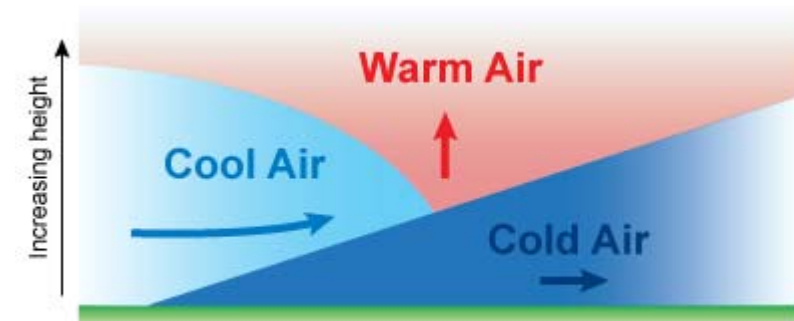


# CLOUD FRONTS

## Cold, warm and occluded



**Cold Occlusion**



**Warm Occlusion**

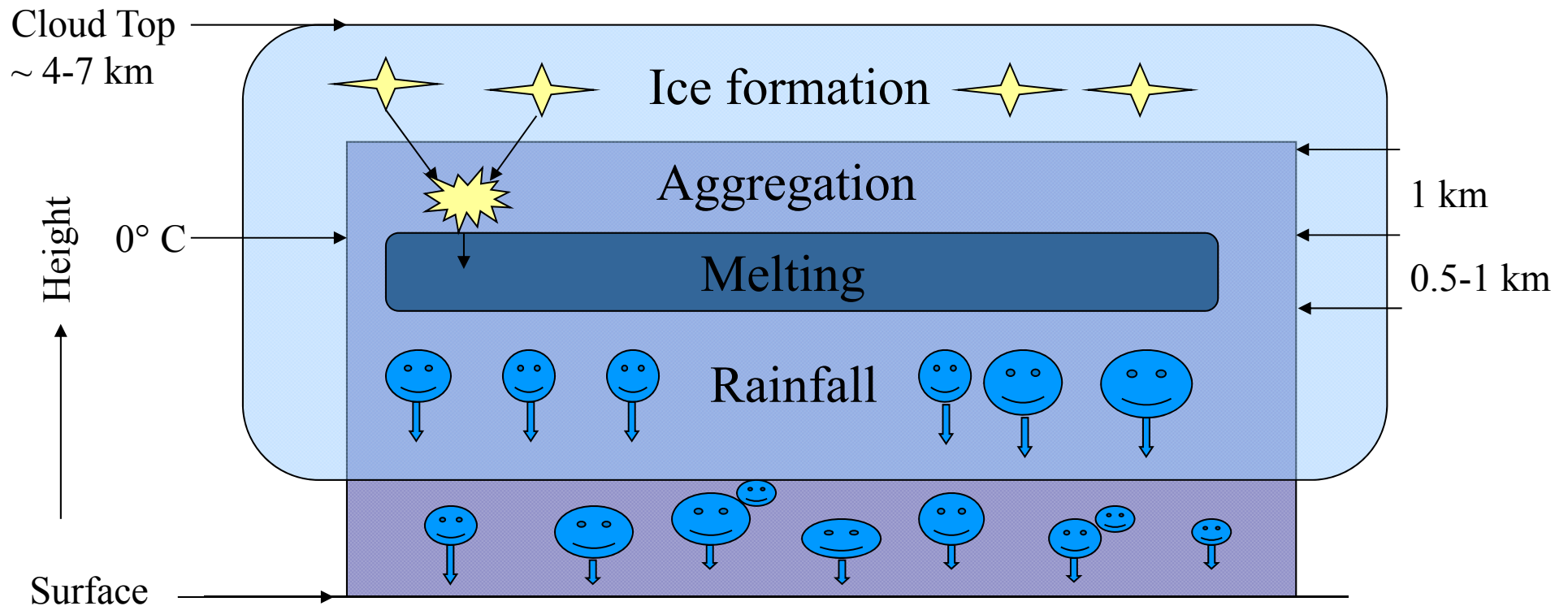
# PRECIPITATING CLOUDS

## Time and space scales

		CHARACTERISTIC TIME			
		1 day- 1 month	1 hour – 1 day	1 hour – 1 minute	
H O R I Z O N T A L  S C A L E	200 - 2000 km	- <i>Fronts</i> - <i>Hurricanes</i> (tens of days)			Meso- $\alpha$ scale
	20 - 200 km		- <i>Squall lines</i> - <i>Cloud clusters</i> - <i>Mountain and lake disturbances</i> (2 hours to 1 day)		Meso- $\beta$ scale
	2 - 20 km			- <i>Thunderstorms</i> (tens of minutes to hours)	Meso- $\gamma$ scale
	0.2 - 2 km			- <i>Tornadoes</i> - <i>Deep convection</i> (few minutes to 1 hour)	Micro- $\alpha$ scale
	0.02 - 0.2 km			- <i>Thermals</i> (few minutes)	Micro- $\beta$ scale
	0.001 - 0.02 km			- <i>Plumes</i> - <i>Turbulence</i> (less than a few minutes)	Micro- $\gamma$ scale

# PRECIPITATING CLOUDS

## Stratiform process

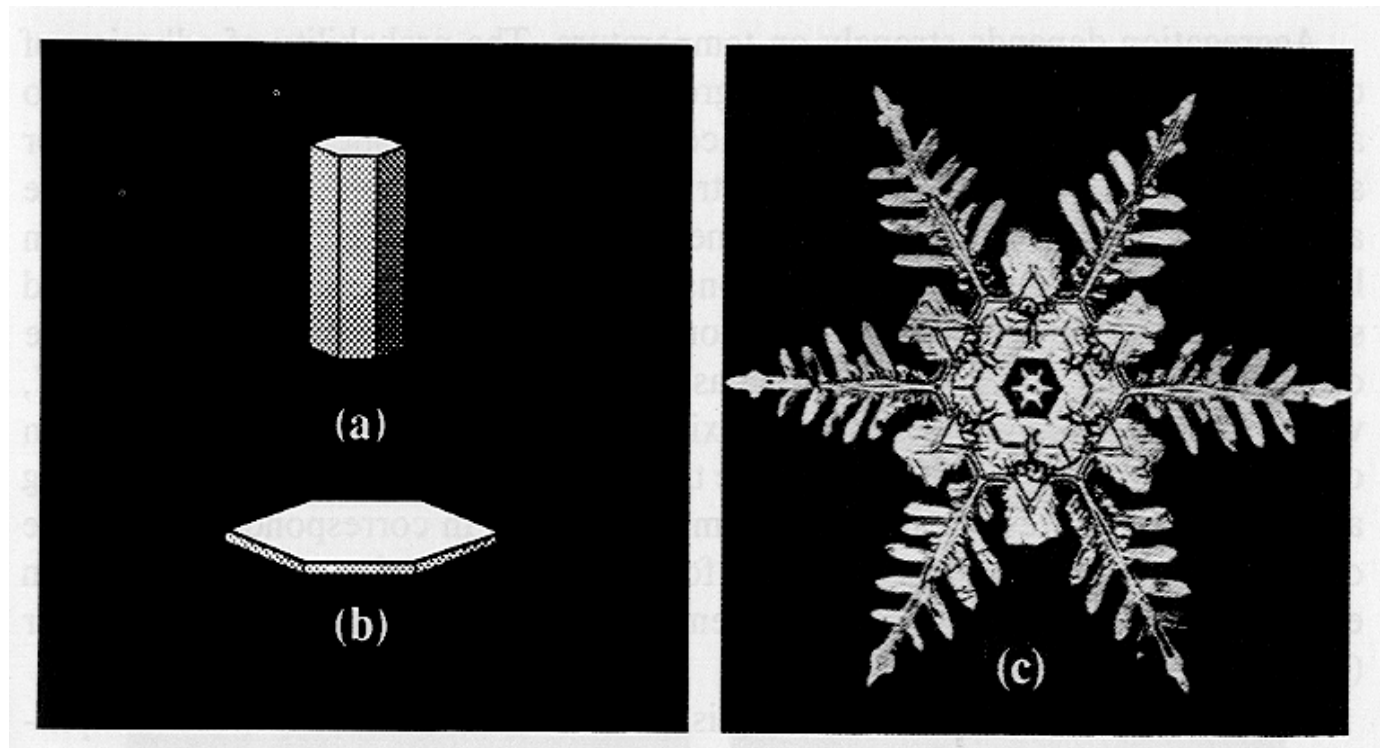


# PRECIPITATING CLOUDS

## Ice crystals

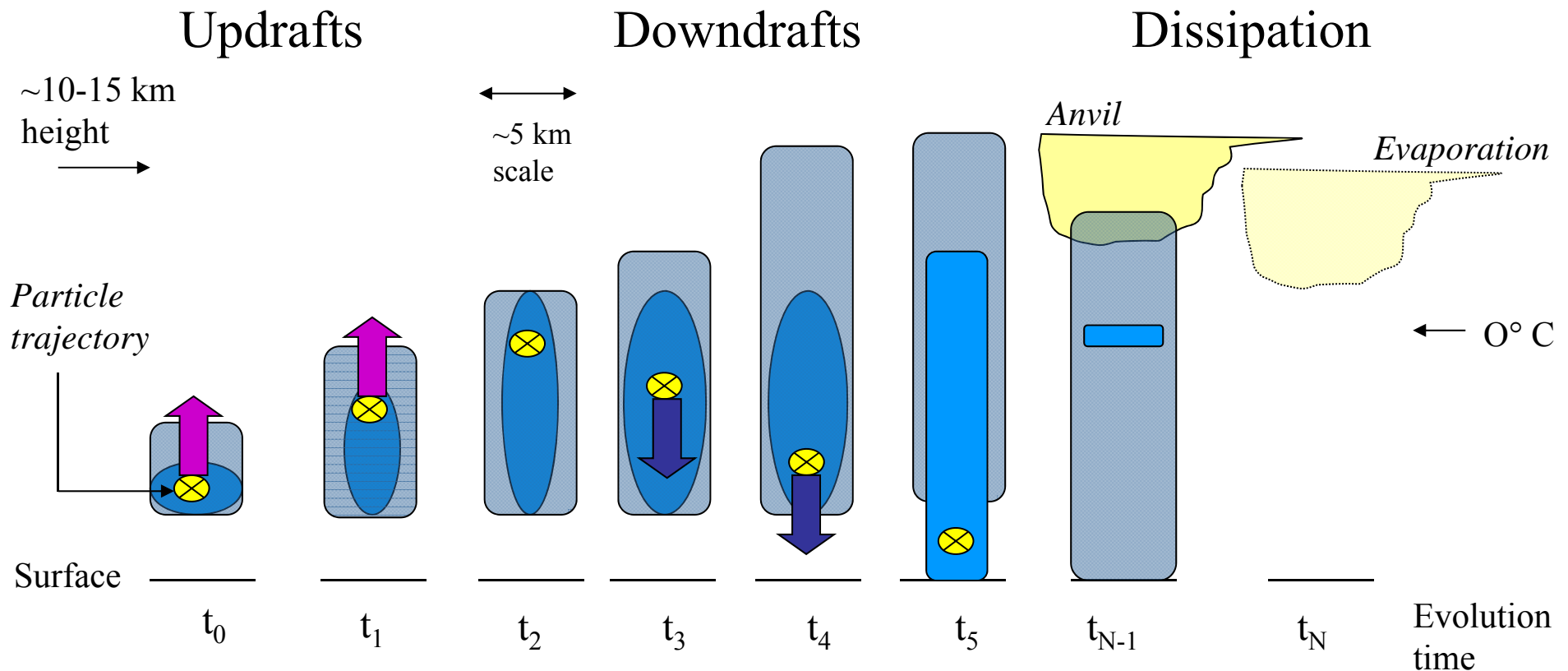
As for droplets (condensation/evaporation), ice crystals can initially grow by *vapor deposition* and evaporate by *sublimation*. The probability of presence of ice particles in a cloud increases with decreasing temperature below zero. At  $-10^{\circ}\text{C}$  the probability to find ice is around 50%, while at  $-20^{\circ}\text{C}$  it raises to 95%.

**Main ice crystal shapes are: (a) columnar or prismatic, (b) disc, (c) dendrite**



# PRECIPITATING CLOUDS

## Convective process



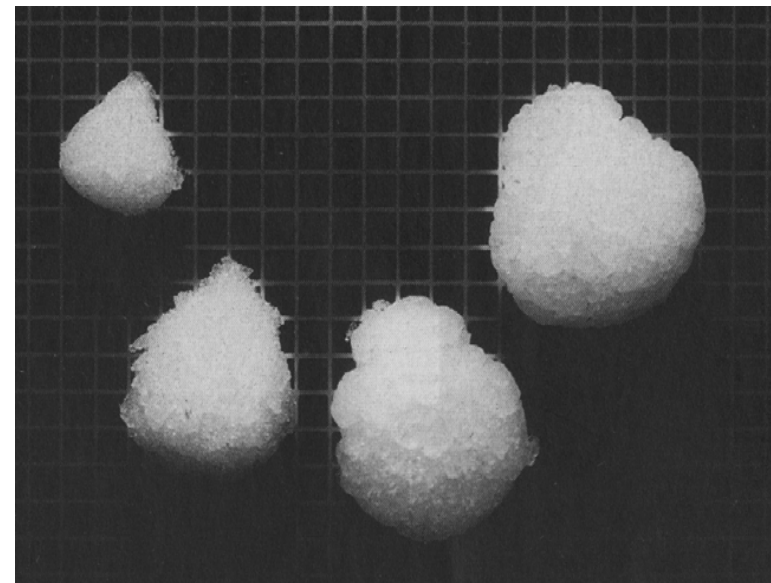
# PRECIPITATING CLOUDS

## Graupel particles

### GRAUPEL FORMATION BY RIMING

- Ice crystals fall through supercooled cloud droplets
- Supercooled droplets, that hit ice crystals, freeze to it
- Eventually the frozen droplet can hide the original shape
- Due to strong updrafts, hailstones can form

Natural graupels (boxes are 2 mm on a side)



	Rimed crystal		Hexagonal graupel	Conelike graupel
	Without tumbling (rotation)	Plane crystal	a <sub>1</sub>	a <sub>2</sub>
b <sub>1</sub>				b <sub>2</sub>
c <sub>1</sub>				c <sub>2</sub>
Columnar crystal		d <sub>1</sub>		d <sub>2</sub>
Radiating assemblage of plane branches		e <sub>1</sub>		e <sub>2</sub>
Frozen drop		f <sub>1</sub>		f <sub>2</sub>
With tumbling (rotation)	Rimed crystal		Lump graupel	
	Columnar crystal	g <sub>1</sub>	g <sub>2</sub>	x
	Radiating assemblage of plane branches	h <sub>1</sub>	h <sub>2</sub>	z
	Frozen drop	i <sub>1</sub>	i <sub>2</sub>	w

# RADAR APPLICATIONS

## Drop size distributions

### ➤ Modified Gamma distribution

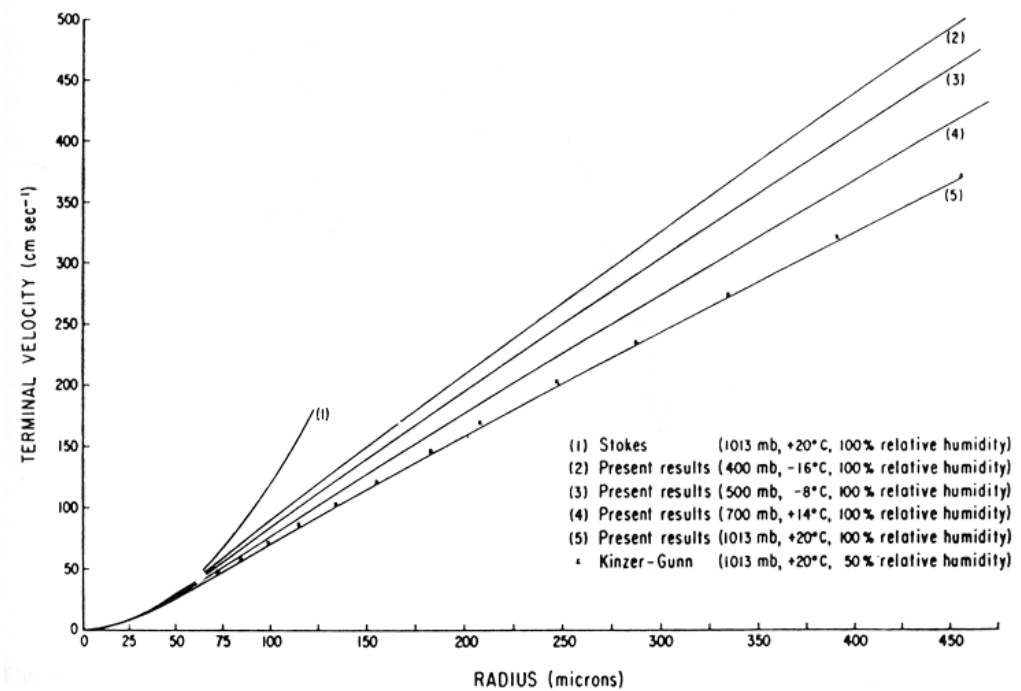
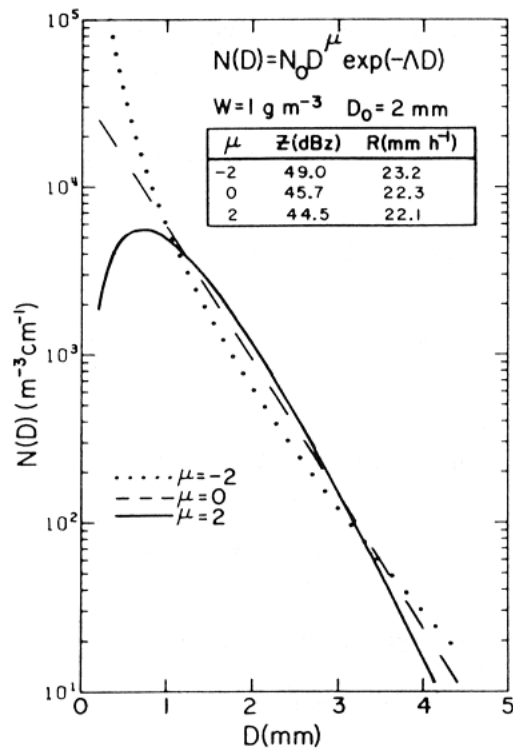
$$N(D) = N_0 D^\mu e^{-\Lambda D} \quad \text{with } D : \text{drop diameter; } N_0, \mu, \Lambda : \text{parameters}$$

- Moments of order  $n$ : 
$$m_n = \int_0^{\infty} D^n N(D) dD = \frac{N_0 \Gamma(n + \mu + 1)}{\Lambda^{n + \mu + 1}}$$
- Drop concentration [ $\#/m^3$ ]: 
$$N_T = m_0$$
  - Equivalent water content [ $g/m^3$ ]: 
$$M = (\pi/6) \rho m_3$$
  - Median volumetric diameter [m]: 
$$D_0 = M / 2$$
  - Precipitation intensity [mm/h]: 
$$R_\rho = (\pi/6) \rho \int_0^{\infty} D^3 N(D) [u_t(D) - w] dD \quad [\text{kg/m}^2 \text{s}]$$
  

$$u_t(D) = aD^b \Rightarrow R = R_\rho / \rho = (\pi/6) a m_{3+b} \quad [\text{m/s} = \text{mm/h}]$$

# RADAR APPLICATIONS

## Examples of size distributions





# RADAR APPLICATIONS

## Atmospheric attenuation

$$\alpha = k_a = 0.4343 \int_0^{\infty} \sigma_a(D) N(D) dD \quad [\text{dB/km}] \Rightarrow$$

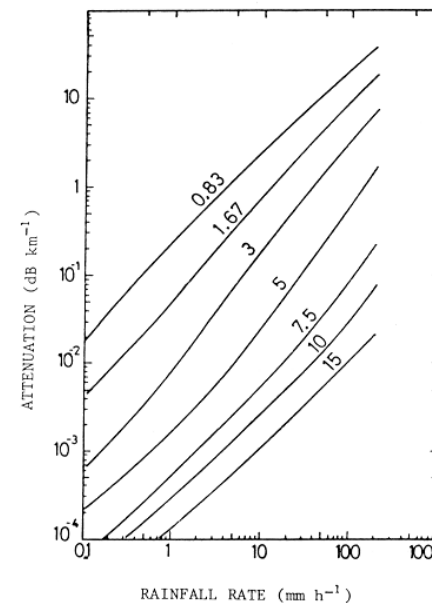
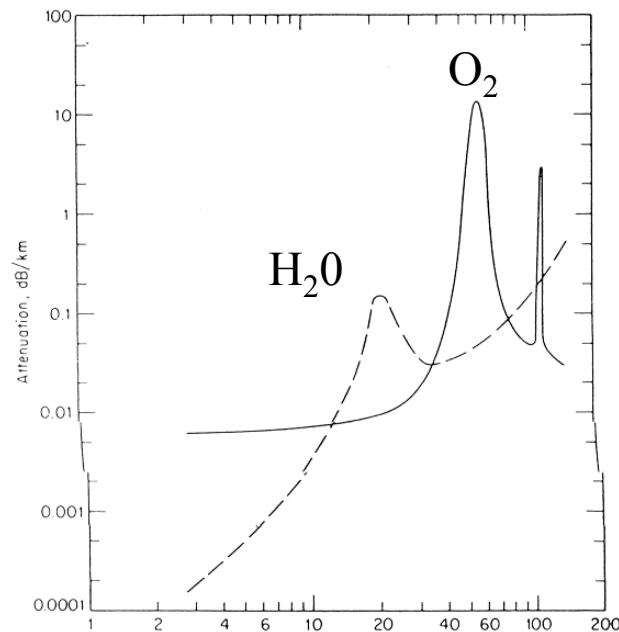
$\alpha_g$ : atmospheric gases

$\alpha_c$ : cloud and ice

$\alpha_p$ : rain and graupel (precipitating)

$\alpha_t = \alpha_g + \alpha_c + \alpha_p$

$$\alpha_p = k_p R$$



# RADAR APPLICATIONS

## Radar reflectivity

- Volumetric reflectivity of particles with different size

$$\eta = \frac{\sum_{i=1}^{N_{part}} \sigma_{bi}}{dV} \rightarrow \frac{\sum_{i=1}^{N_D} \sigma_b(D_i) N_p(D_i)}{dV} = \sum_{i=1}^{N_D} \sigma_b(D_i) N(D_i) \cong \int_0^{\infty} \sigma_b(D) N(D) dD \quad \left[ \frac{1}{m} \right]$$

- Rayleigh reflectivity and reflectivity factor

$$Se D \ll \lambda : \quad \eta = \int_0^{\infty} \left[ \frac{\pi^5 |K|^2}{\lambda^4} D^6 \right] N(D) dD \Rightarrow Z \equiv \int_0^{\infty} D^6 N(D) dD$$

- Mie reflectivity and equivalent reflectivity factor

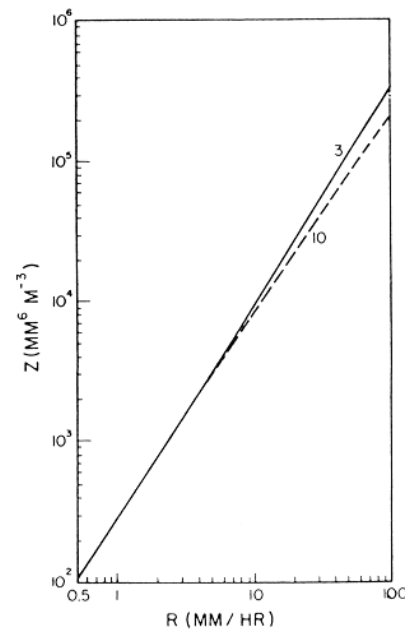
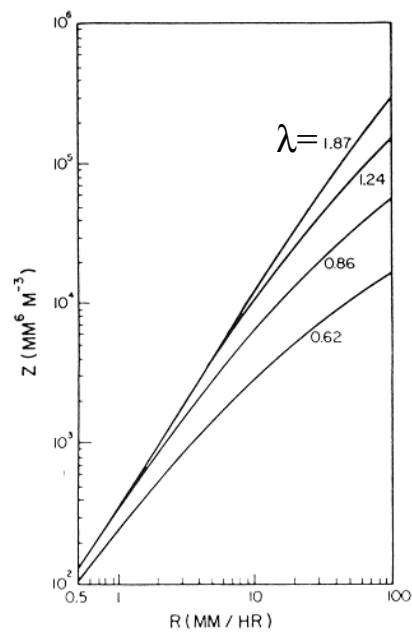
$$Se D > \lambda : \quad \eta = \frac{\pi^5 |K|^2}{\lambda^4} Z_e \Rightarrow Z_e \equiv \frac{\lambda^4}{\pi^5 |K|^2} \int_0^{\infty} \sigma_b(D) N(D) dD$$

- Dielectric factor K
 

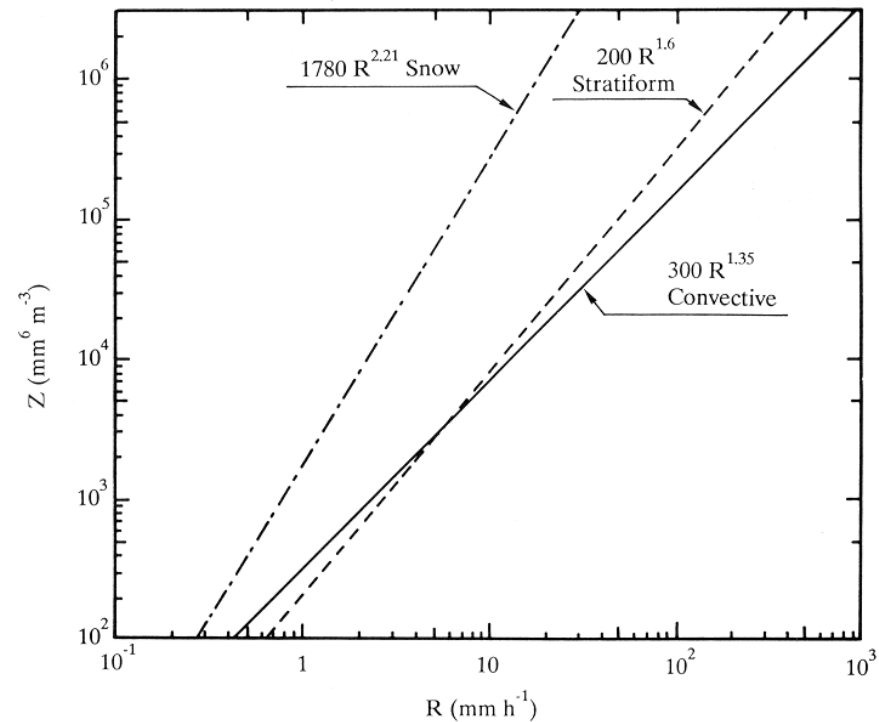
Water	$ K ^2=0.930$
Ice	$ K ^2=0.176$
Melt	$ K ^2=0.208$

# RADAR APPLICATIONS

## Examples of radar reflectivity



**Power-law relation:  $Z = aR^b$**



# RADAR APPLICATIONS

## Hydrologic relationships

<i>Parameter</i>	<i>Stratiform</i> <i>[Marshall and Palmer,</i> <i>1948]</i>	<i>Thunderstorm</i> <i>[Sekhon and</i> <i>Srivastava, 1971]</i>	<i>Snow</i> <i>[Sekhon and</i> <i>Srivastava,</i> <i>1970]</i>
<i>Z-R</i>	$Z = 200 R^{1.6}$	$Z = 300 R^{1.35}$	$Z = 1780 R^{2.21}$
<i>M-R</i>	$M = 0.072 R^{0.88}$	$M = 0.052 R^{0.94}$	$M = 0.250 R^{0.86}$
<i>D<sub>0</sub>-R</i>	$D_0 = 0.09 R^{0.21}$	$D_0 = 0.13 R^{0.14}$	$D_0 = 0.14 R^{0.45}$
<i>Γ-R</i>	$\Gamma = 41 R^{-0.21}$	$\Gamma = 38 R^{-0.14}$	$\Gamma = 22.9 R^{-0.45}$
<i>N<sub>0</sub>-R</i>	$N_0 = 0.08$	$N_0 = 0.07 R^{0.37}$	$N_0 = 0.025 R^{-0.94}$
<i>Γ-M</i>	$\Gamma = 22 M^{-0.24}$	$\Gamma = 25 M^{-0.15}$	$\Gamma = 11 M^{-0.52}$

*Note:*

*Z* is in  $\text{mm}^6 \text{m}^{-3}$ , *R* is in  $\text{mm h}^{-1}$ , *M* is in  $\text{g m}^{-3} \text{m}^{-3}$ , *D<sub>0</sub>* is in cm, *Γ* is in  $\text{cm}^{-1}$ , *N<sub>0</sub>* is in  $\text{cm}^{-4}$ .

# RADAR APPLICATIONS

## Polarized measurements

### ➤ Backscattering matrix

- The scattering field is related to the incident one by:

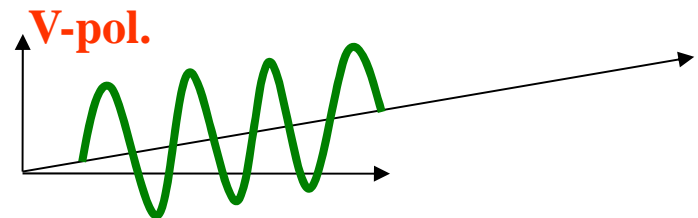
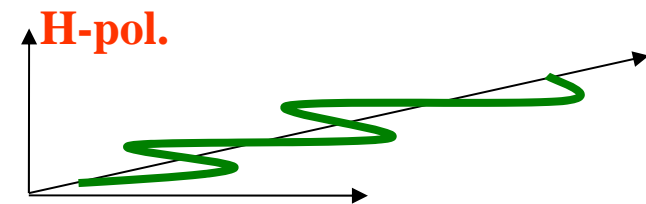
$$\begin{bmatrix} E_h^s \\ E_v^s \end{bmatrix} = \frac{e^{-j\beta r}}{r} \begin{bmatrix} S_{HH} & S_{HV} \\ S_{VH} & S_{VV} \end{bmatrix} \begin{bmatrix} E_h^i \\ E_v^i \end{bmatrix}$$

- The S matrix is the complex backscattering matrix, whose terms are:

$$S_{pp} = |S_{pp}| e^{j\varphi_{pp}}$$

### ➤ Interpretation and Symmetry

- Diagonal terms  $S_{HH}$  and  $S_{VV}$  represent linear co-polarized components
  - $S_{HH}$  : H transmission, H reception
- Off-diagonal terms are cross-polarized components
  - $S_{VH}$  : H transmission, V reception
- For a spherical particles, the cross-polarized terms are zero.
- The cross-pol. terms are equal for reciprocity (unless medium is anisotropic).



# RADAR APPLICATIONS

## Polarimetric definitions

### ➤ Polarimetric form of radar equation

- Radar equation can be generalized to polarimetric form
- Complex envelope  $V$  will be polarization dependent, i.e.  $V_{pp}$

### ➤ Co-polar radar reflectivity ( $\text{mm}^6\text{m}^{-3}$ )

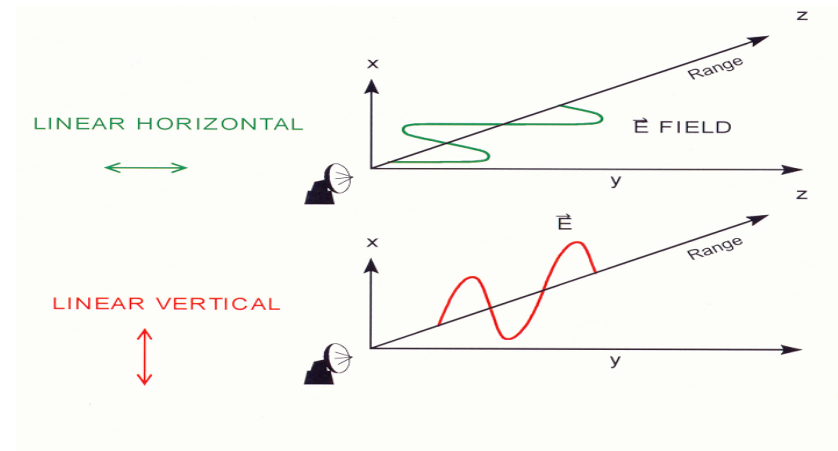
$$Z_{HH} = (\lambda^4 / \pi^5 |K|^2) \langle |S_{HH}|^2 \rangle, \quad Z_{VV} = (\lambda^4 / \pi^5 |K|^2) \langle |S_{VV}|^2 \rangle$$

### ➤ Differential reflectivity

$$Z_{DR} = 10 \text{Log} \frac{\langle |S_{HH}|^2 \rangle}{\langle |S_{VV}|^2 \rangle}$$

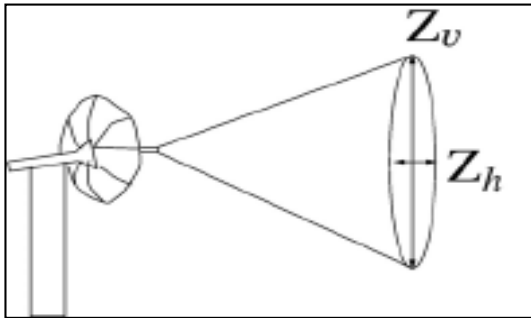
### ➤ Linear Depolarization Ratio (LDR)

$$LDR = 10 \text{Log} \frac{\langle |S_{HV}|^2 \rangle}{\langle |S_{HH}|^2 \rangle}$$



# RADAR APPLICATIONS

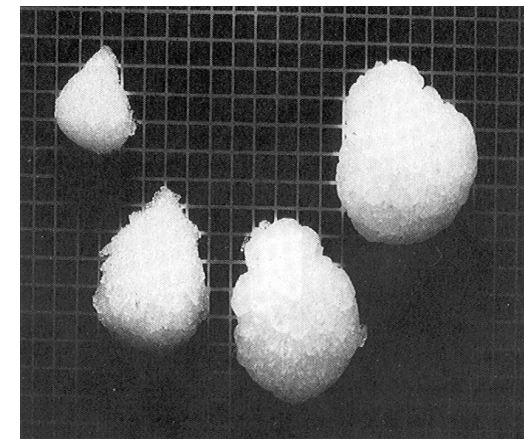
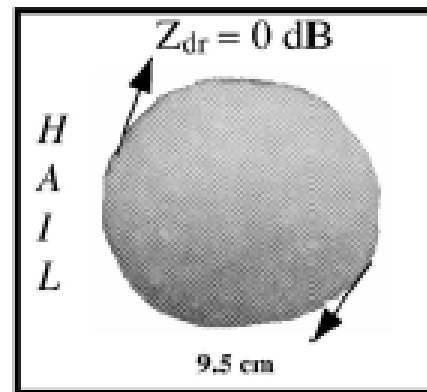
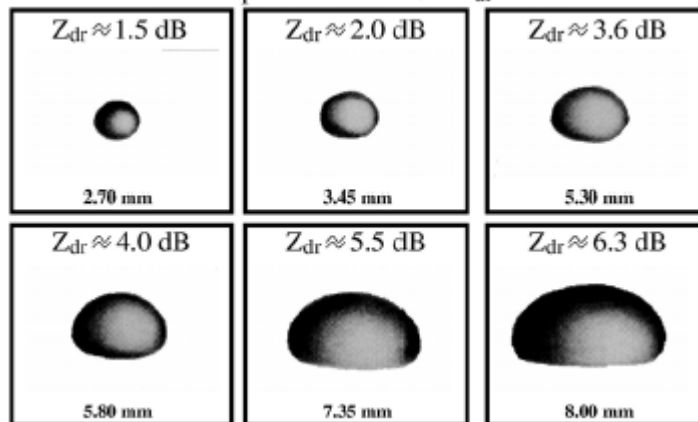
## Polarimetric examples



$$Z_{DR} = 10 \text{Log} \frac{\langle |S_{HH}|^2 \rangle}{\langle |S_{VV}|^2 \rangle}$$

Oblate  $\rightarrow Z_{DR} > 0$   
 Sphere  $\rightarrow Z_{DR} = 0$   
 Prolate  $\rightarrow Z_{DR} < 0$

Raindrops  $< 0.3 \text{ mm} \rightarrow Z_{dr} = 0 \text{ dB}$



# RADAR APPLICATIONS

## Polarimetric definitions

### ➤ Correlation coefficient

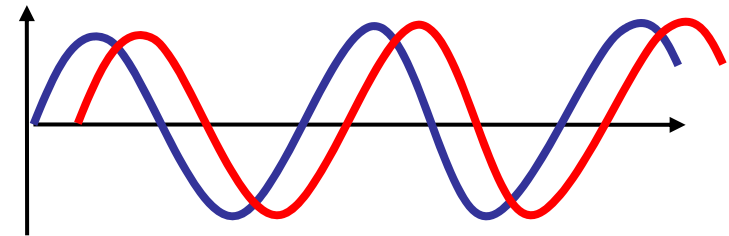
$$\rho_{HV} = \frac{\langle S_{VV} S_{HH}^* \rangle}{\sqrt{\langle |S_{HH}|^2 \rangle \langle |S_{VV}|^2 \rangle}} = |\rho_{hv}| e^{j\delta_{hv}}$$

### ➤ Differential propagation-phase shift (°/km)

$$K_{DP} = \frac{180}{\pi} \lambda \operatorname{Re} \left[ \int_0^D [f_{HH}(D) - f_{VV}(D)] N(D) dD \right]$$

$$\phi_{DP} = 2 \int_0^r K_{DP}(r') dr' + \delta_{hv}$$

$$\delta_{hv} = \arg(\langle S_{VV} S_{HH}^* \rangle)$$



- Oblate particles tend to slow down **H-pol.** waves with respect to **V-pol.** ones
- This effect is larger for liquid than ice particles
- $\delta$  is due to particle-induced phase shift



# RADAR APPLICATIONS

## Polarimetric definitions

- Specific attenuation (1/km)

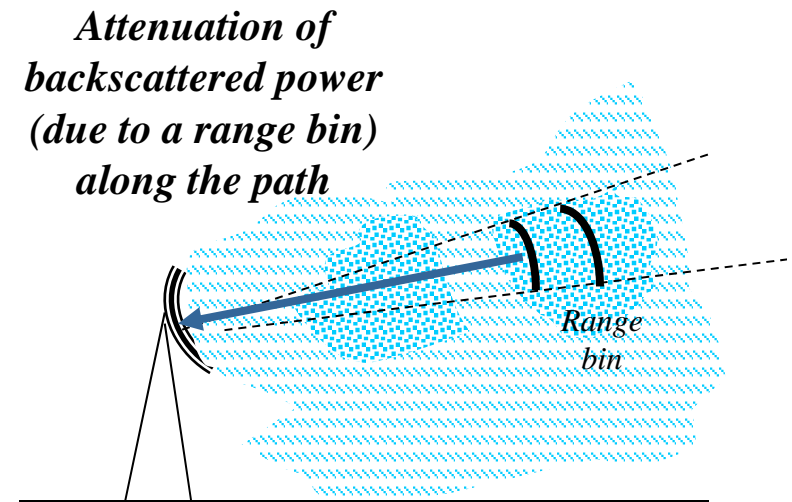
$$\alpha_{HH} = 2 \cdot 10^{-3} \frac{180}{\pi} \lambda \operatorname{Im} \left[ \langle 4\pi S_{HH}^f(D_e, \phi) \rangle \right]$$

- One-way path attenuation

$$A_{HH}(r) = \int_0^r \alpha_{HH}(r') dr'$$

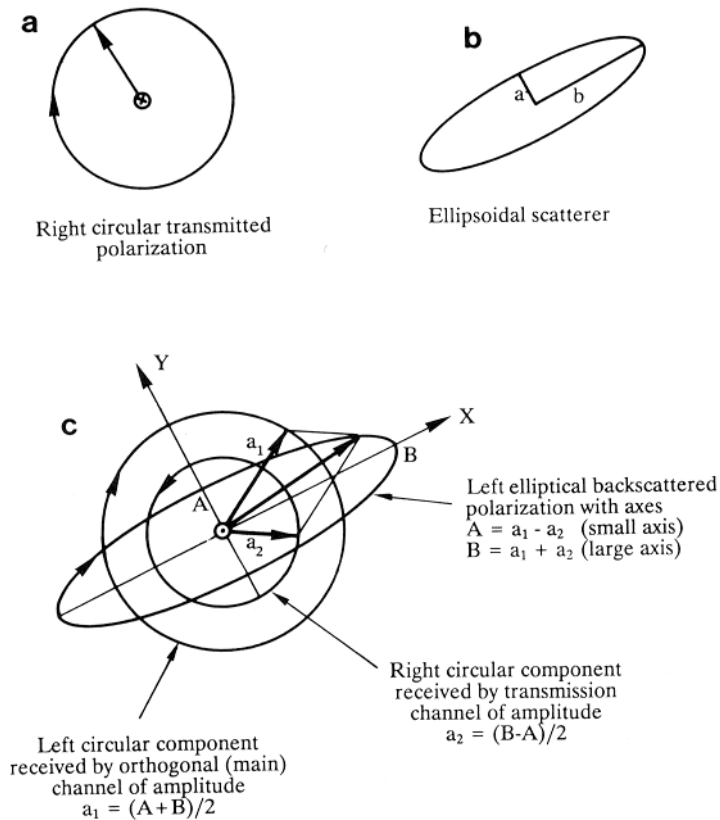
- Application

$$\begin{aligned} \langle W_{Rh} \rangle &= C_2 L_{HH}^2(r) \frac{Z_{HH}(r)}{r^2} = \\ &= C_2 \left( e^{-2A_{HH}(r)} \right) \frac{Z_{HH}(r)}{r^2} \end{aligned}$$

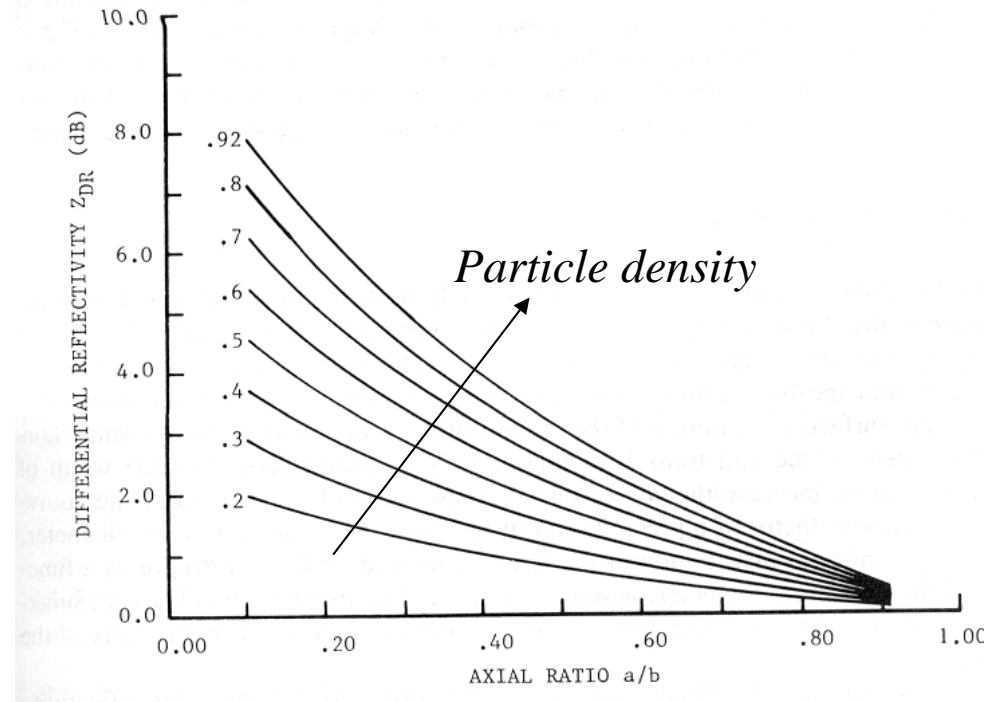


# RADAR APPLICATIONS

## Polarimetric theory



**Differential reflectivity vs. ellipsoid axial ratio**



# RADAR APPLICATIONS

## Polarimetric features

Table 1. Values of polarimetric variables for precipitation types (from Doviak and Zrnić 1993).

	$Z_h$ (dBZ)	$Z_{DR}$ (dB)	$\rho_{hv}$	$K_{DP}$ ( $^{\circ}$ km $^{-1}$ )	LDR (dB)
Drizzle	< 25	0	> 0.99	0	< -34
Rain	25 to 60	.5 to 4	> 0.97	0 to 10	-27 to -34
Dry snow	< 35	0 to .5	> 0.99	0 to 0.5	< -34
Dense snow	< 25	0 to 5	> 0.95	0 to 1	-25 to -34
Wet snow	< 45	0 to 3	0.8 to 0.95	0 to 2	-13 to -18
Dry graupel	40 to 50	-0.5 to 1	> 0.99	-0.5 to 0.5	< -30
Wet graupel	40 to 55	-0.5 to 3	> 0.99	-0.5 to 2	-20 to -25
Wet hail (< 2 cm)	50 to 60	-0.5 to 0.5	> 0.95	-0.5 to 0.5	< -20
Wet hail (> 2 cm)	55 to 70	< -0.5	> 0.96	-1 to 1	-10 to -15
Rain/hail	50 to 70	-1 to 1	> 0.90	0 to 10	-10 to -20

# RADAR APPLICATIONS

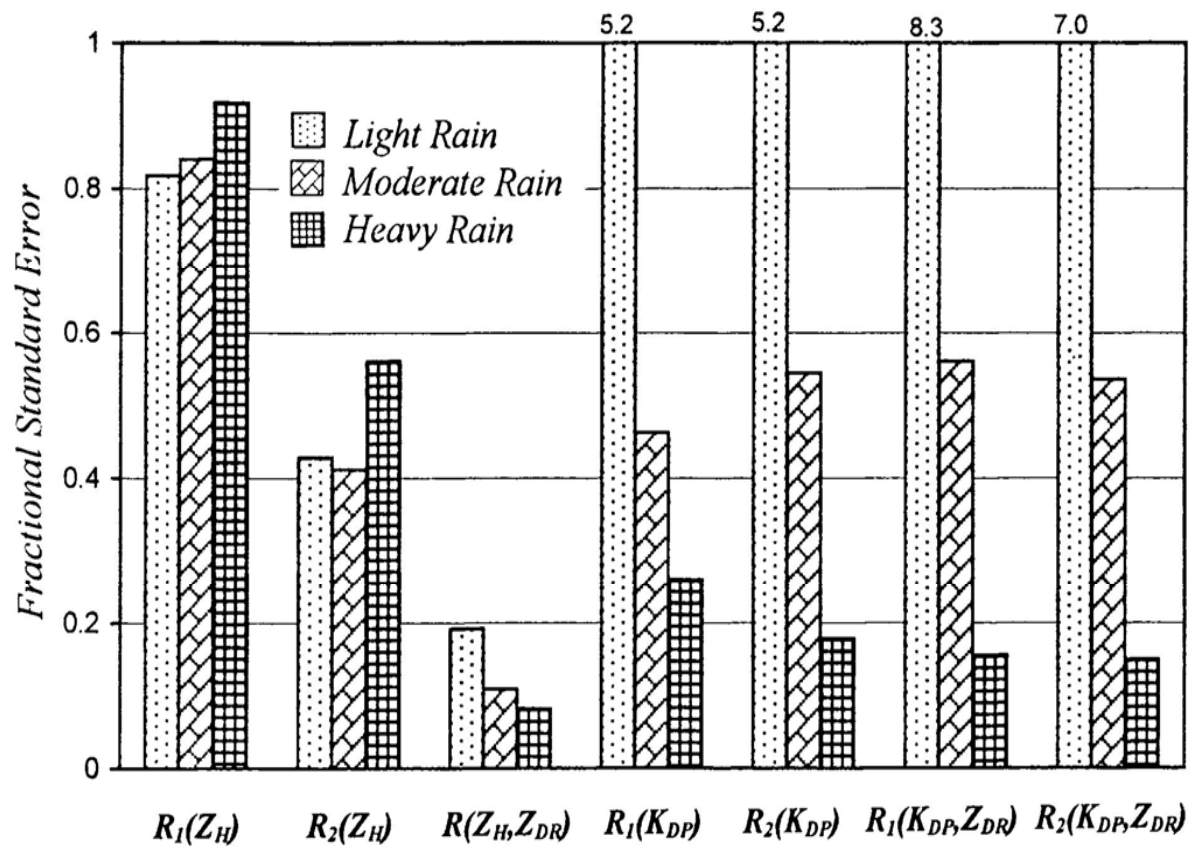
## Polarimetric advantages

TABLE 1. Attributes of polarimetric variables (for 5- and 10-cm wavelengths).

Attribute Variable	Independent of absolute radar	Immune to propagation effects calibration	Immune to noise bias	Used for quantitative estimation	Independent of concen- tration
$Z_h$	no	no	no	yes	no
$Z_{DR}$	yes	no	no	yes	yes
$K_{DP}$	yes	yes	yes	yes	no
$\rho_{hv}$	yes	yes	no	no	yes
$\delta$	yes	no	yes	no	yes
LDR	yes	no	no	no	yes

# RADAR APPLICATIONS

## Rainfall polarimetric estimators



**From single-pol.**

$$R = R(Z_h, Z_{dr})$$

**To dual-pol. retrieval**

$$R = R(Z_h, Z_{dr})$$

$$R = R(K_{dp}, Z_{dr})$$

$$R = R(K_{dp})$$

# RADAR APPLICATION

## Measurement issues

### Dual-pol. radar equation

$$\langle W_{Rh,v}(r, \theta, \varphi) \rangle = \frac{1}{M_{pulse}} \sum_{i=1}^{M_{pulse}} W_{iRh,v}(r, \theta, \varphi) - N_R$$

$$Z_{HH,VV}(r, \theta, \varphi) = C_2 r^2 \langle W_{Rh,v} \rangle (r, \theta, \varphi) e^{2A_{HH,VV}(r, \theta, \varphi)}$$

where  $P_R$ : received power at  $h$  or  $v$  pol.

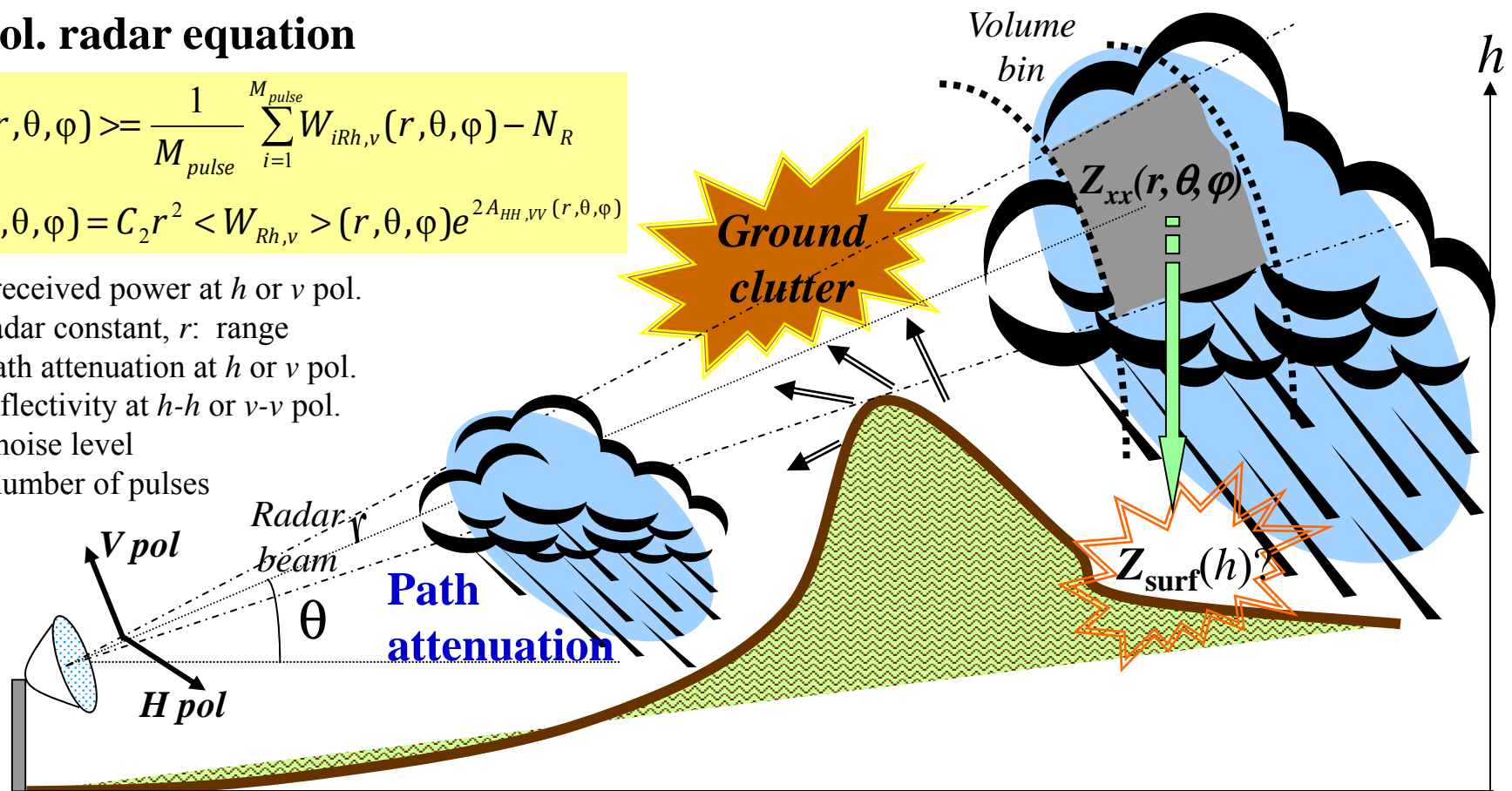
$C$ : radar constant,  $r$ : range

$A$ : path attenuation at  $h$  or  $v$  pol.

$Z$ : reflectivity at  $h-h$  or  $v-v$  pol.

$N_R$ : noise level

$M$ : number of pulses



# RADAR APPLICATIONS

## Operational issues

### ➤ Calibration objective

- Determine with precision and accuracy the instrumental constant  $C$  in radar equation equation
  - Transmitter power, loss of radiofrequency chain (e.g., guides, mixers, couplers)
  - Noise level in the receiver chain, receiver gain and linear dynamics (no distortions)

### ➤ Calibration techniques

- Internal calibration: measurement of each radar component and module
  - Suitable instrumentation (e.g., synthetic pulse generator, network analyzer)
  - Periodic controls (antenna, TX) vs. frequent controls (oscillators, noise)
- External calibration: measurement of reference radar targets
  - Large metallic spheres such that  $\sigma_b$  is equal to the geometric section
  - Rain-gauge time-series data
  - Difficult measurement set up and ambiguities
- Overall calibration obtainable:  $\geq 1$  dB

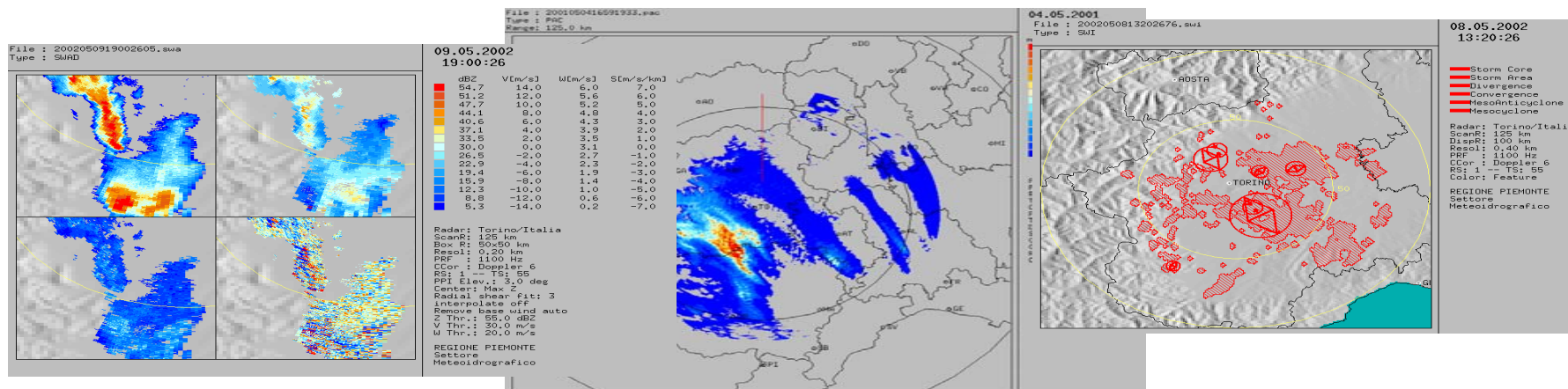
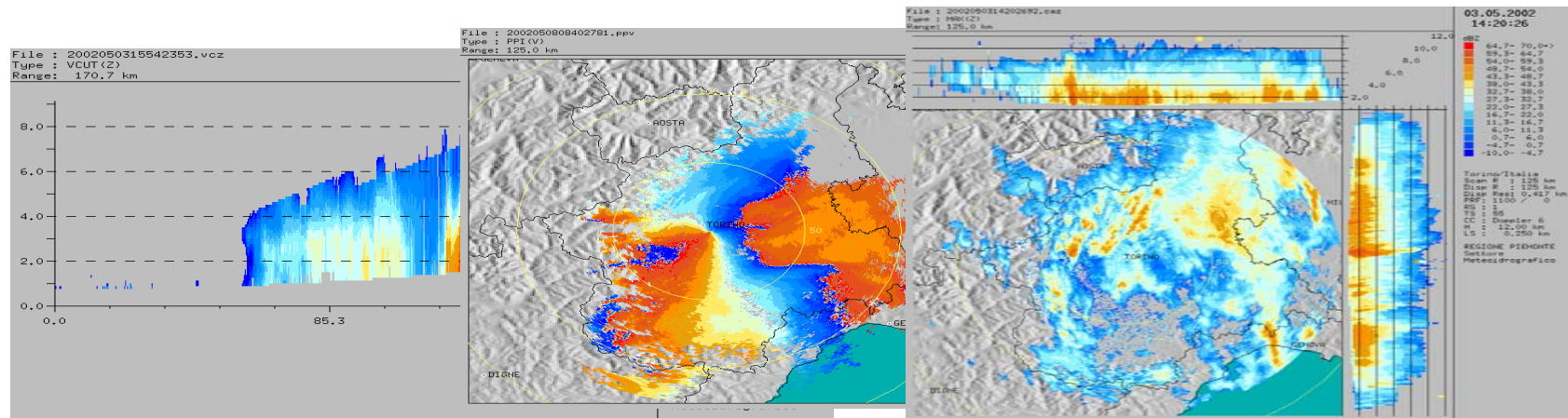
### ➤ Ground-clutter contamination

- Spurious echoes backscattered by ground and obstacles
- Spectral, statistical and static removal techniques
- Coupling with beam occultation, anomalous propagation, reflectivity gradients



# RADAR PRODUCTS

## An overview

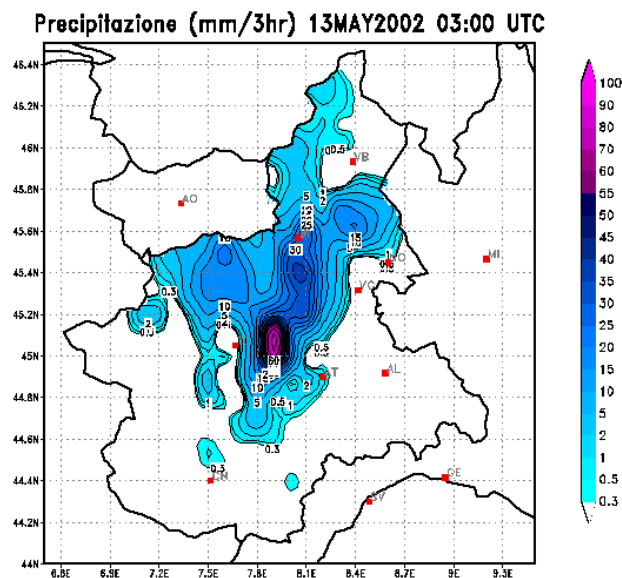




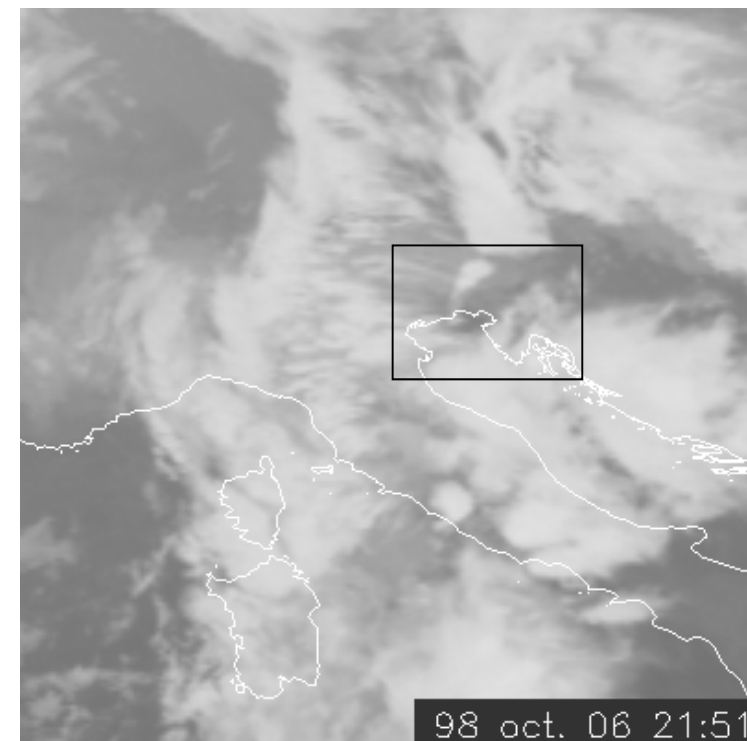
# RADAR PRODUCTS

## Comparison among instruments

### Rain gauges



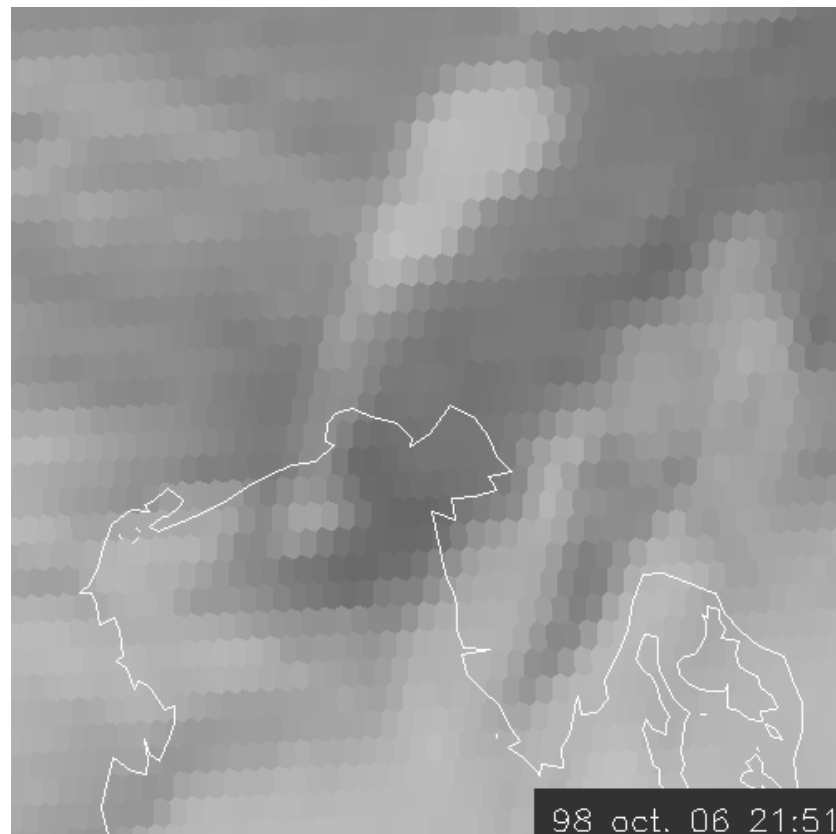
### Meteosat IR



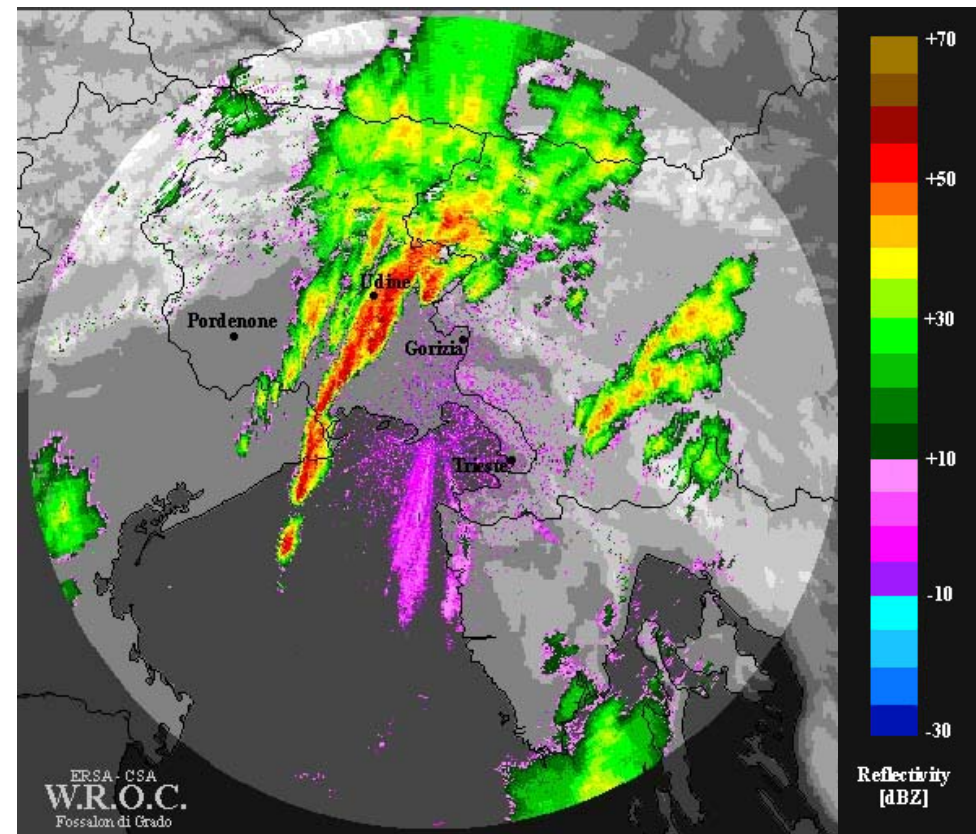
# RADAR PRODUCTS

## Comparison among instruments

**Meteosat IR**



**C-band Radar (ARPA-FVG)**



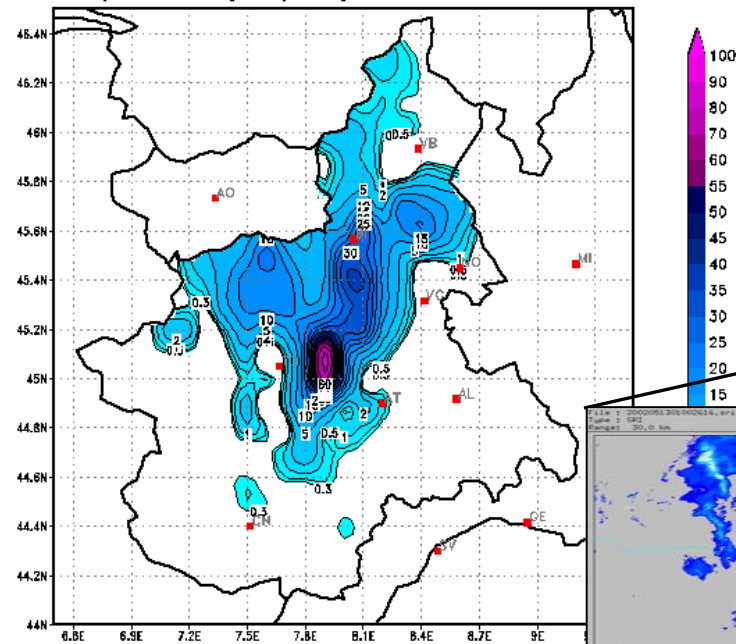
# RADAR PRODUCTS

## Comparison among instruments

Rain field spatially interpolated  
from rain gauges

Rain field derived from ARPAP  
C-band radar

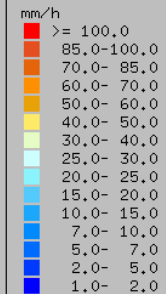
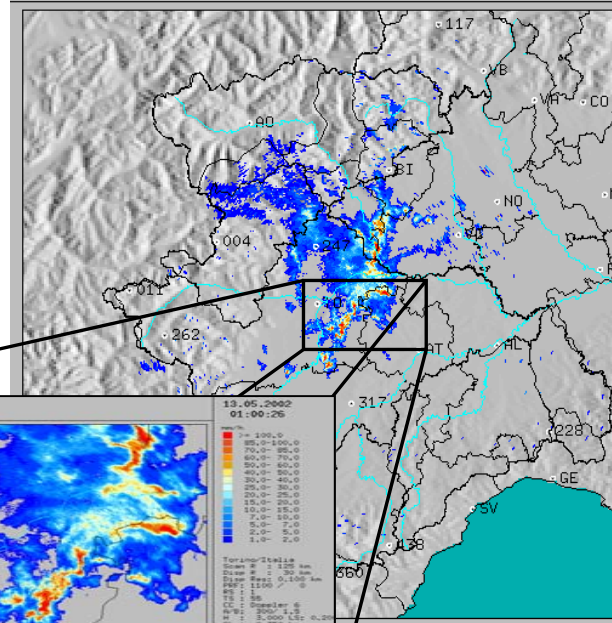
Precipitazione (mm/3hr) 13MAY2002 03:00 UTC



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Type : SRI  
Range : 125,0 km

Range = 125 km

13.05.2002  
01:00:26



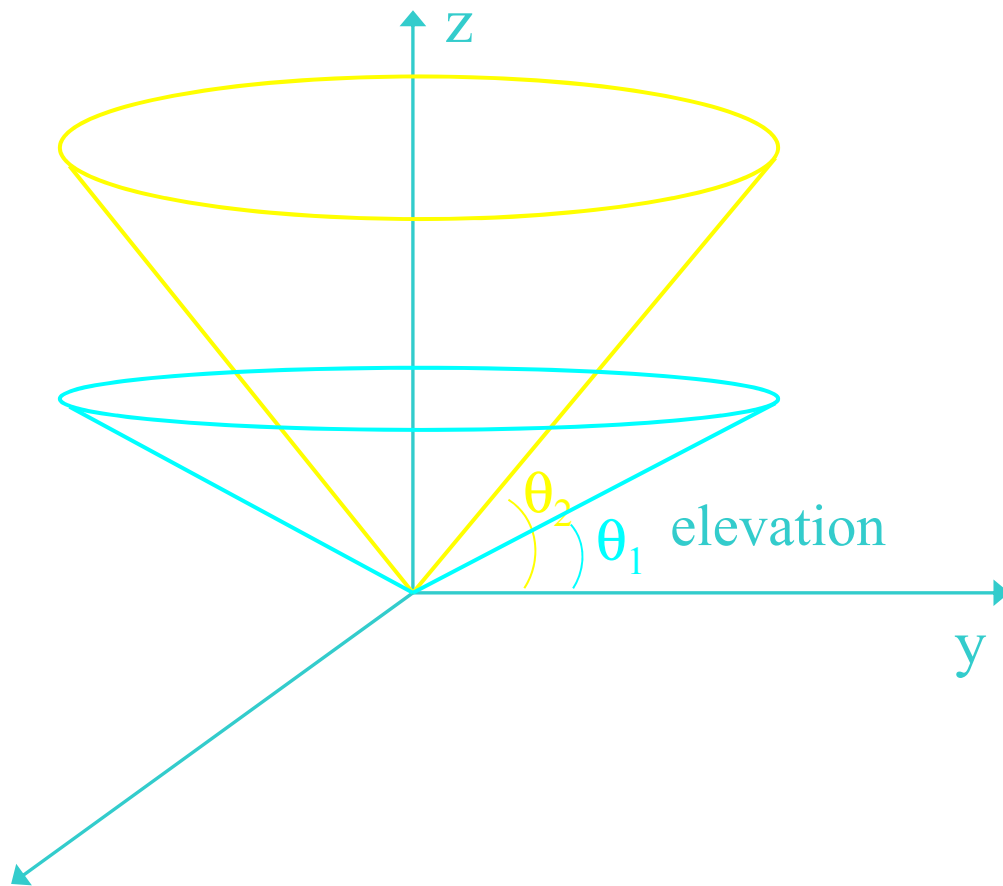
Torino/Italia  
Scan R : 125 km  
Disp R : 125 km  
Disp Res: 0,500 km  
PRF: 1100 / 0  
RS : 1  
TS : 55  
CC : Doppler 6  
A/B: 300/ 1,5  
H : 10,000 LS: 0,100  
SL : 1,000 km

REGIONE PIEMONTE  
Settore  
Meteoidrografico

Range = 30 km

# RADAR PRODUCTS

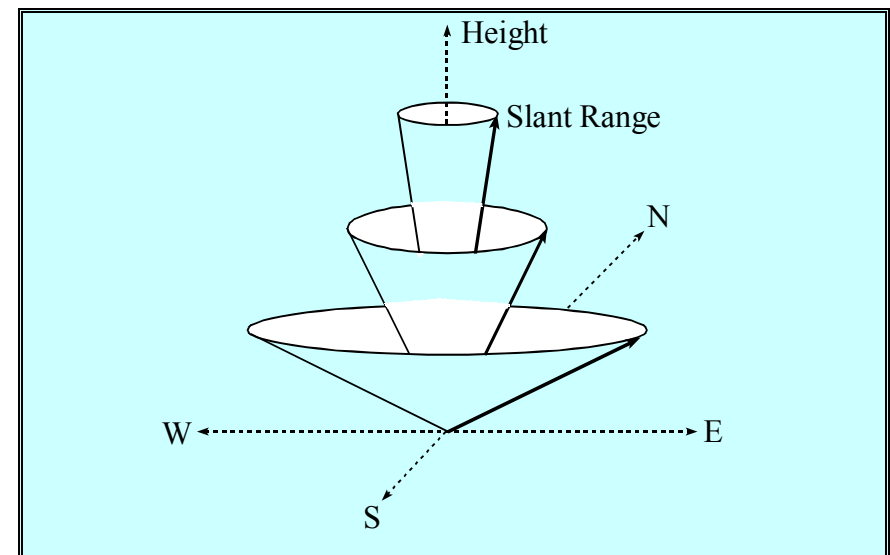
## Polar volume scanning



### SCANNING STRATEGY

Elevation angle  $\theta$ : 0-20 with variable step

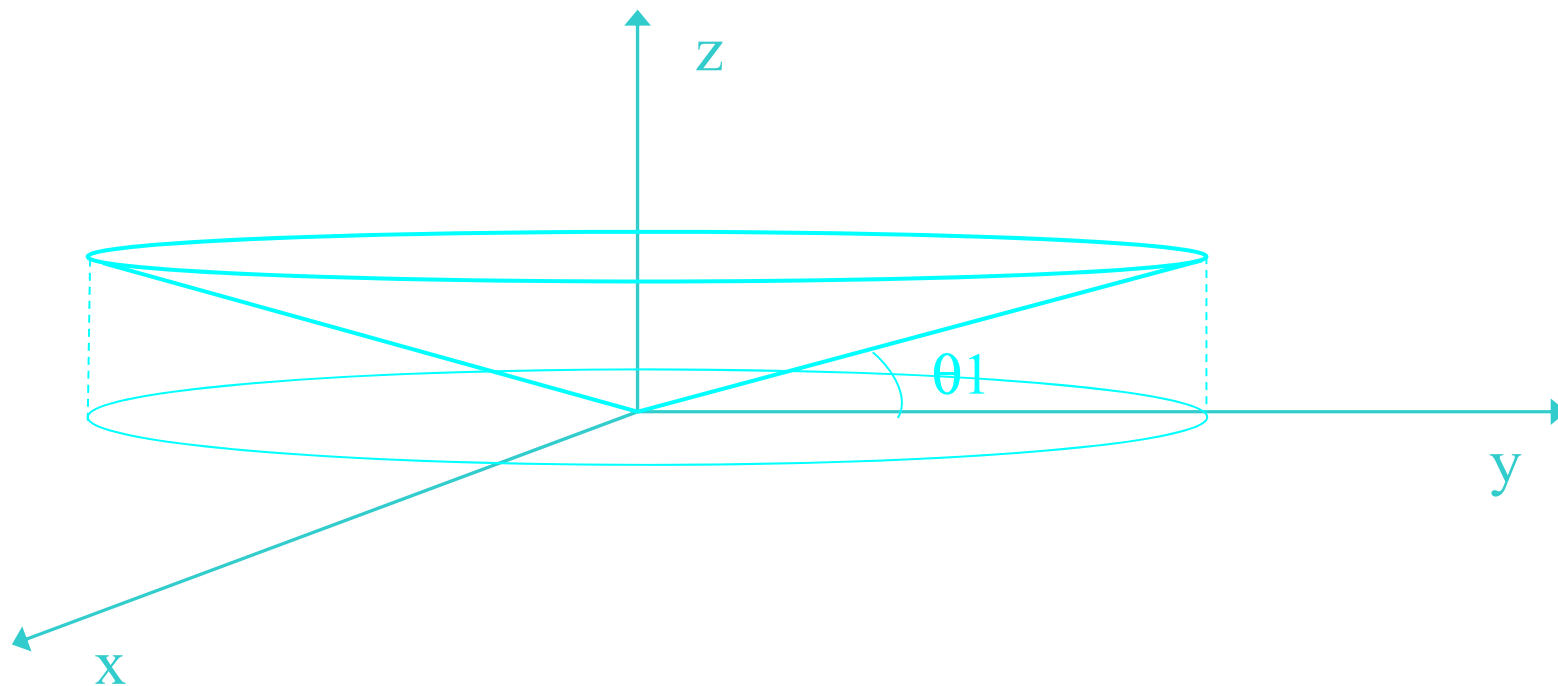
Azimuth angle  $\phi$ : 0-360° with 1° step



# RADAR PRODUCTS

## Plan Position Indicator (PPI)

Projection on the ground plane of radar 3-D polar data acquired upon a conical surface (constant elevation, azimuth between  $0^\circ$  and  $360^\circ$ )

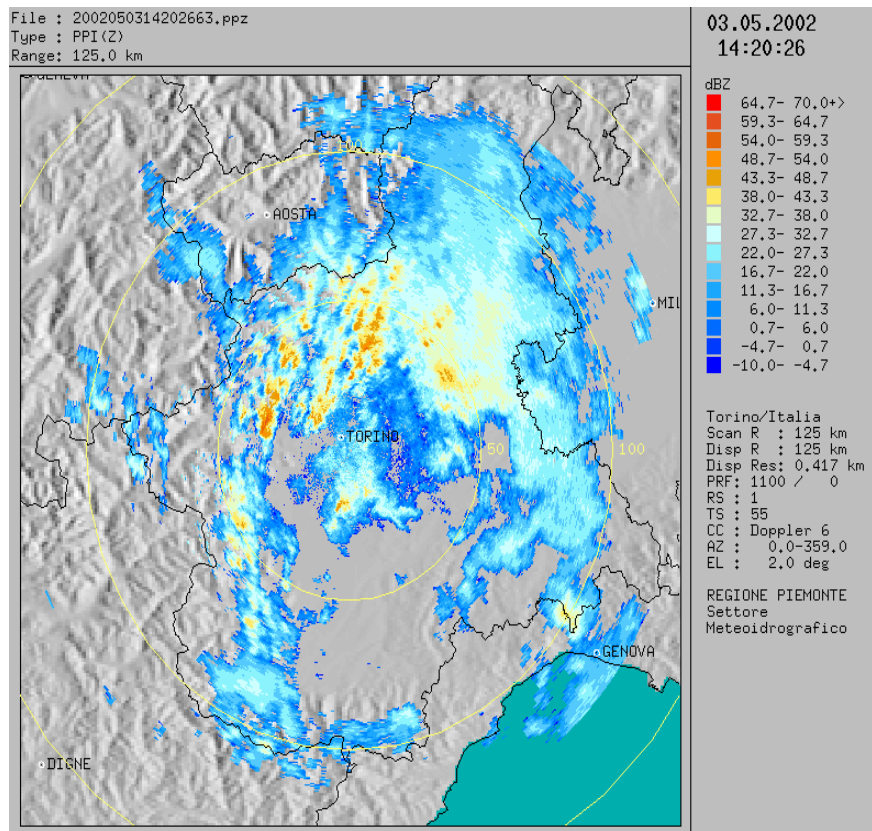




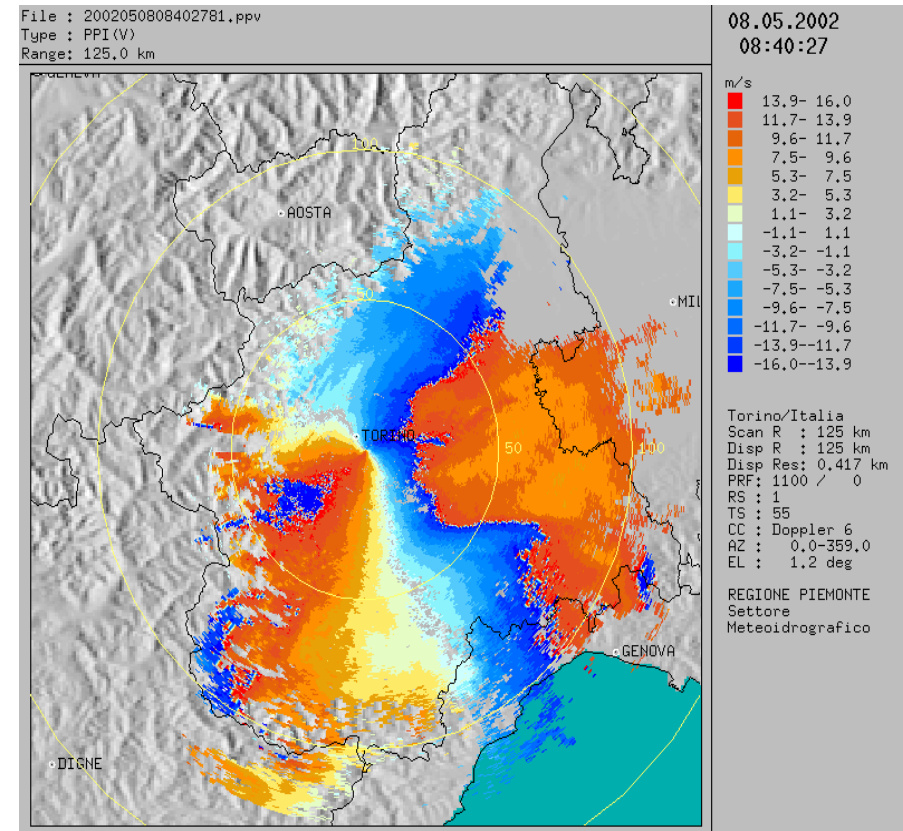
# RADAR PRODUCTS

## Example of PPI

### H-pol. Reflectivity at 2° elevation



### Radial velocity at 2° elevation

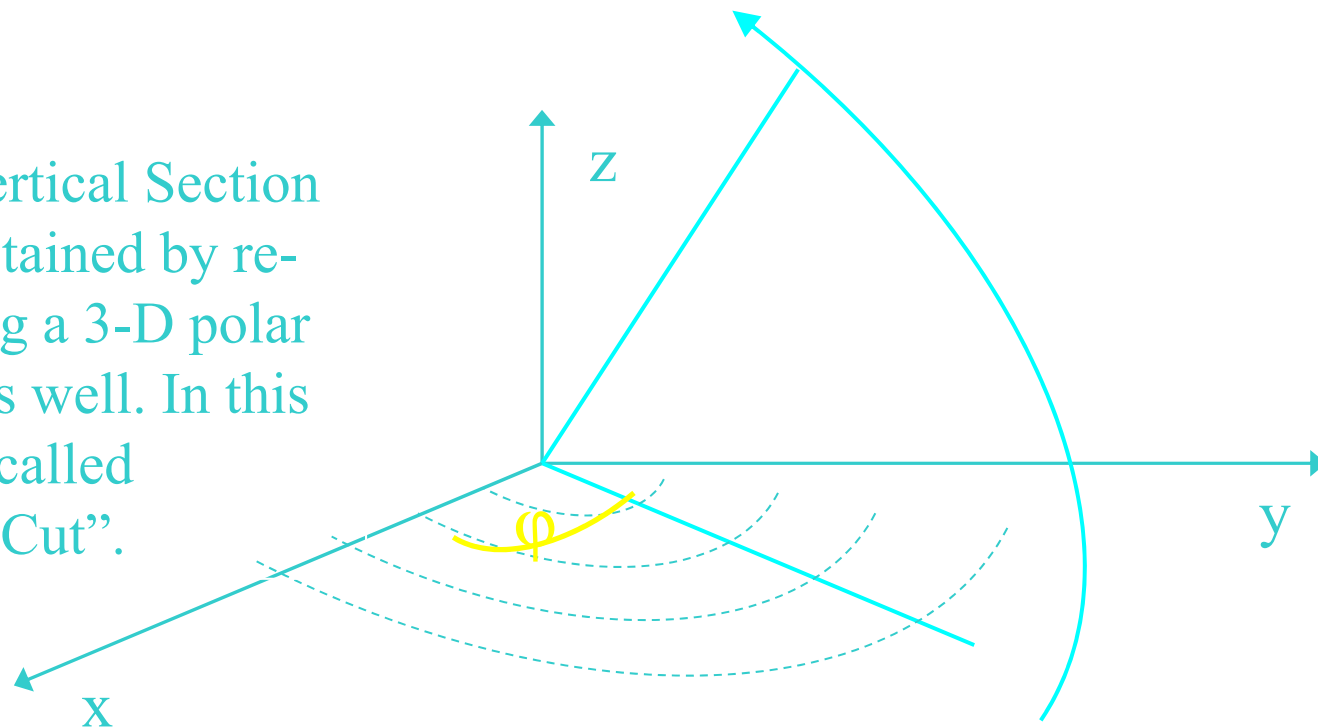


# RADAR PRODUCTS

## Range Height Indicator (RHI)

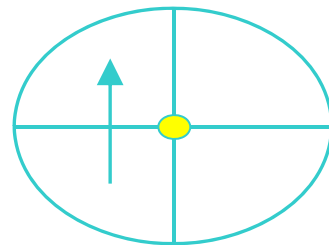
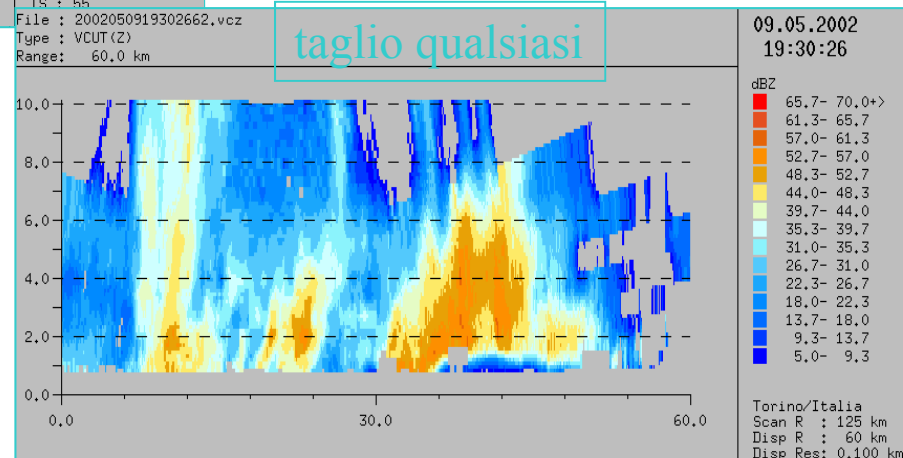
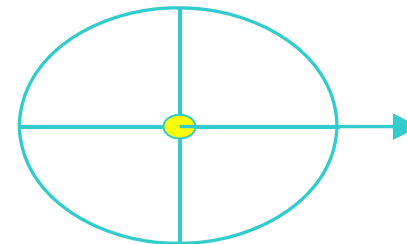
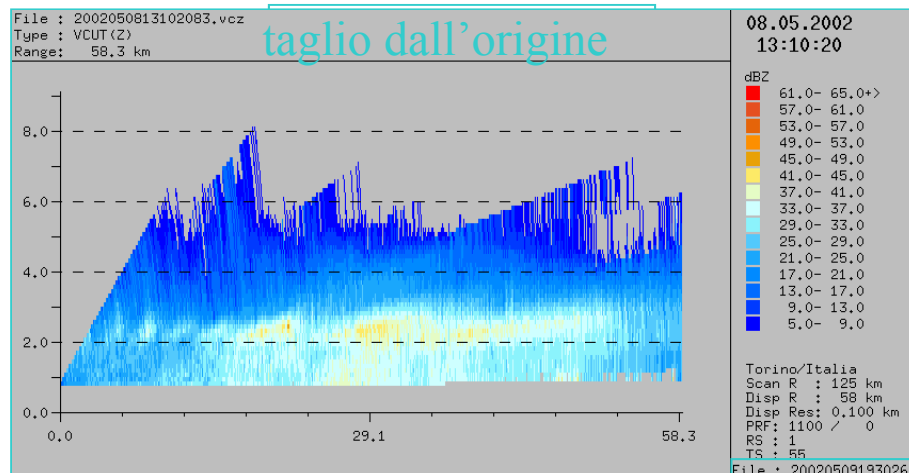
RHI: Vertical scan for constant azimuth angle  $\phi$

N.B. A vertical Section can be obtained by re-processing a 3-D polar volume as well. In this case it is called “Vertical Cut”.



# RADAR PRODUCTS

## Example of RHI and VCut

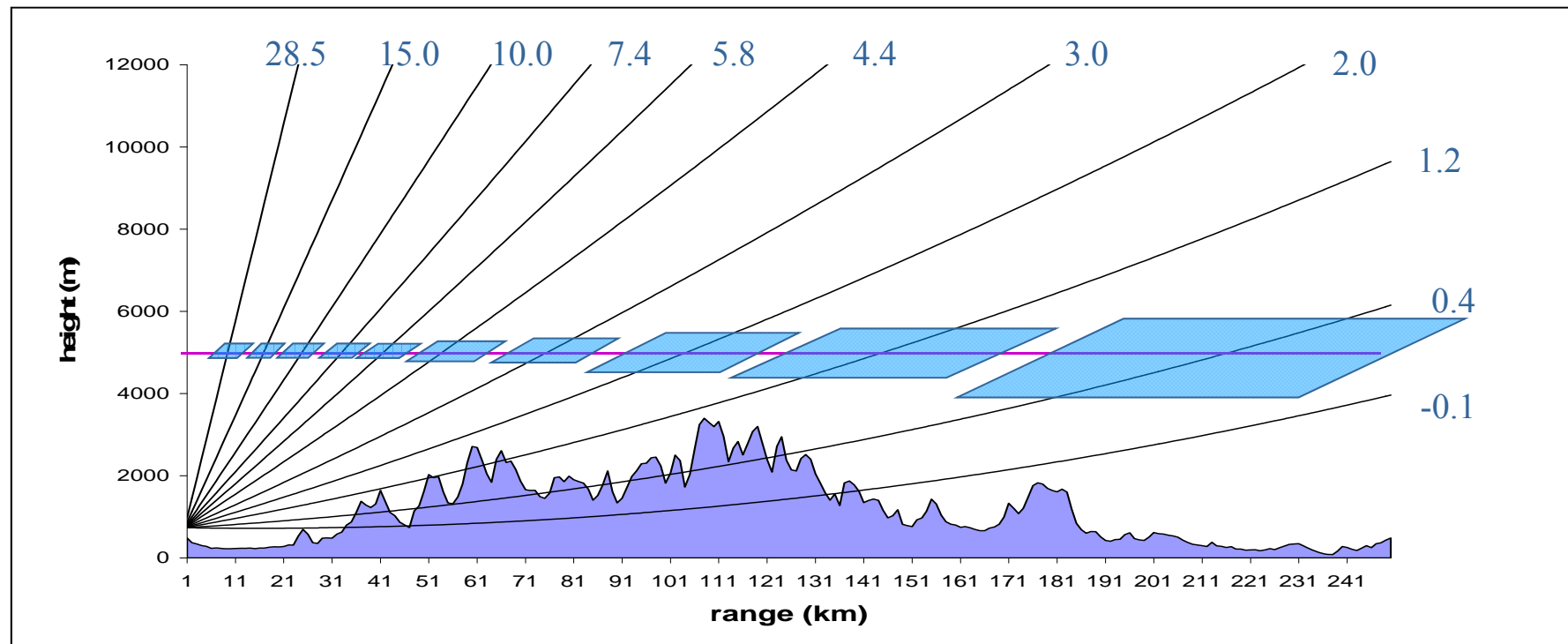




# RADAR PRODUCTS

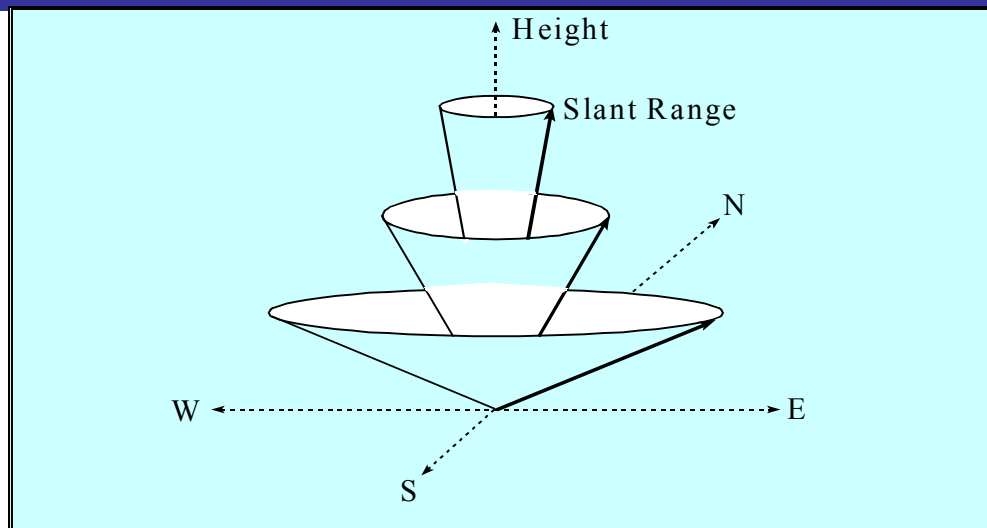
## Constant Altitude PPI

CAPPI: Horizontal cut of polar volume at a constant altitude

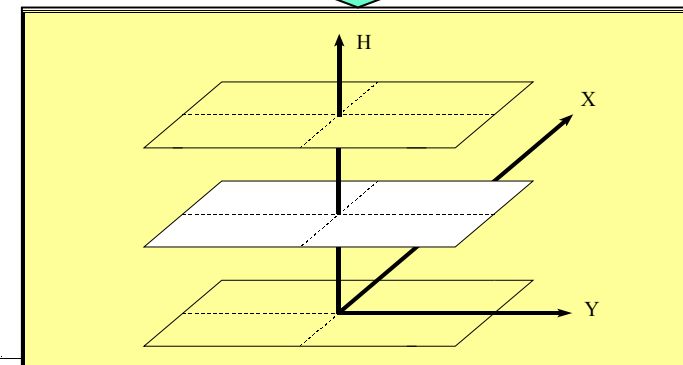
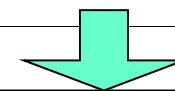
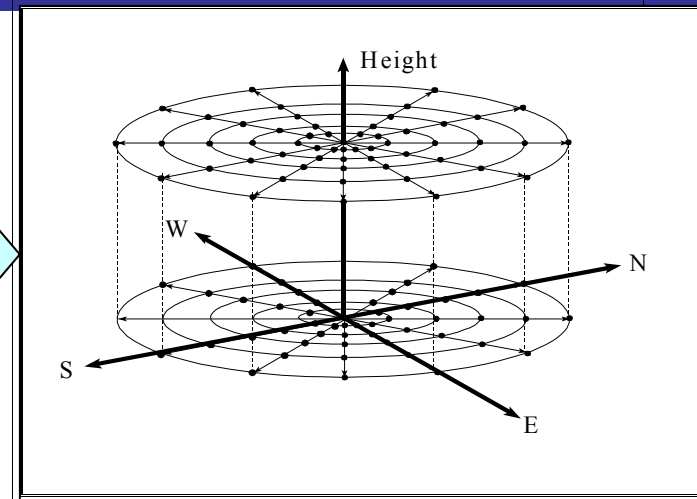
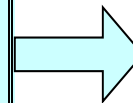


# RADAR PRODUCTS

## CAPPI processing



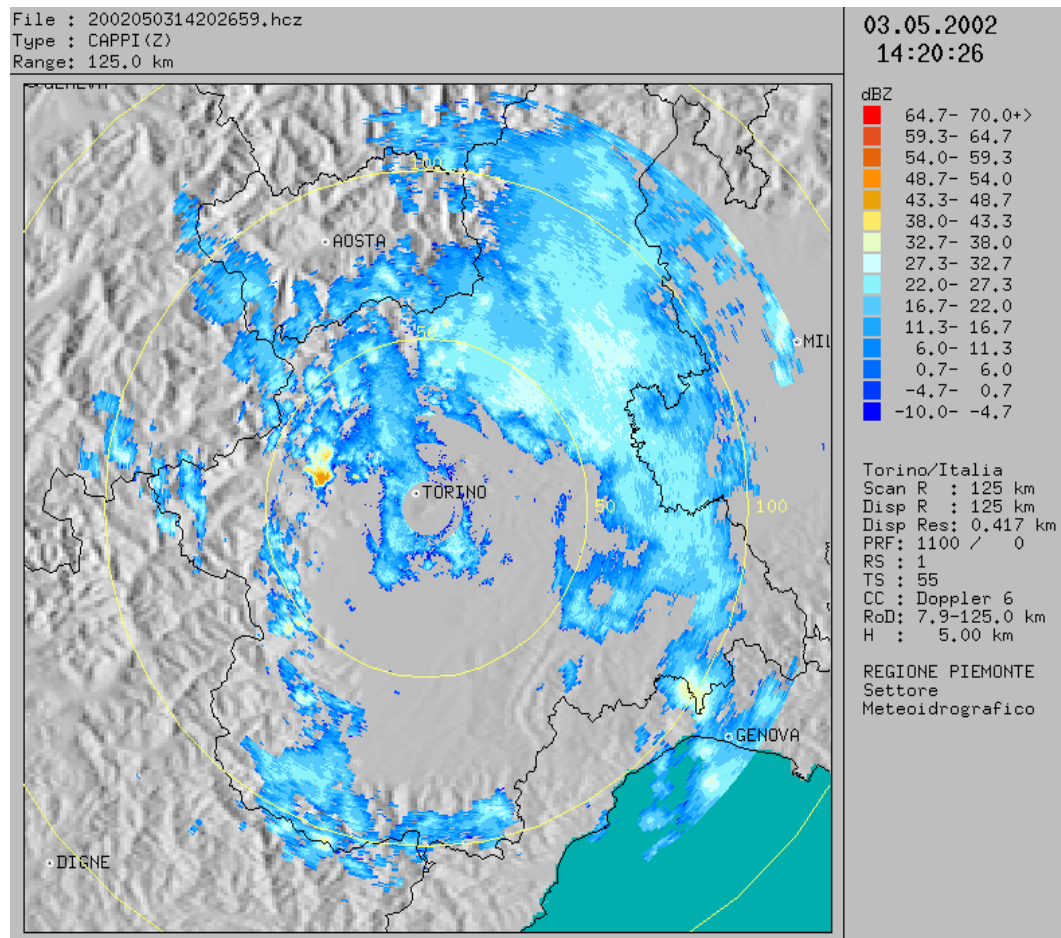
Volume polare



Volume  
cartesiano

# RADAR PRODUCTS

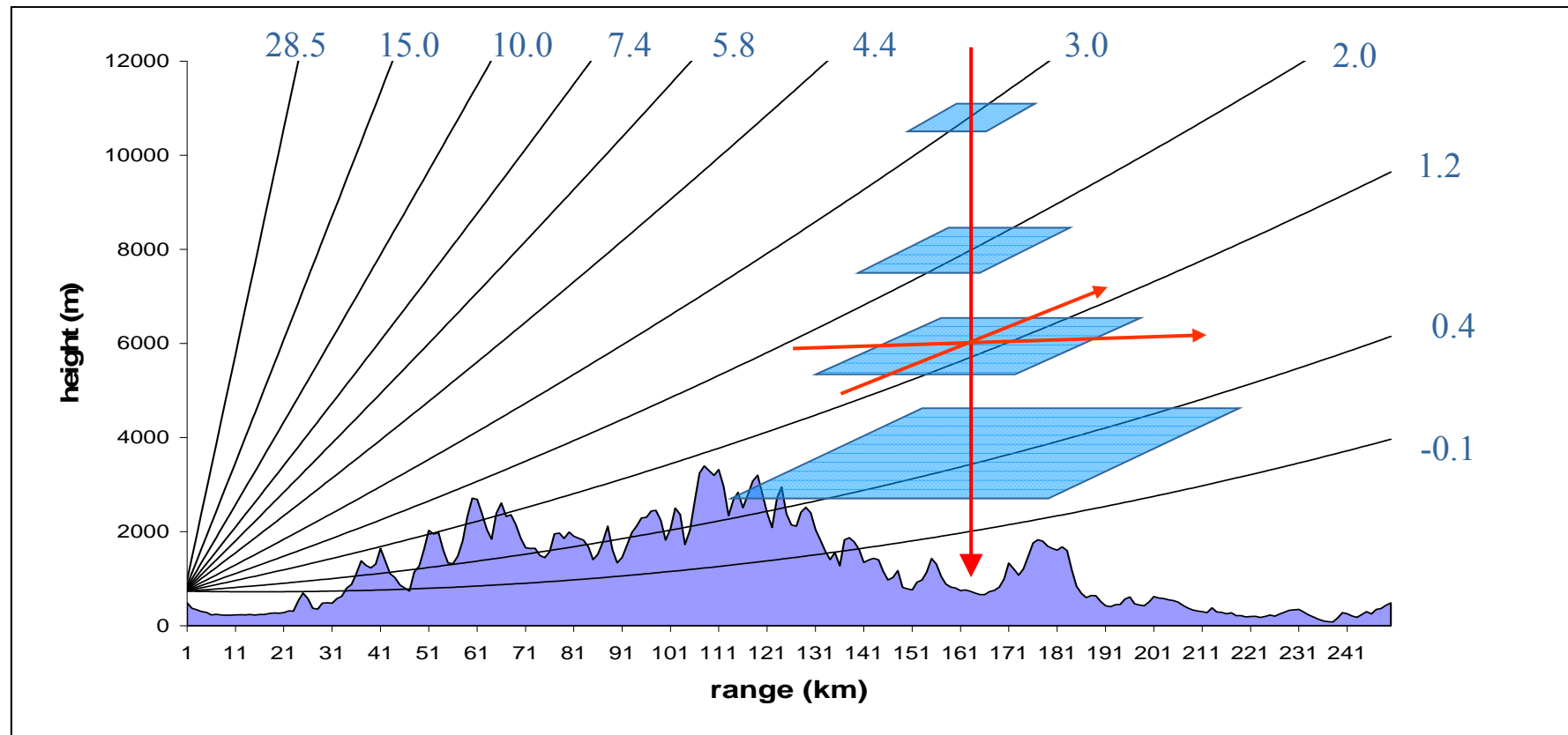
## Example of CAPPI



# RADAR PRODUCTS

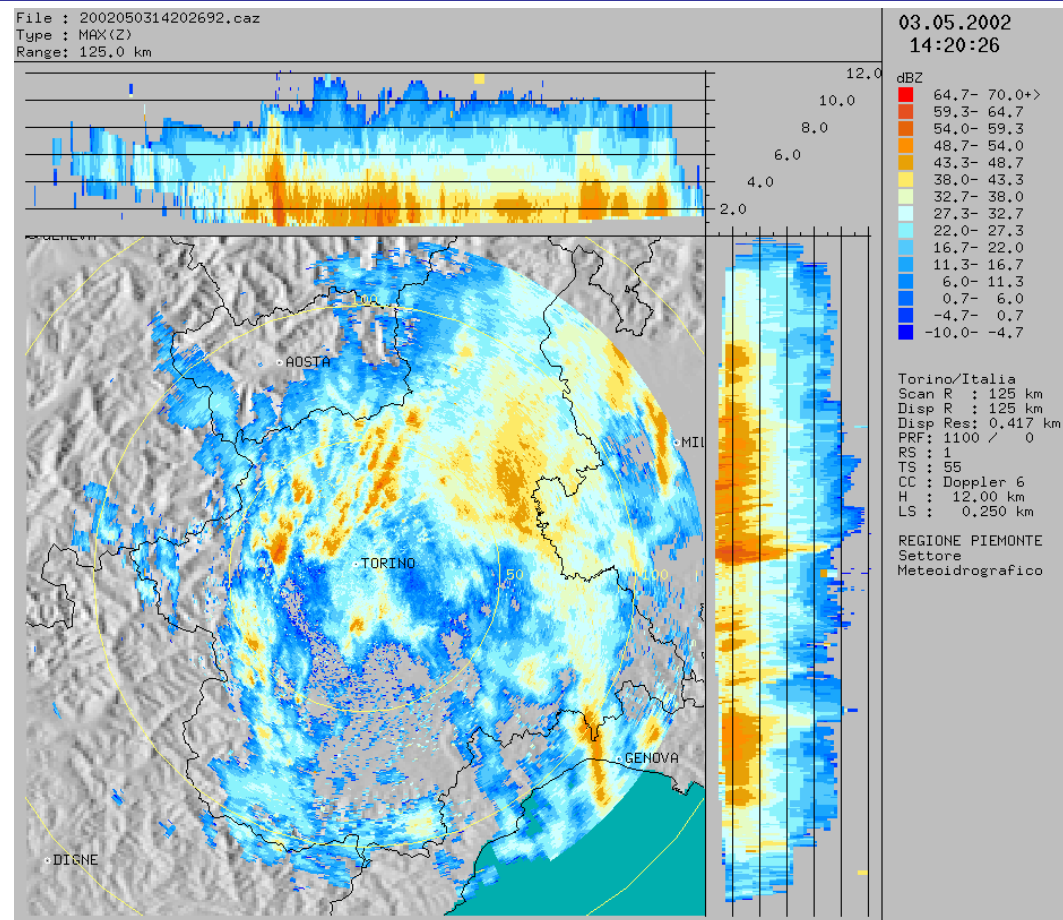
## Hor.-Vertical Maximum Intensity

HVMI: Horizontal-vertical projection (3 axes) of maximum reflectivity



# RADAR PRODUCTS

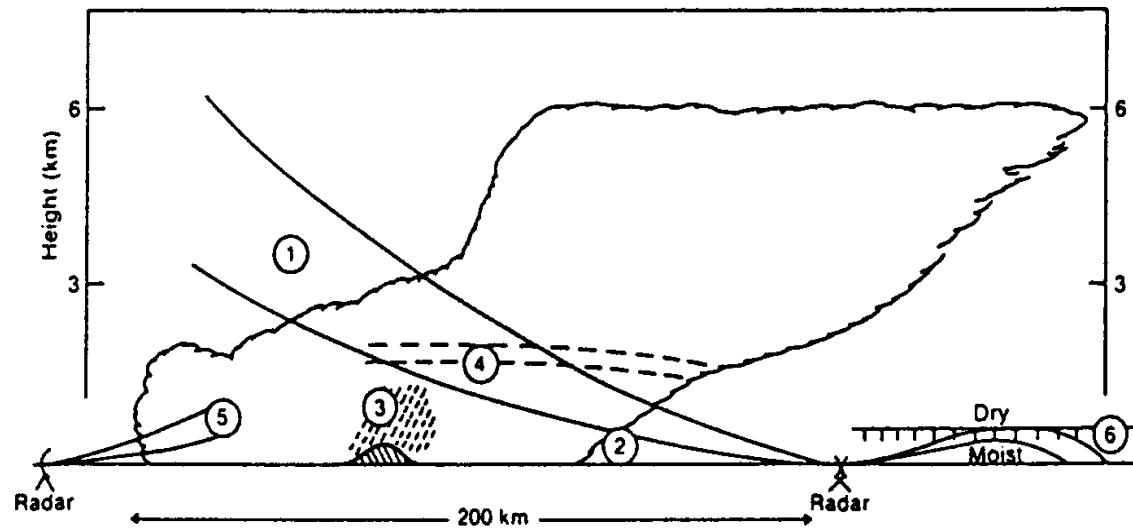
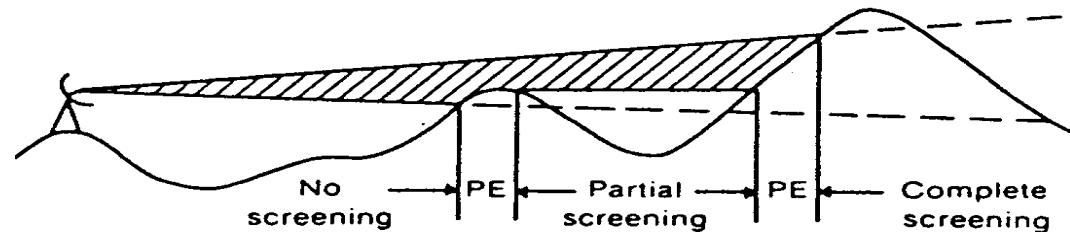
## Example of VMI



# RADAR PRODUCTS

## Errors in rainfall retrieval

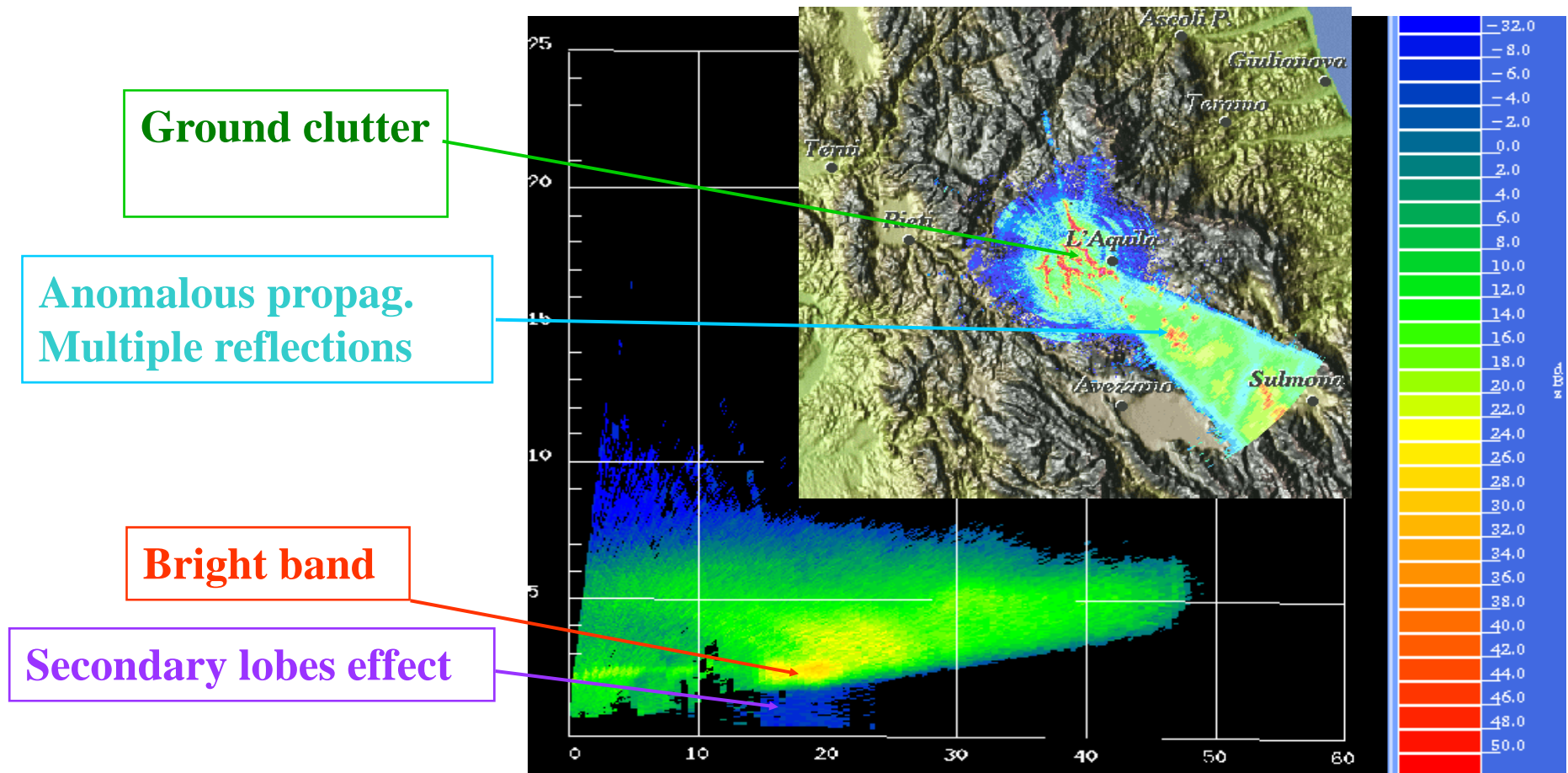
1. Inhomogeneity of rain cloud within the radar beam
2. Evaporation of rainfall below the radar beam
3. Underestimation due to orographic effects on rainfall process
4. Bright-band effects due to melting layer
5. Path attenuation due clear-air gases (water vapor and oxygen) and rain
6. Anomalous propagation due to air refractivity – second-trip echoes
7. Beam screening due to presence of obstacles (hills, mountains)
8. Calibration errors within the radar system





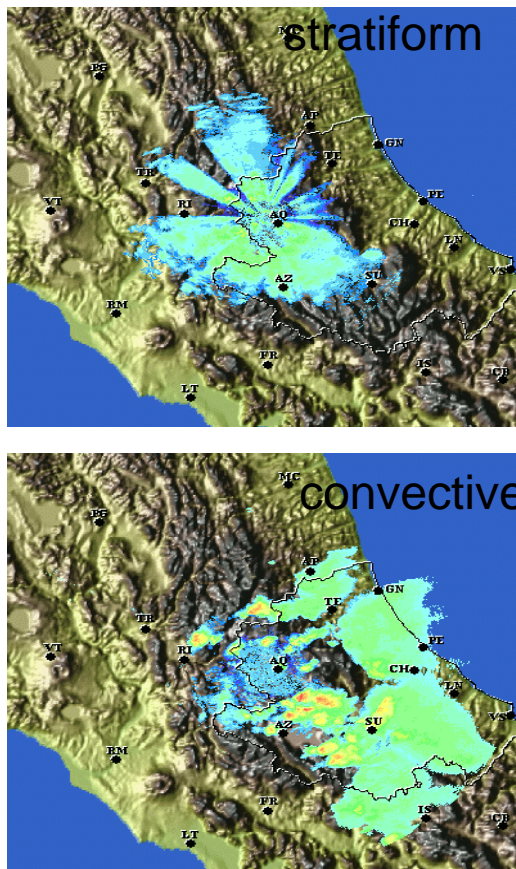
# RADAR PRODUCTS

## Examples of radar artifacts (1)

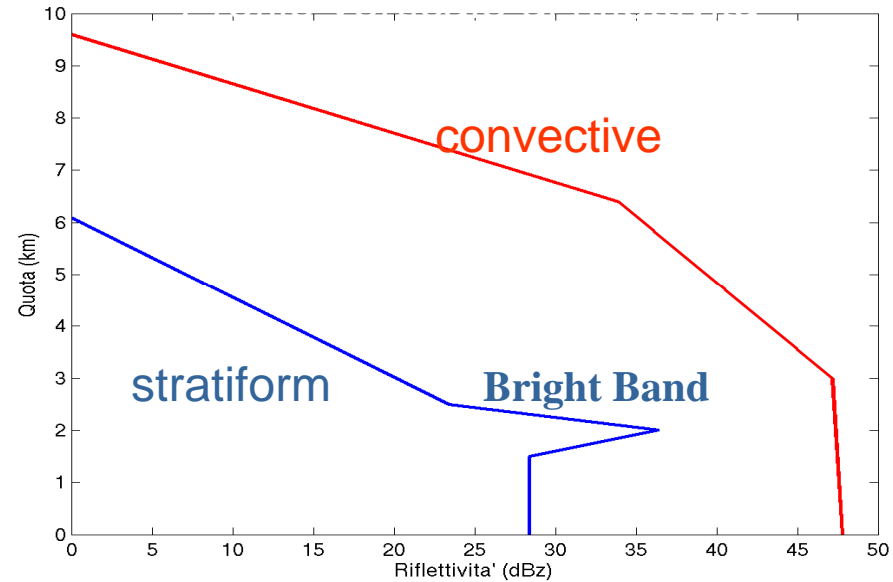


# RADAR PRODUCTS

## Examples of radar artifacts (2)



Estimate of surface rainfall from stratiform or convective Z-R relation ?

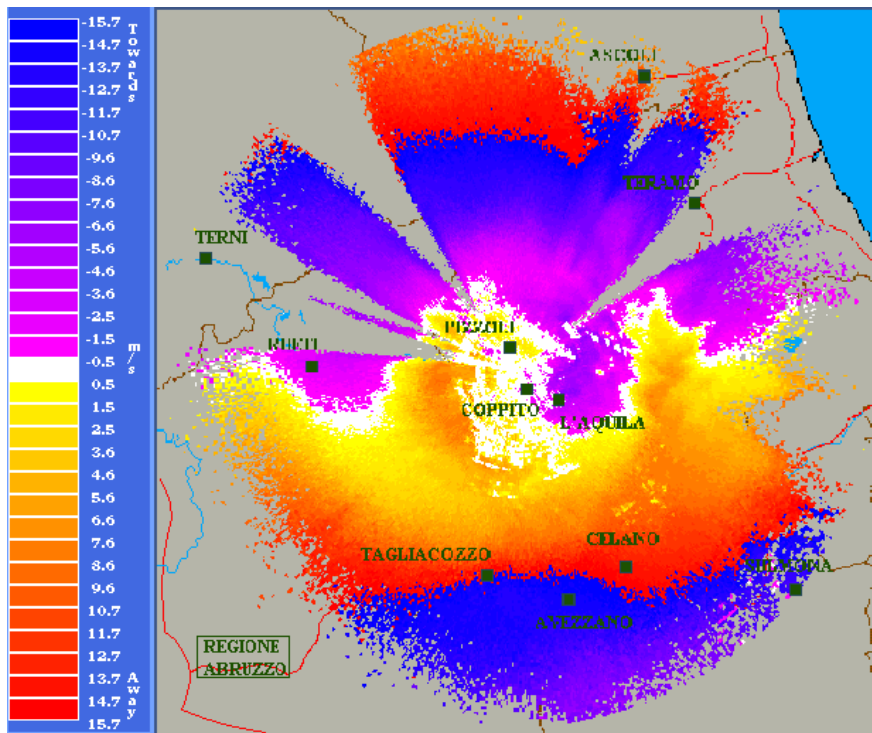




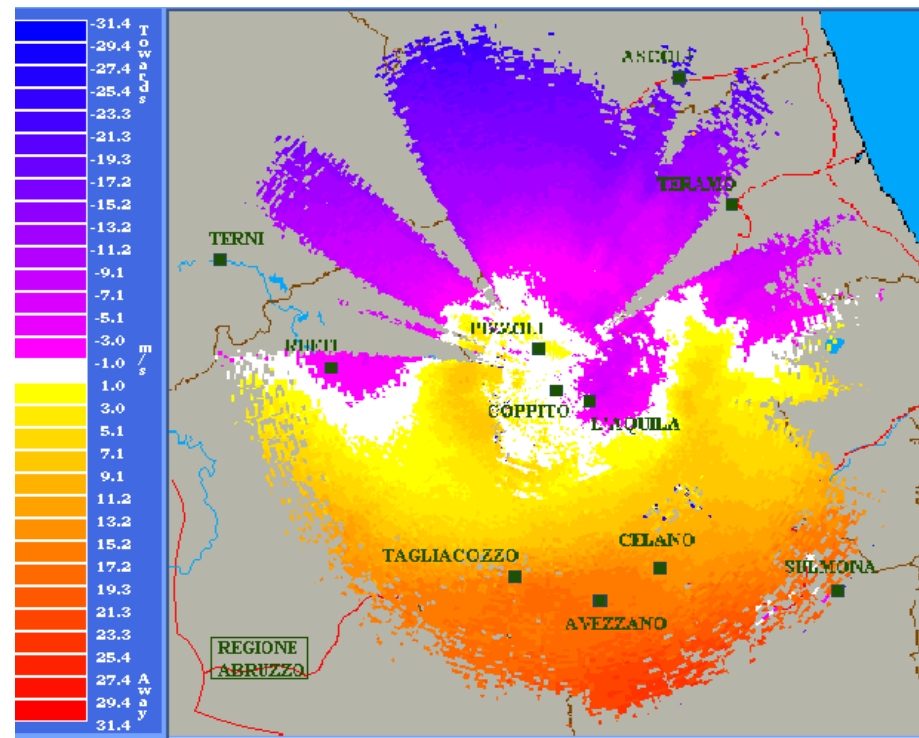
# RADAR PRODUCTS

## Examples of radar artifacts (3)

Max. unamb. velocity:  $u_{\max} = \pm 16\text{m/s}$   
Mode: single PRF=1180 Hz



Max. unamb. velocity:  $u_{\max} = \pm 32\text{m/s}$   
Mode: dual PRF=1180 Hz, 787 Hz

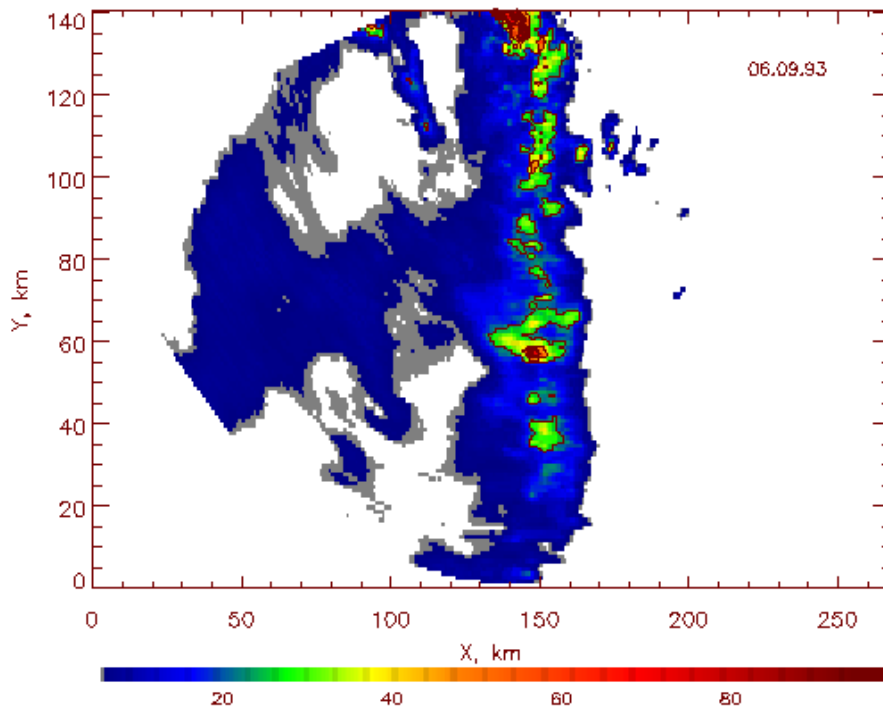


# RADAR PRODUCTS

## Examples of radar artifacts (4)

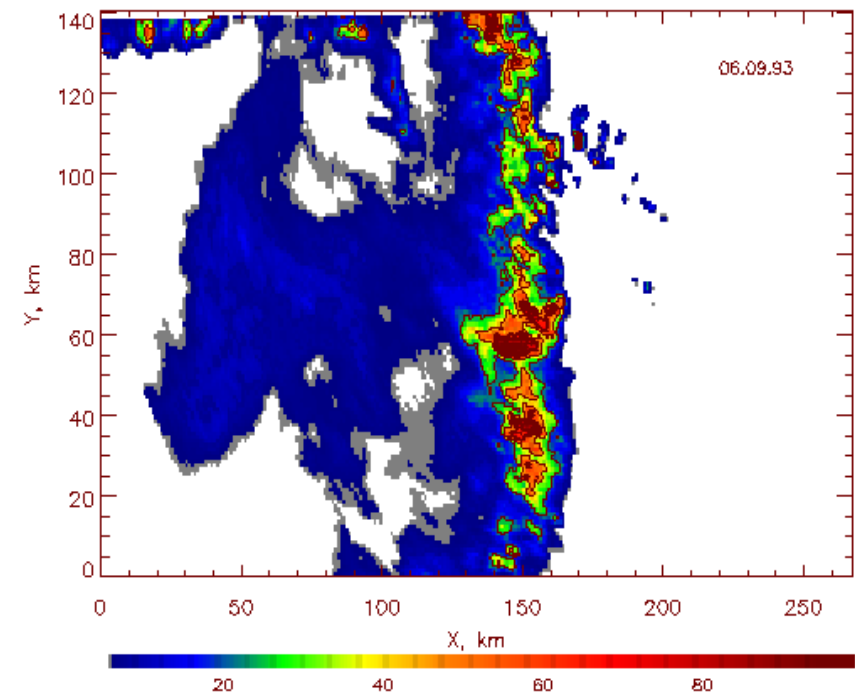
**C-band radar:** affected by rain path attenuation

R(Z) – Cimarron



**S-band radar:** immune to rain path attenuation

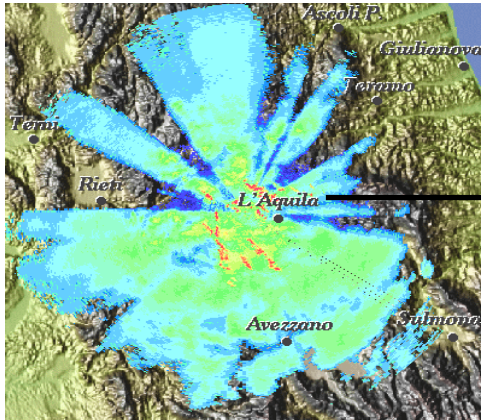
R(Z) – WSR-88D



# RADAR PRODUCTS

## Radar data processing

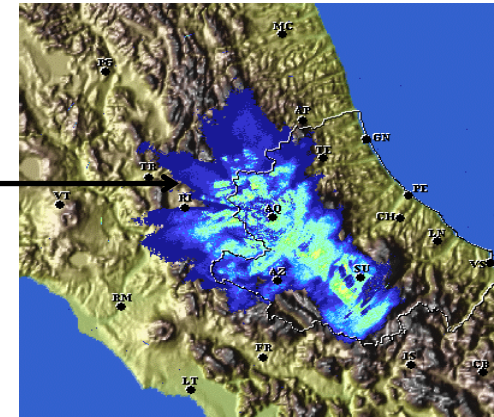
radar volume



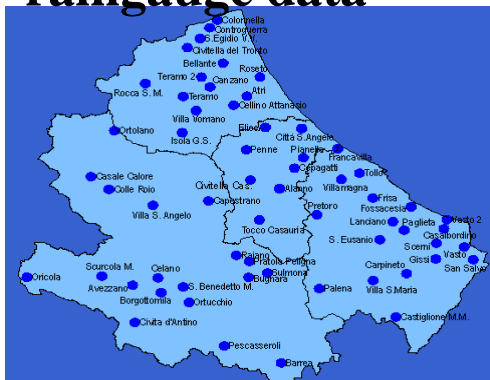
Clutter, ana-prop., vel. correction

VPR, atten. correction

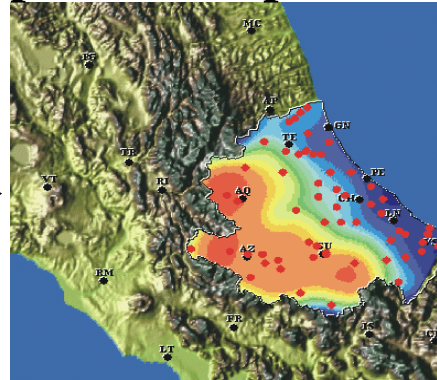
rain rate estimation



raingauge data



spatial interpolation



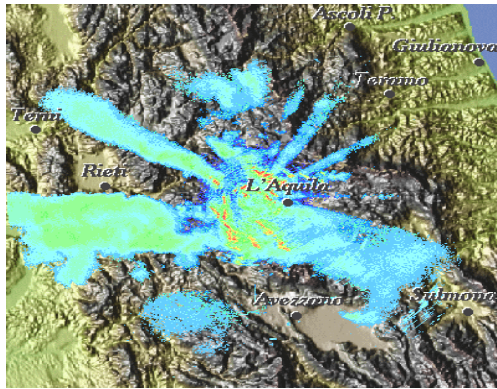
Error Estimation



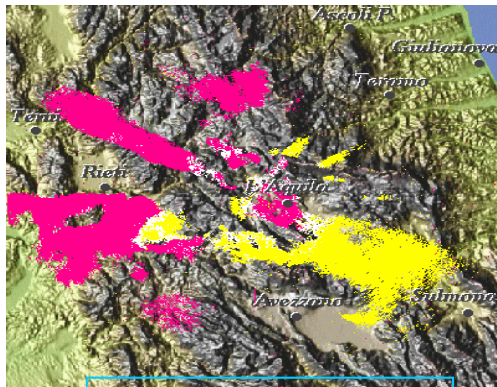
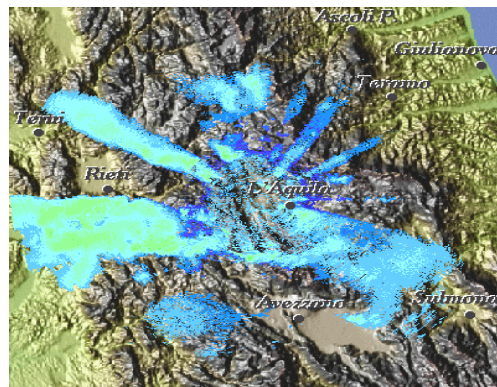
# RADAR PRODUCTS

## Example of data processing

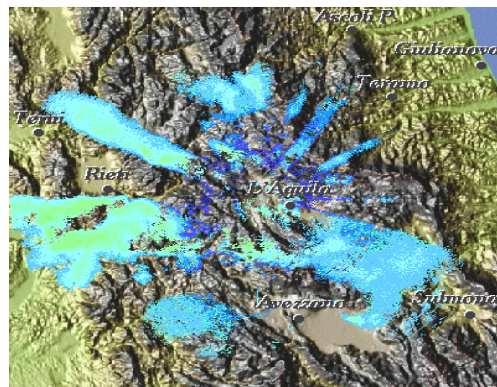
Z uncorrected



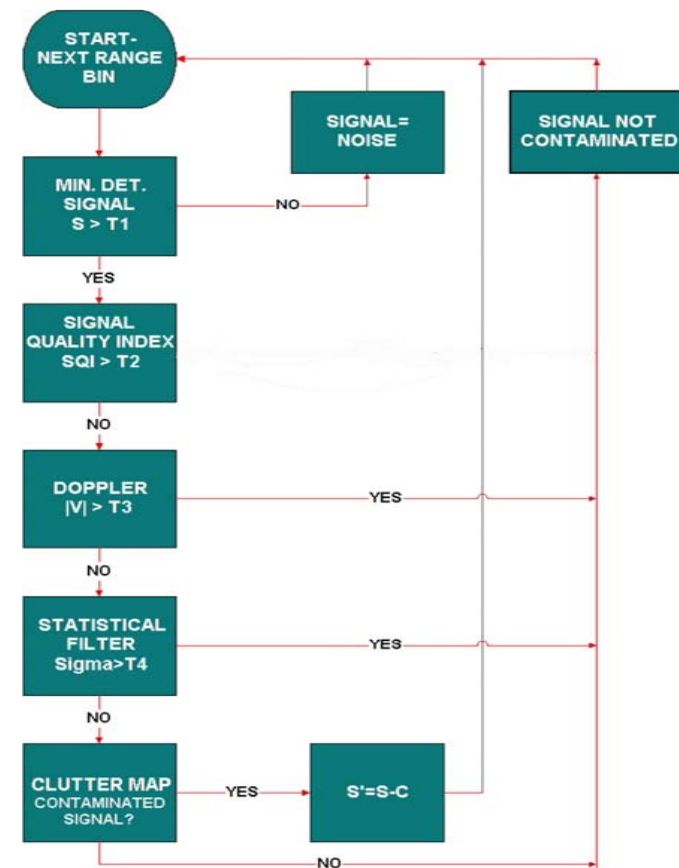
Z corrected



Radial velocity

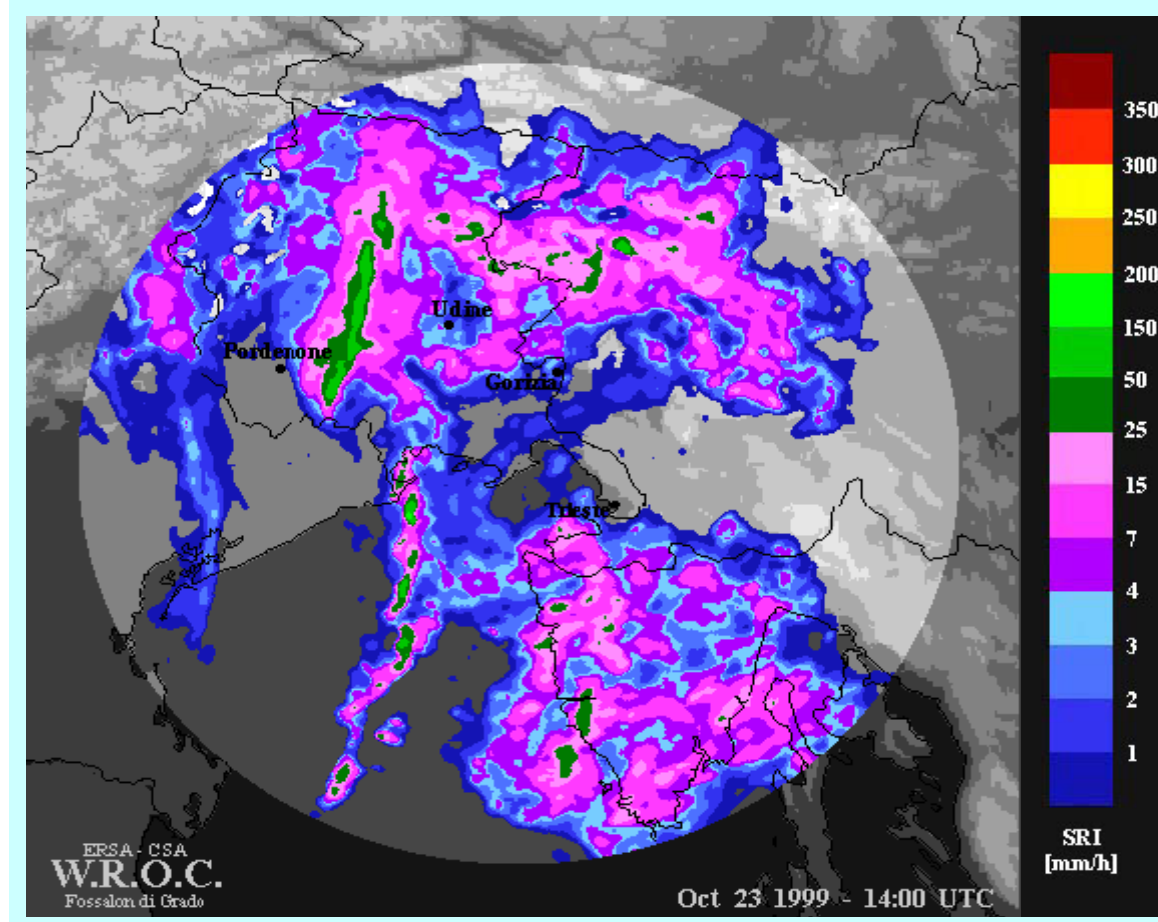


Spectral filter



# RADAR PRODUCTS

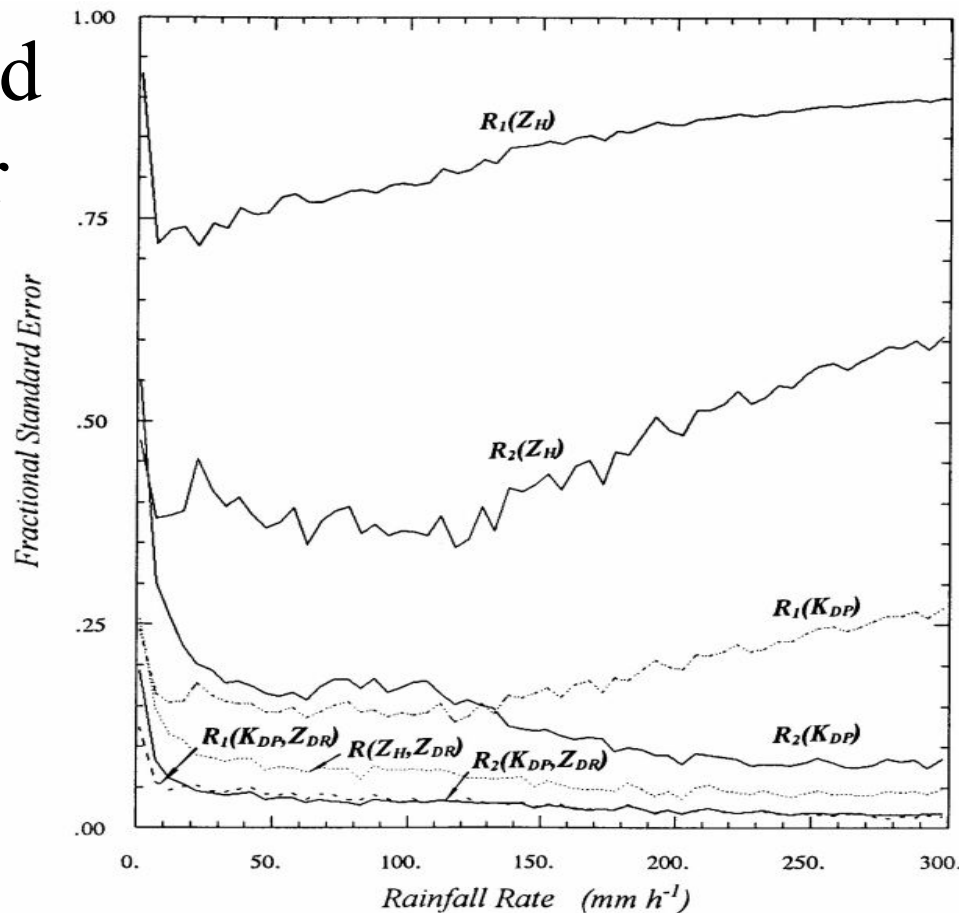
## Surface rainfall C-band retrieval



# RADAR PRODUCTS

## Retrieval expected accuracy

S-band  
radar



*Fractional Standard Error*

$$FSE = \frac{\left\langle \left[ \hat{R}_{est} - R_{sim} \right]^2 \right\rangle^{1/2}}{\langle R_{sim} \rangle}$$

*Algorithms:*

R<sub>1</sub>=M-P

R<sub>2</sub>=WSR-88D

R<sub>dp</sub>=G. et al. (rob)

R<sub>1</sub>(K<sub>dp</sub>)=S-Z

R<sub>2</sub>(K<sub>dp</sub>)=G-S (lin)

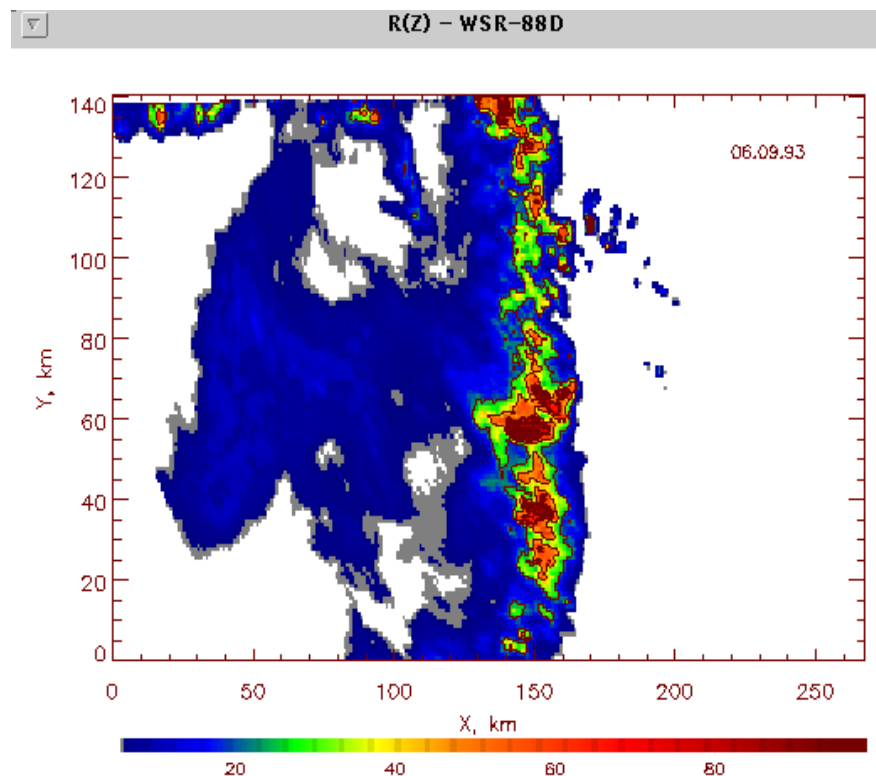
R<sub>dp1</sub>=G-S (rob)

R<sub>dp2</sub>=R-Z (log)

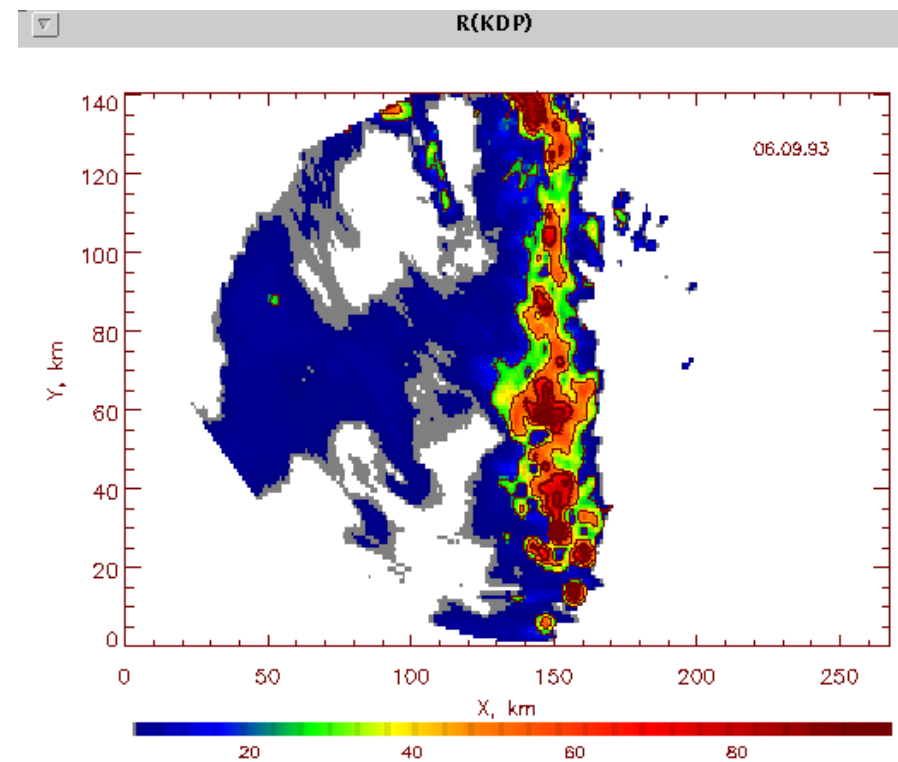
# RADAR PRODUCTS

## Path attenuation mitigation

**S-band radar:** immune to rain path attenuation



**C-band radar:** corrected for rain path attenuation using **fully polarimetric data** ( $Z_h$ ,  $Z_{dr}$ ,  $K_{dp}$ )

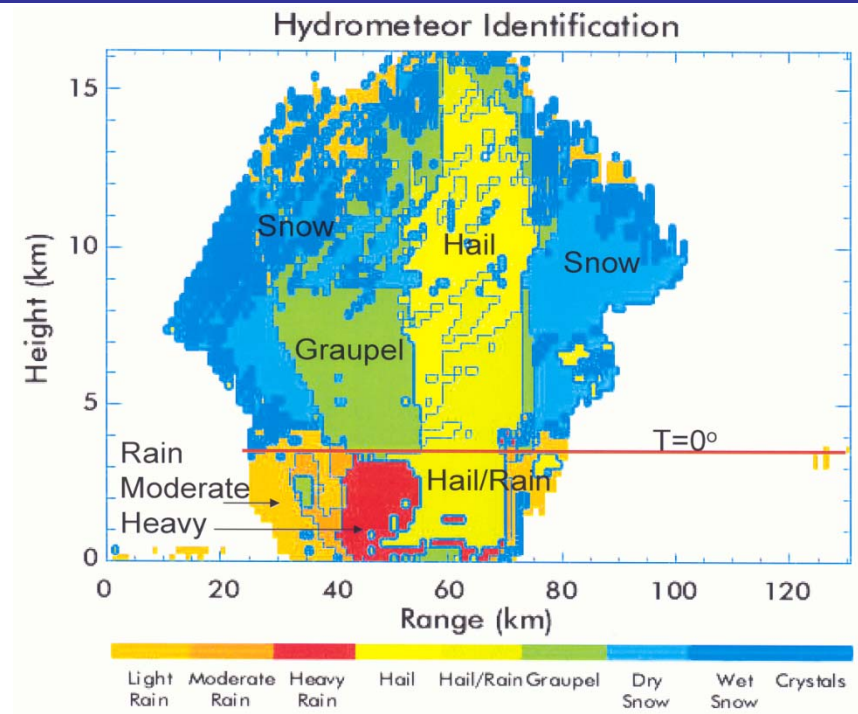
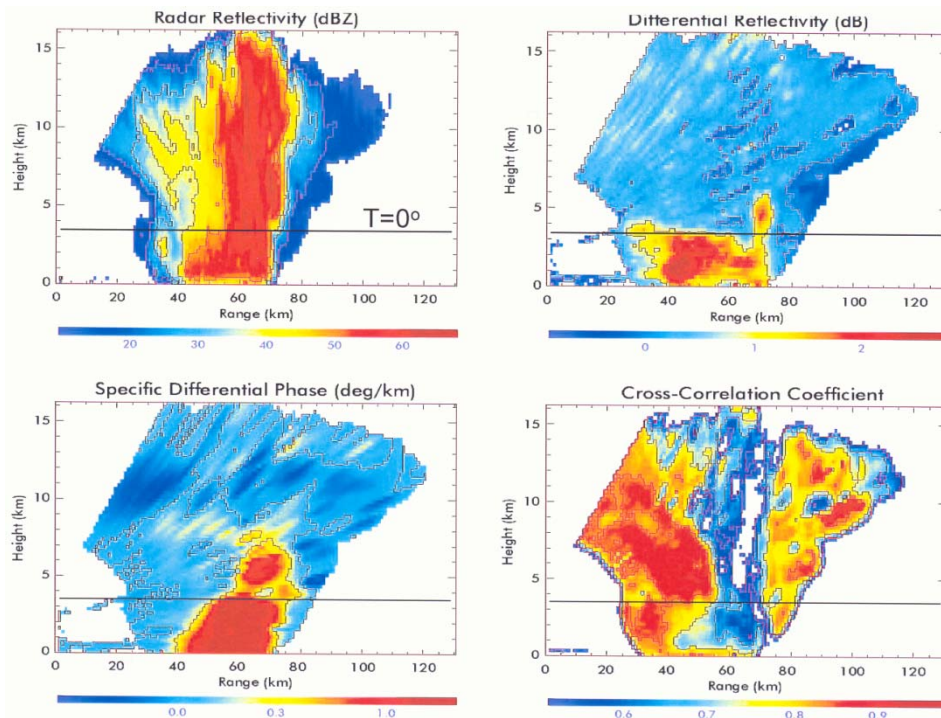




# RADAR PRODUCTS

## Hydrometeor classification

### Hailstorm in Florida observed by S-band polarimetric radar



... and derived hydrometeor classification